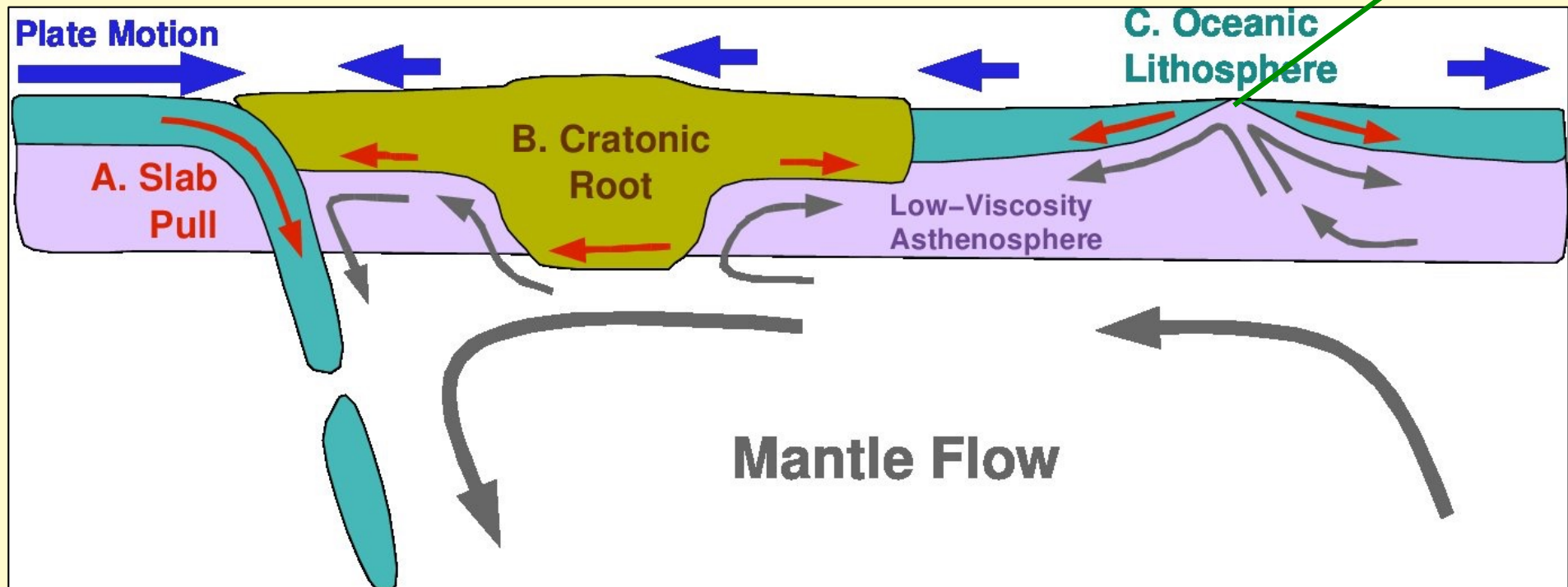


Lithosphere and Asthenosphere: Composition and Evolution

GEO-DEEP9300

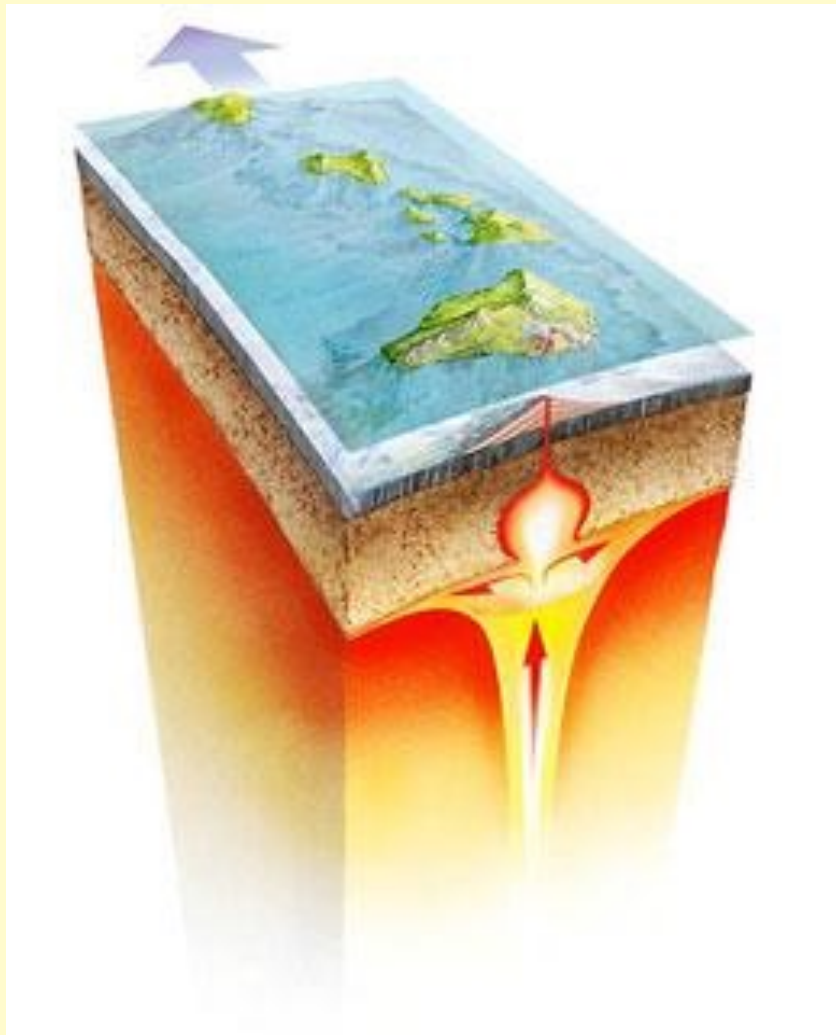
Valerie Maupin
Clint Conrad

Volcanism

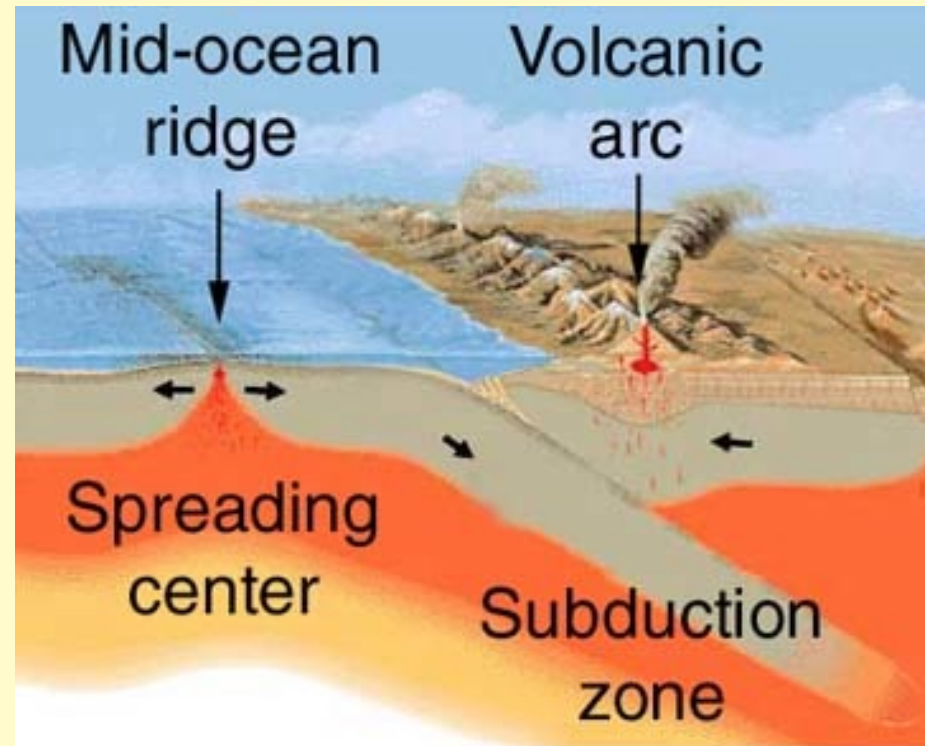


Global Volcanism

Most volcanism results from plumes bringing heat from the deep mantle ...



... or by “shallow” processes at plate boundaries



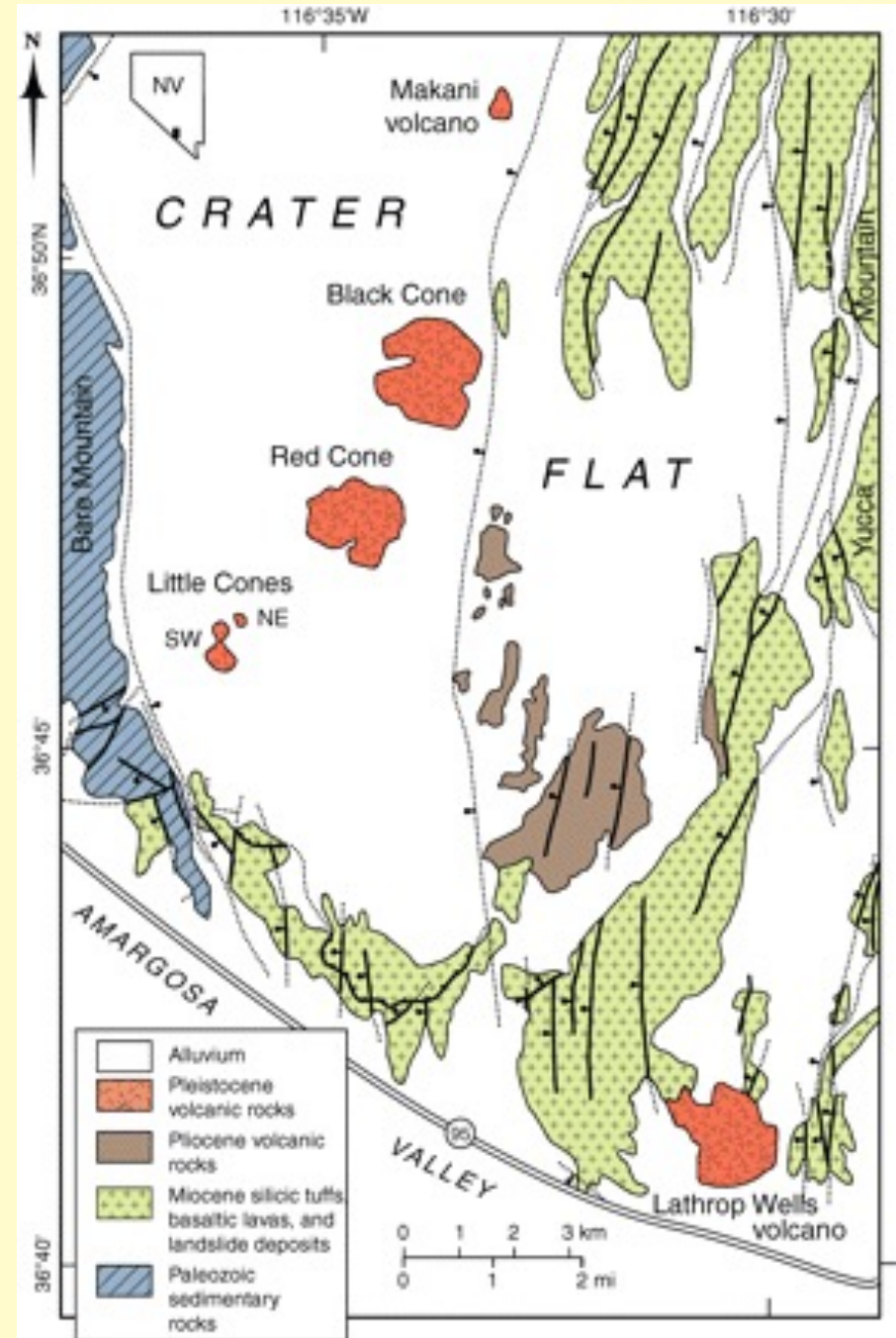
Crater Flat, Southern Nevada



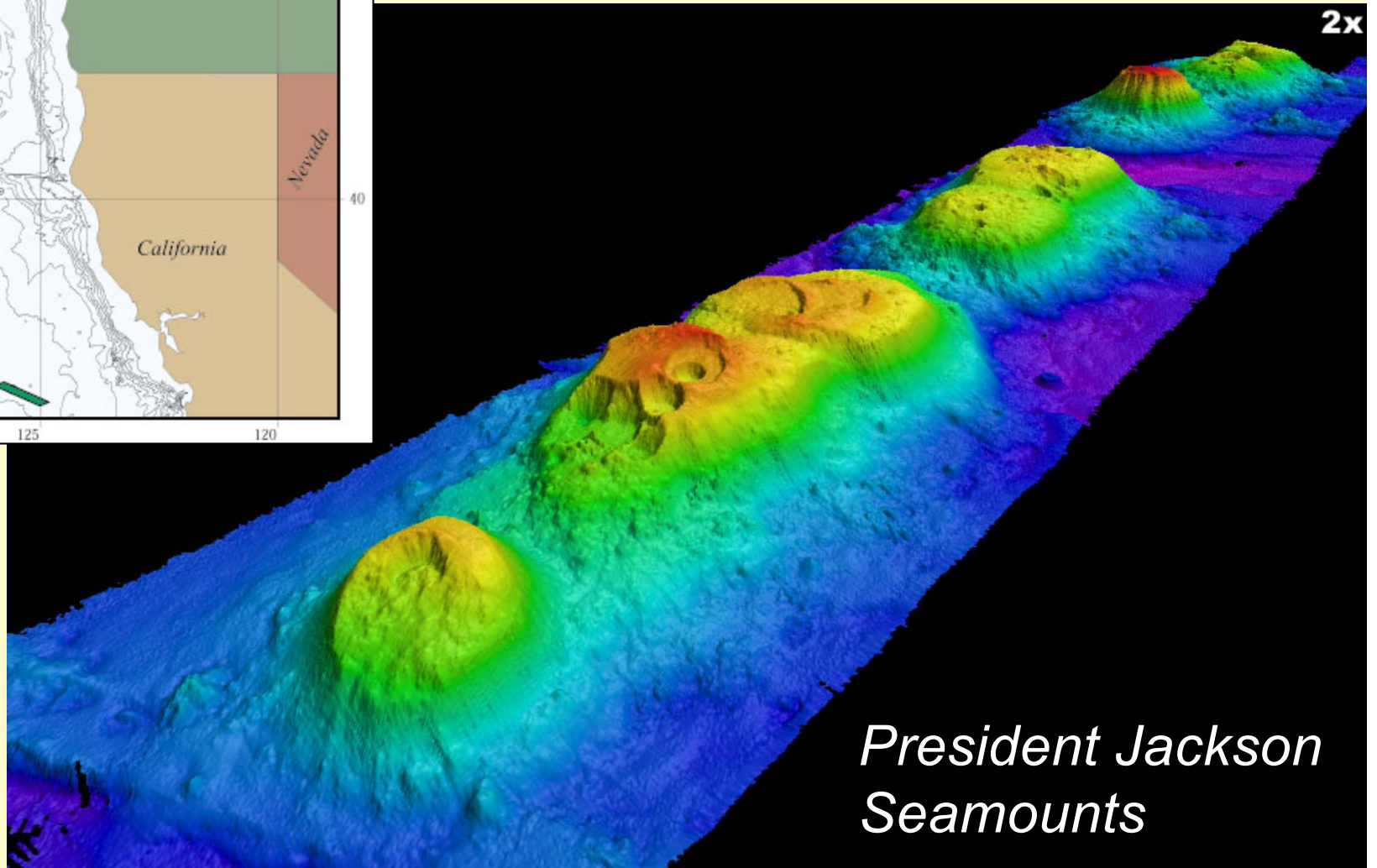
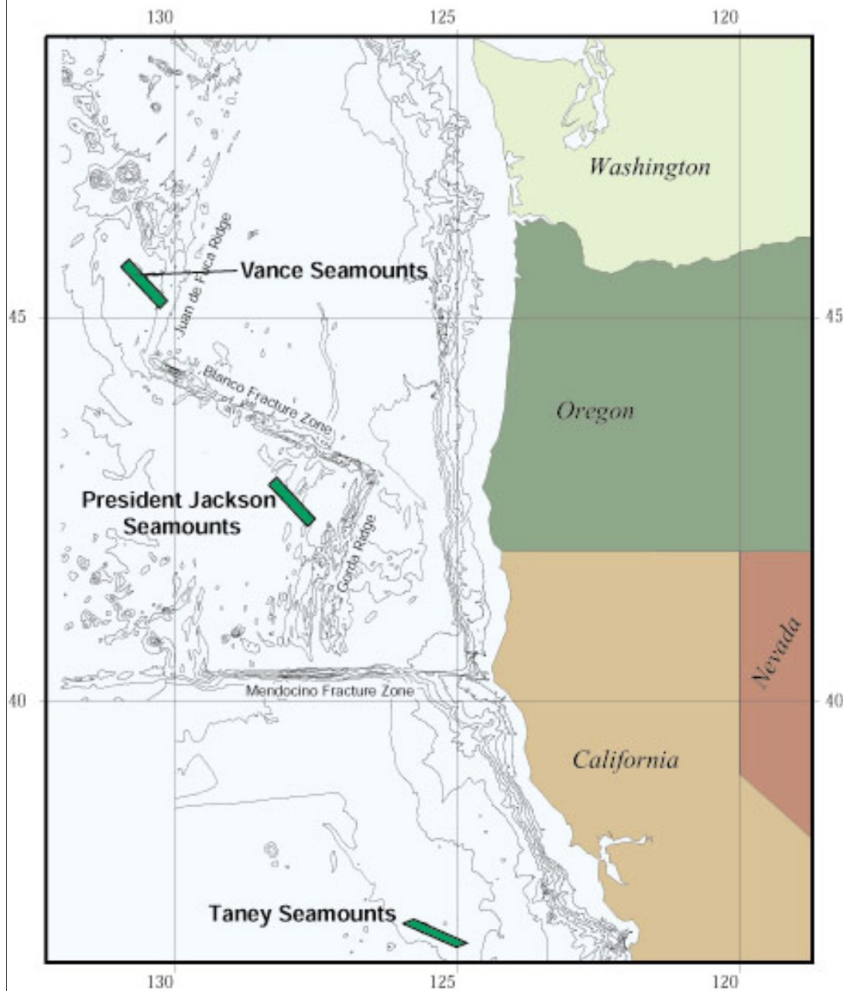
Volcanic History

[Valentine *et al.*, *GSA Bull.*, 2006]

10 Ma	1.5 Myr of basalt flows
4 Ma	smaller basalt flows
1 Ma	5 volcanoes
80 ka	Lanthrop Wells

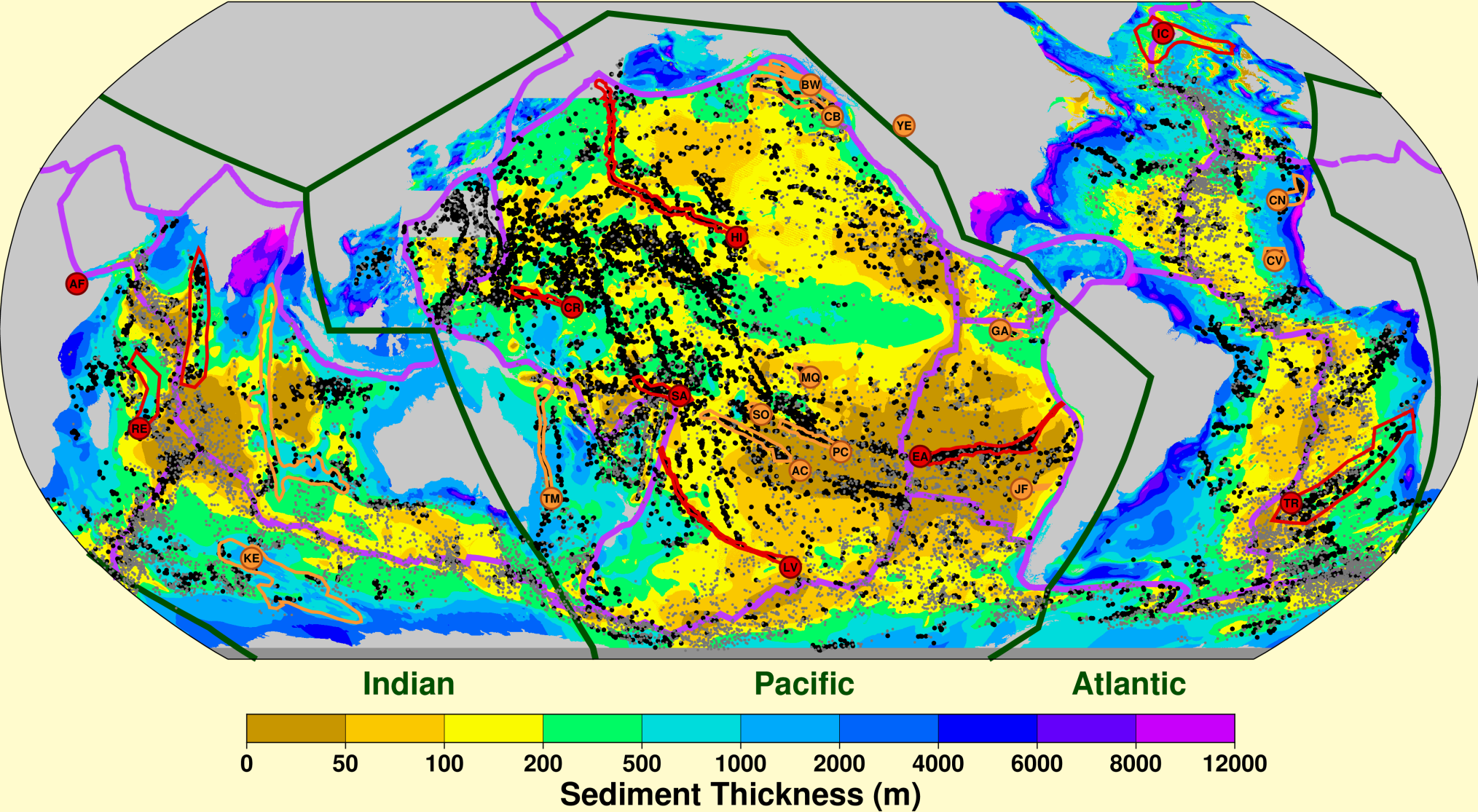


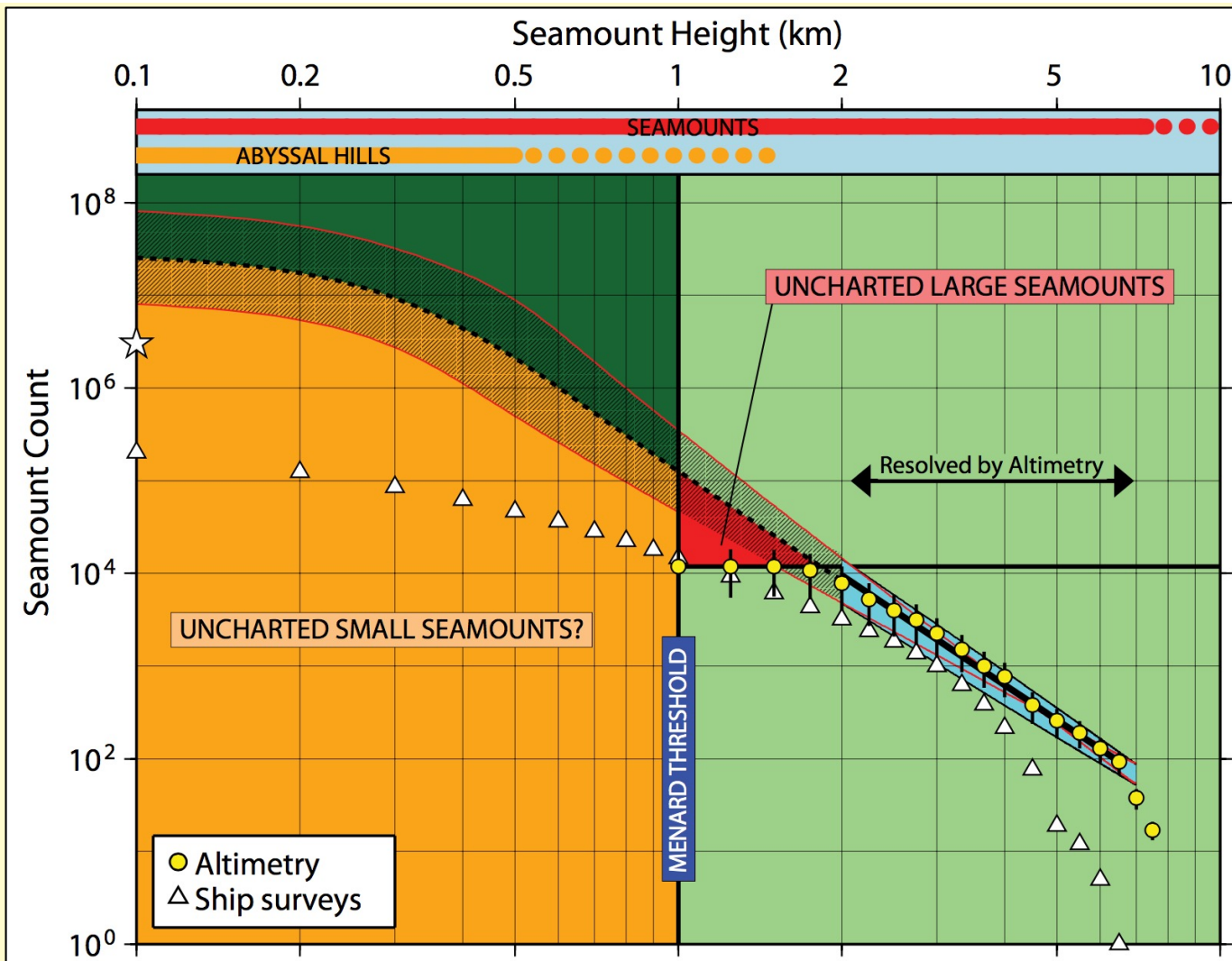
Seamounts represent past volcanism on the seafloor



President Jackson Seamounts

The *Kim & Wessel* [2011] Catalog
24,000+ seamounts detected using satellite gravity

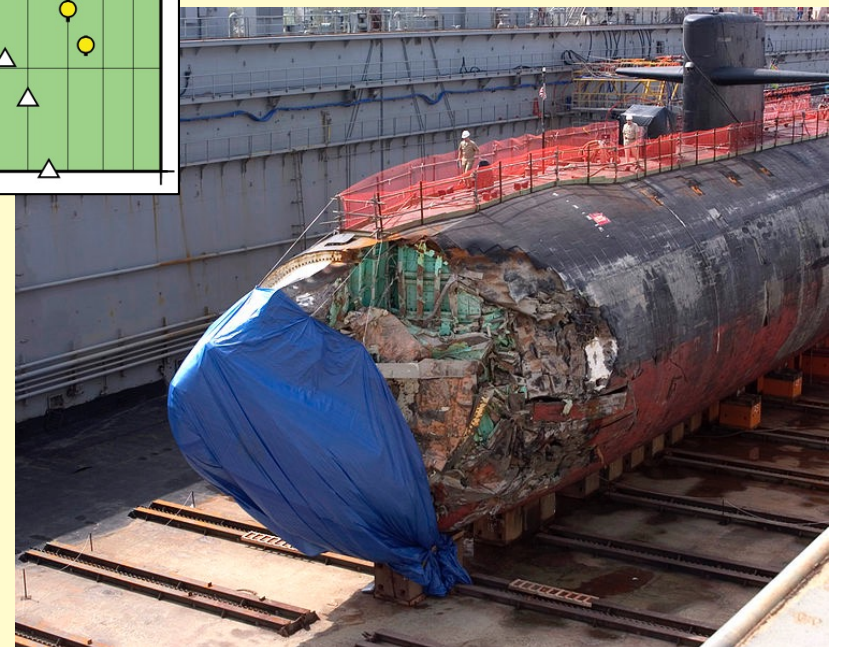




Seamount Census: Height – Frequency Distribution

Wessel et al.
[2010]

*USS San Francisco
Crashed into a uncharted seamount (2005)*

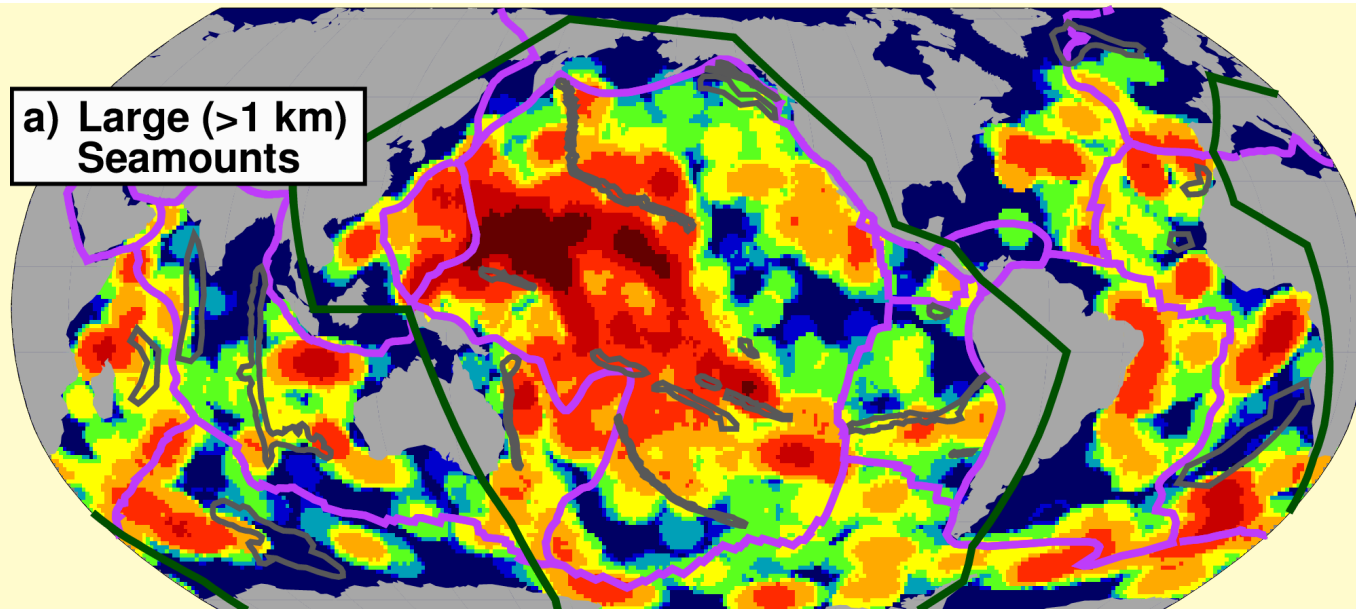


**Seamount
Equivalent
Layer
Thickness:**

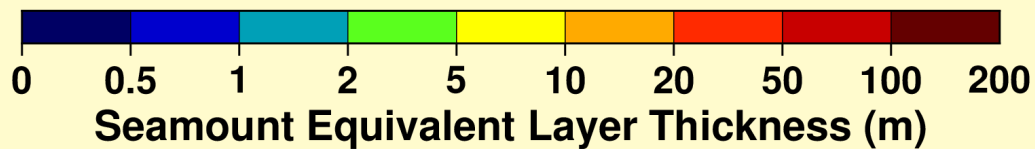
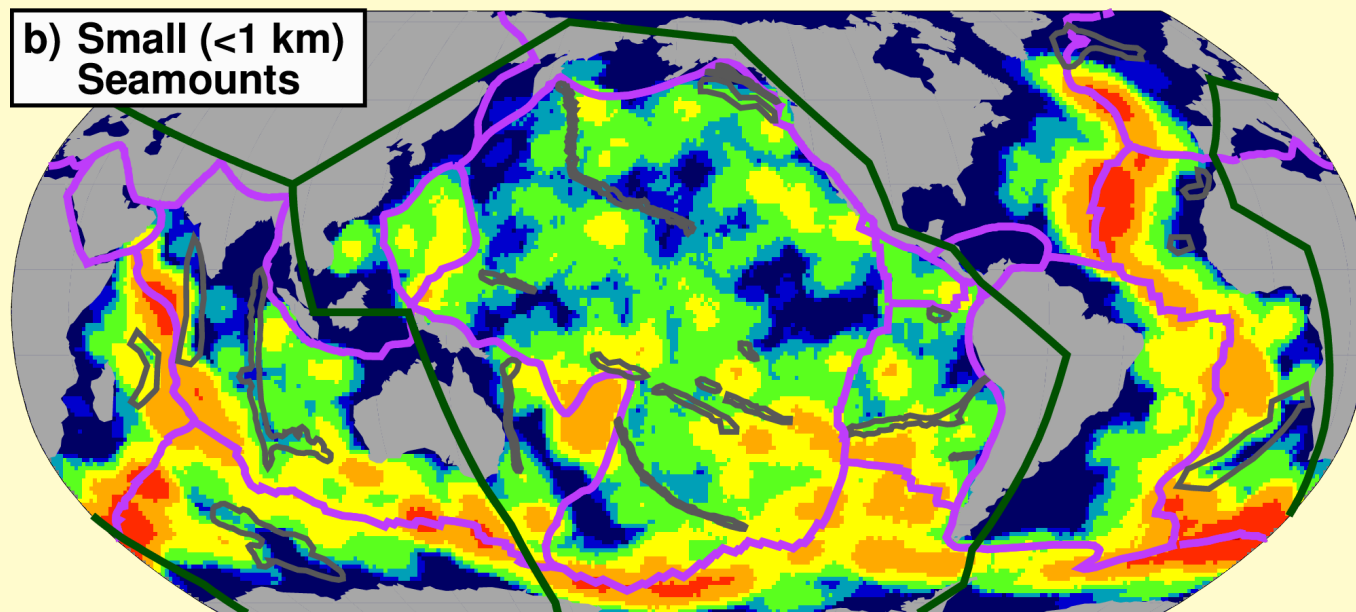
*Thickness of a
volcanic layer if
all seamounts
are spread
evenly across
nearby seafloor*

Conrad et al. [2017]

**a) Large (>1 km)
Seamounts**



**b) Small (<1 km)
Seamounts**



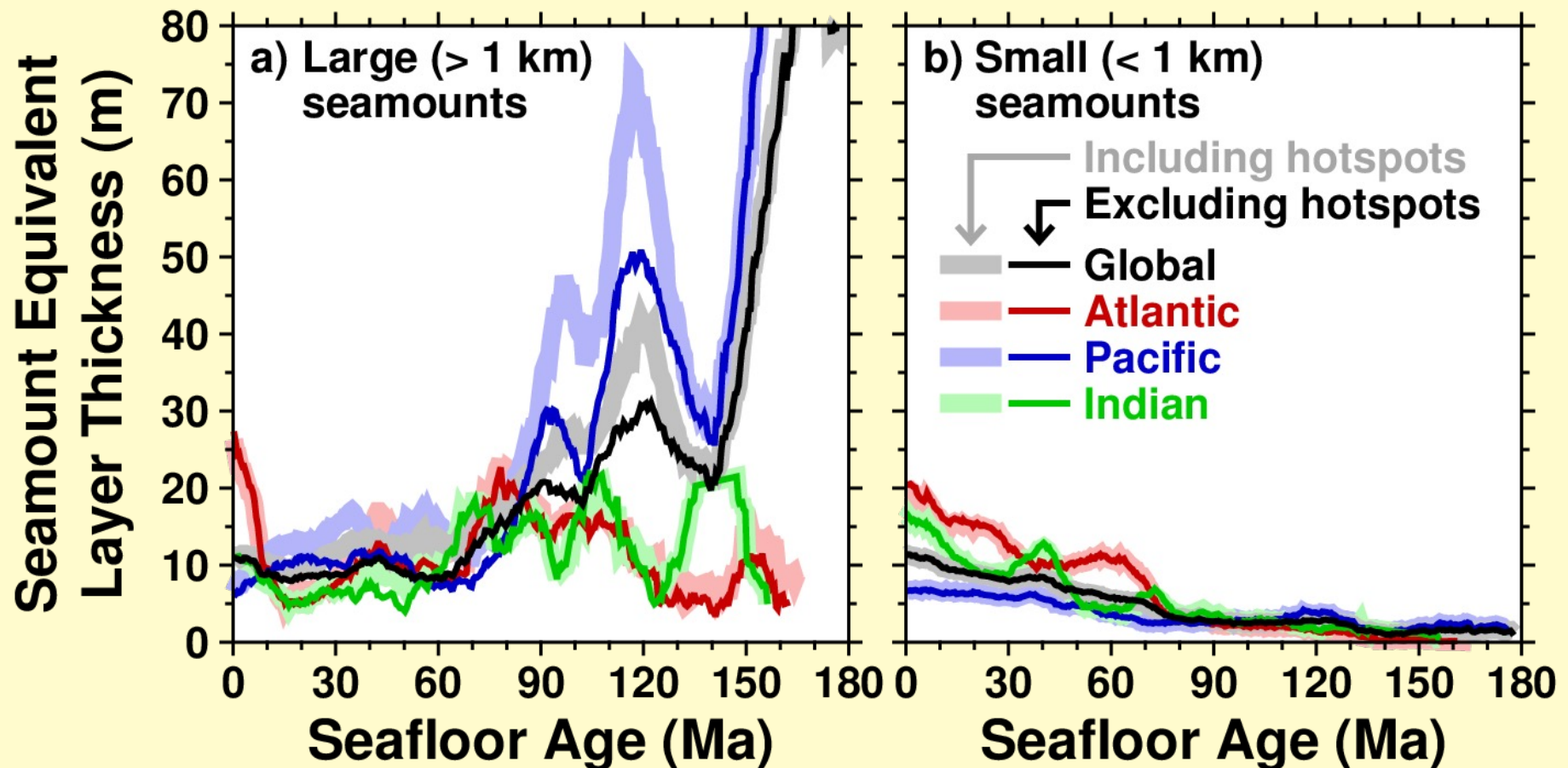
Seamount equivalent thickness vs. seafloor age

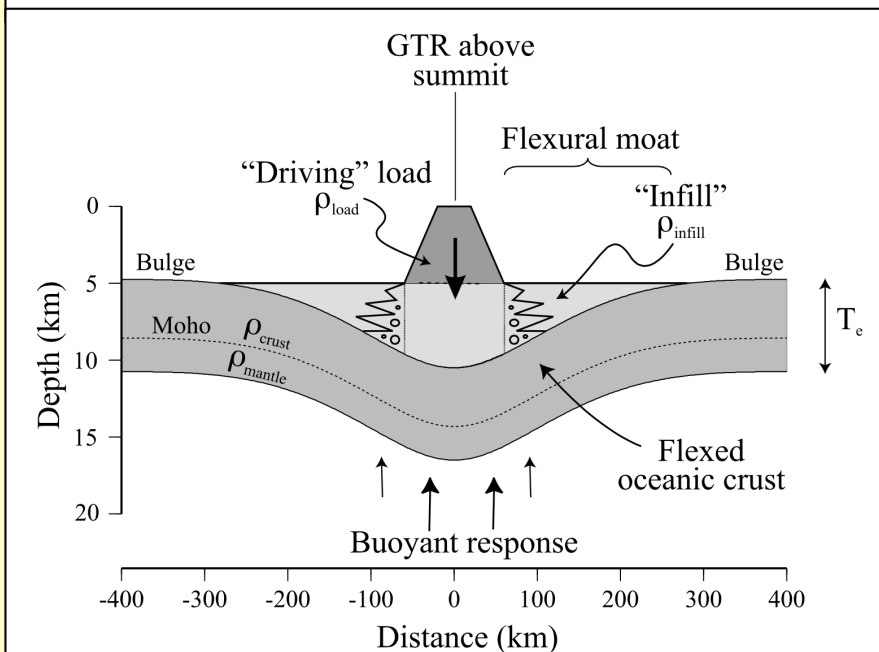
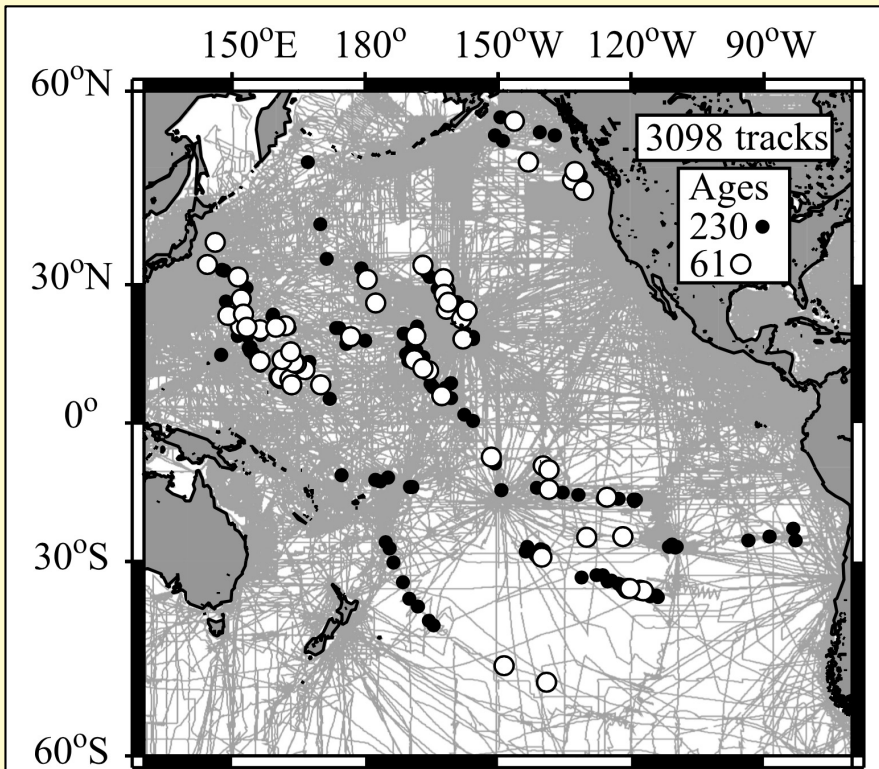
Large seamounts:

- 10 m until ~70 Myr
- Increases after 70 Myr (esp in Pacific)

Small seamounts:

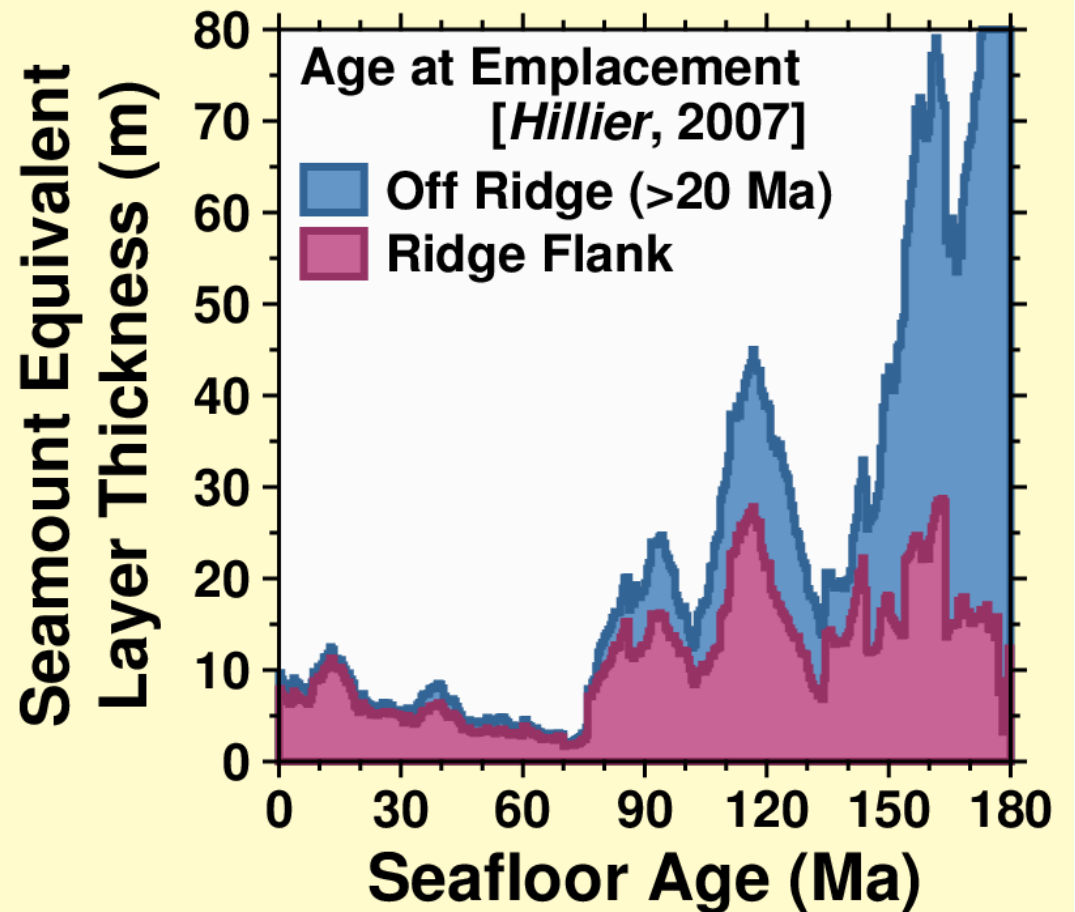
- Decreases away from the ridge due to sampling problems

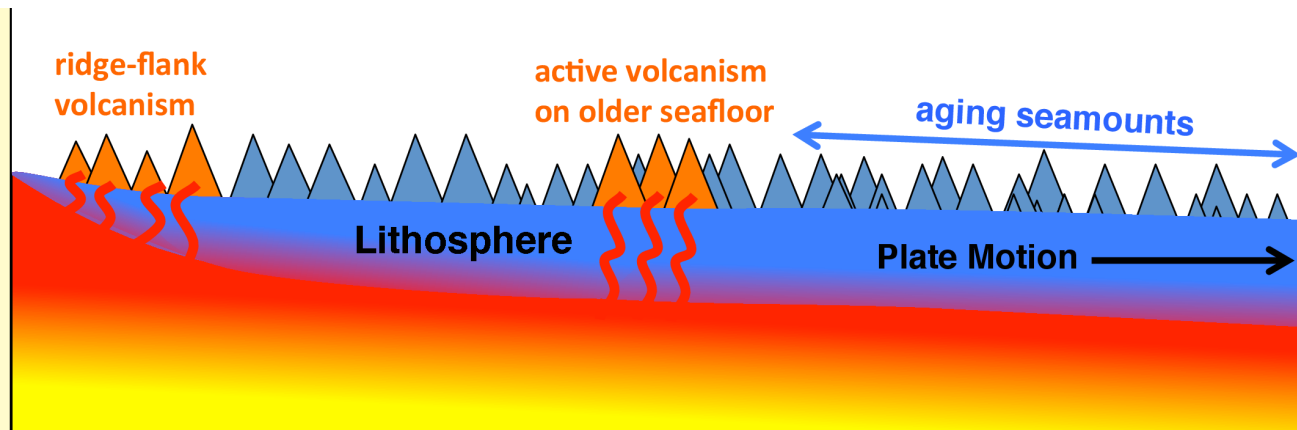




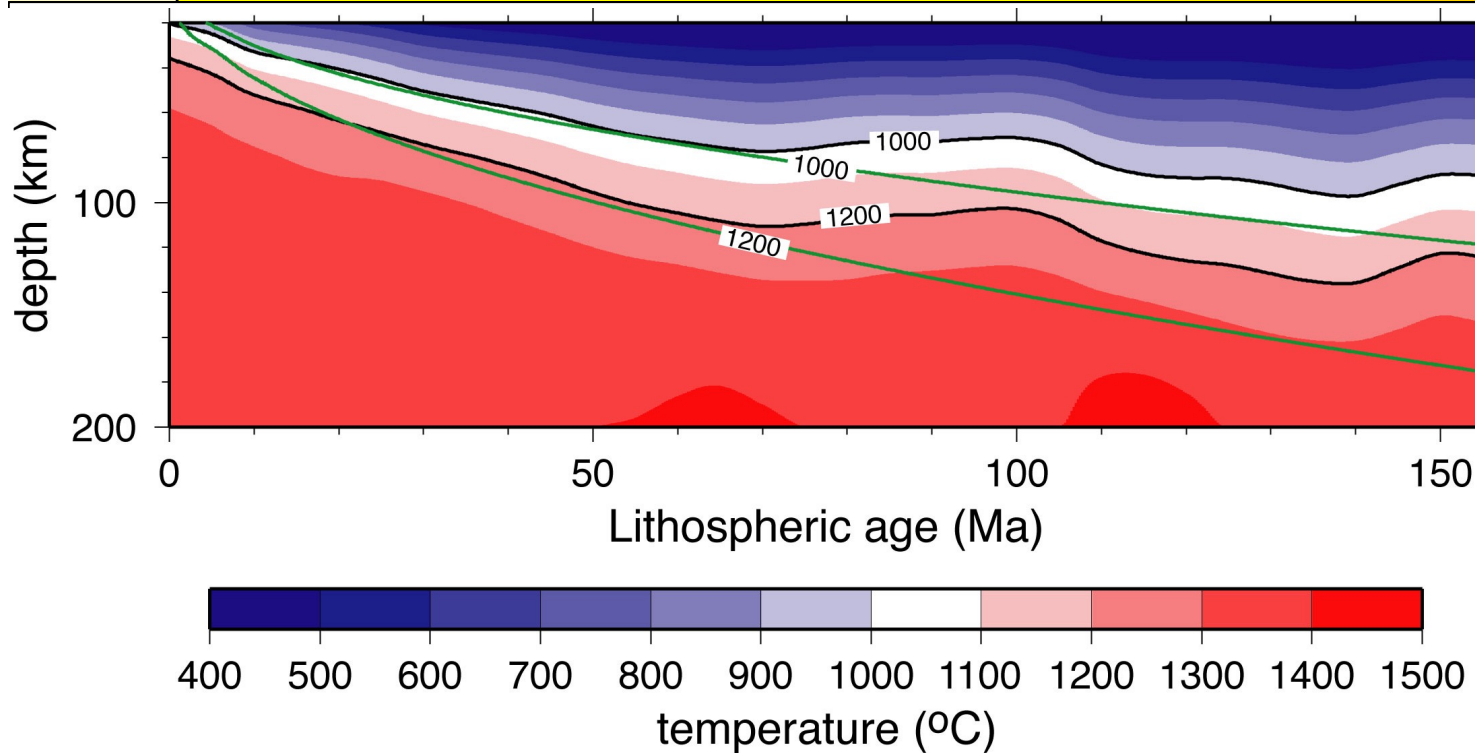
Age of the Seamounts:

- Poorly constrained!
- Must be younger than seafloor
- Constraints for some seamounts from plate flexure





Model of Seamount Emplacement
Conrad et al.
 [2017]



Thermal Profile of the Pacific Lithosphere
Ritzwoller et al.
 [2004]

Model for Seamount formation:

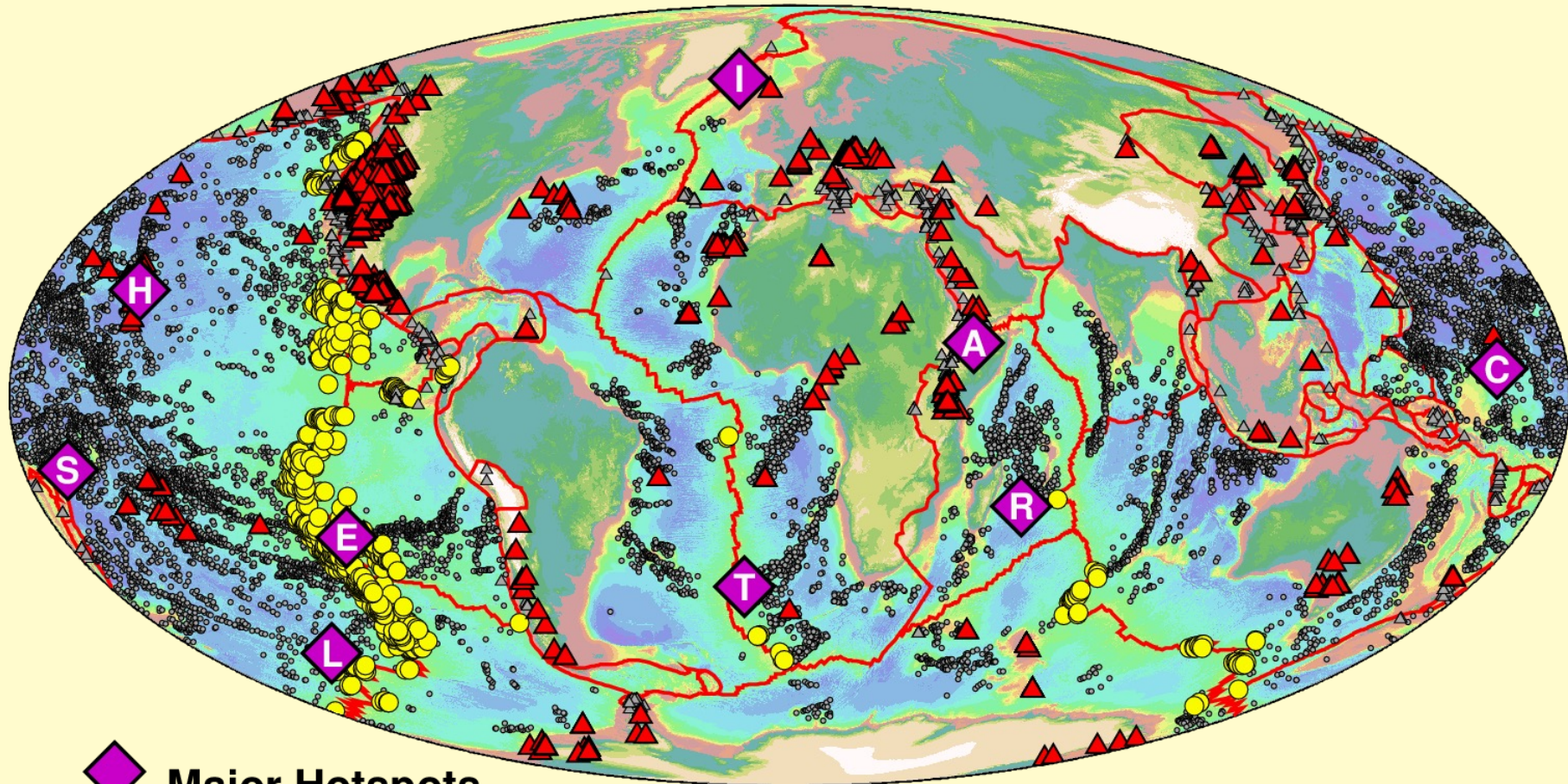
- One population formed on the ridge flanks
- Another population formed on seafloor > 70 Myr old
- Pacific seamount production was faster in the Cretaceous

Intraplate Volcanism

- Mostly minor and mostly basaltic
- Sometimes exhibits age non-hotspot-like age progression

→ How are the lithosphere and asthenosphere important?

→ What is the mechanism?



Major Hotspots



Young Seamounts (< 10 Myr)



Off-Ridge Seamounts

▲ Intraplate Volcanoes

△ Plate-Boundary Volcanoes

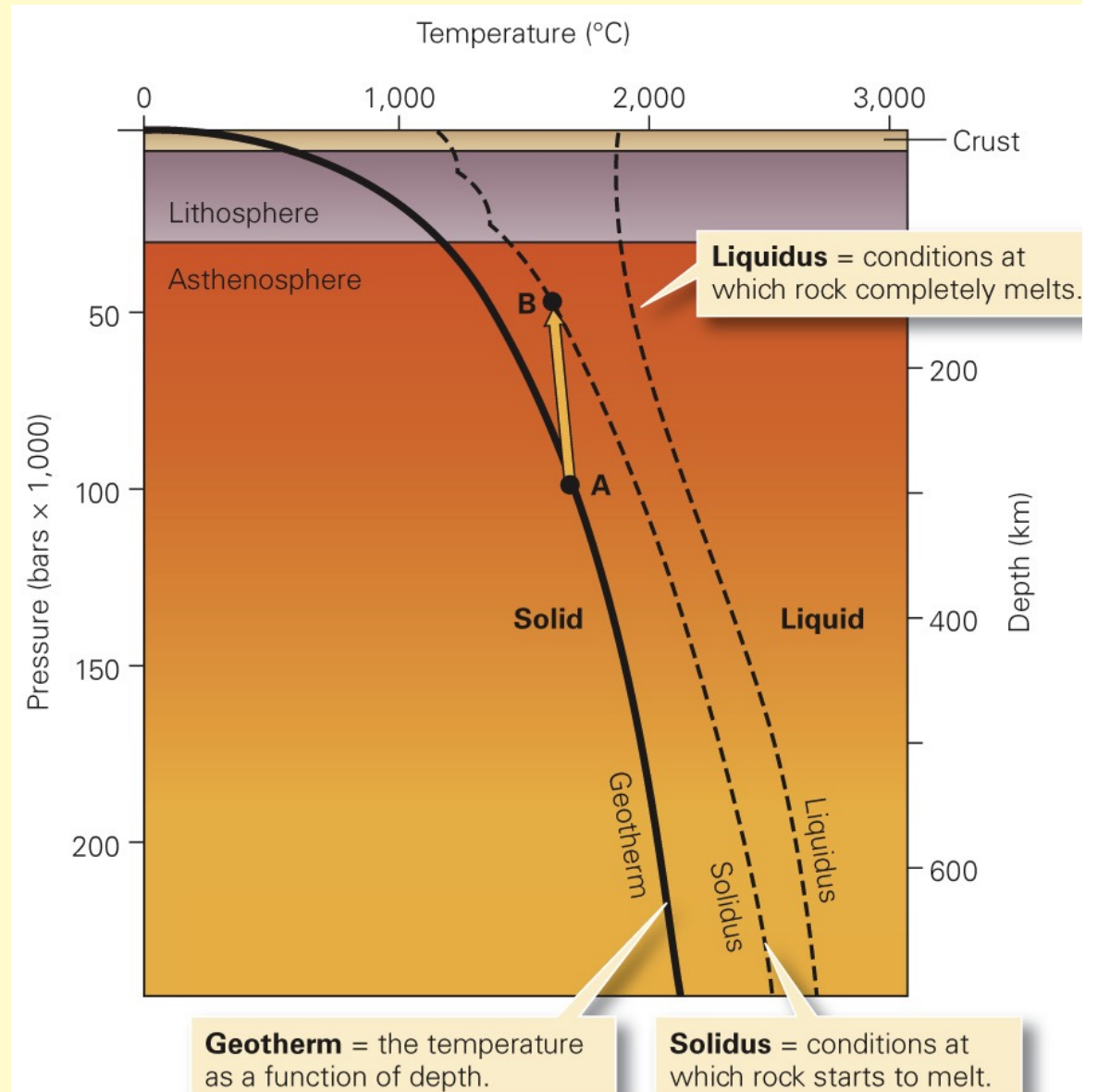
The following can produce melting:

- Hot Mantle Temperatures
- Reduced Mantle Solidus (e.g., due to volatiles)
- Mantle upwelling

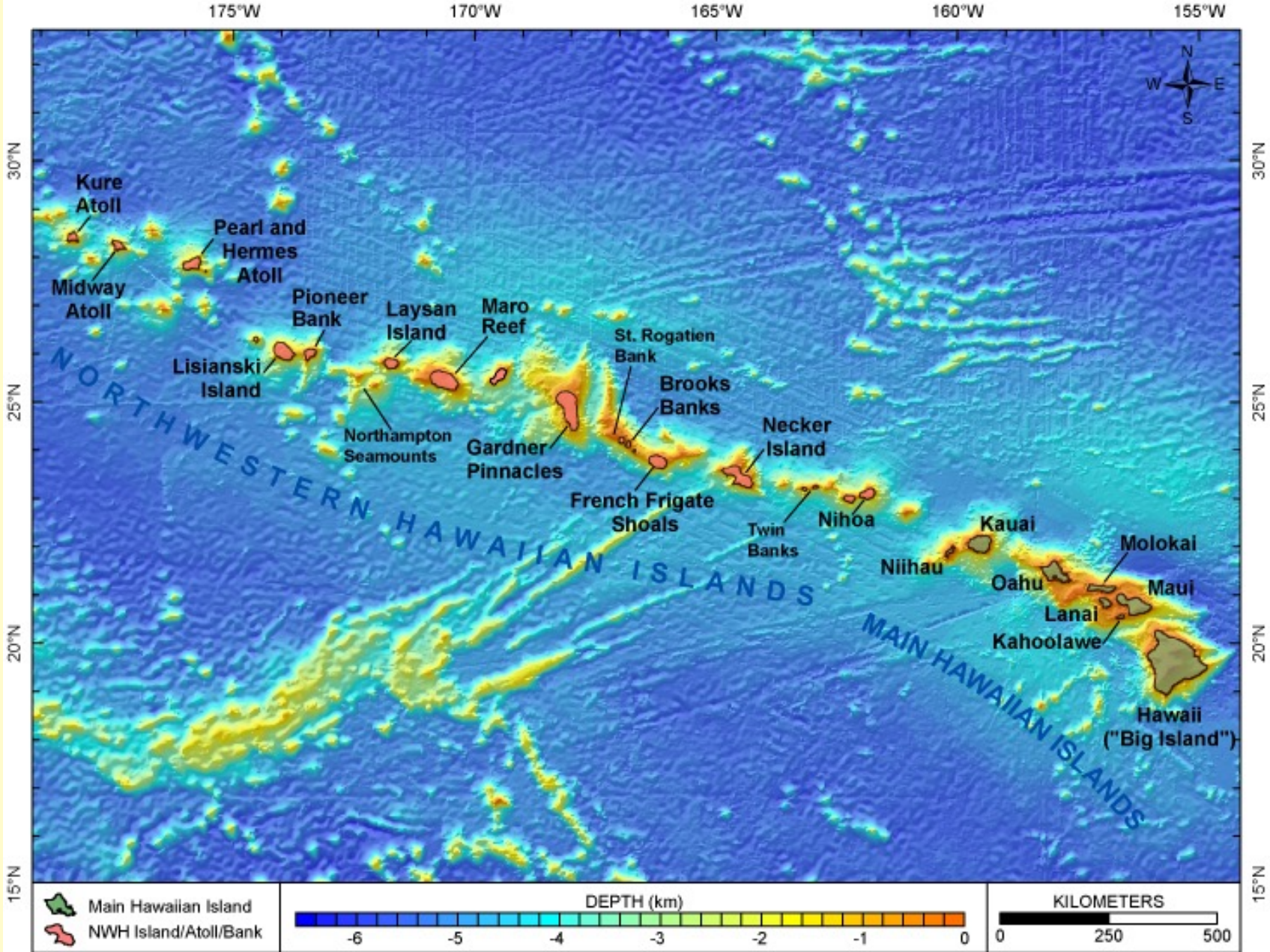
→ Which mechanisms can cause localized mantle upwelling?

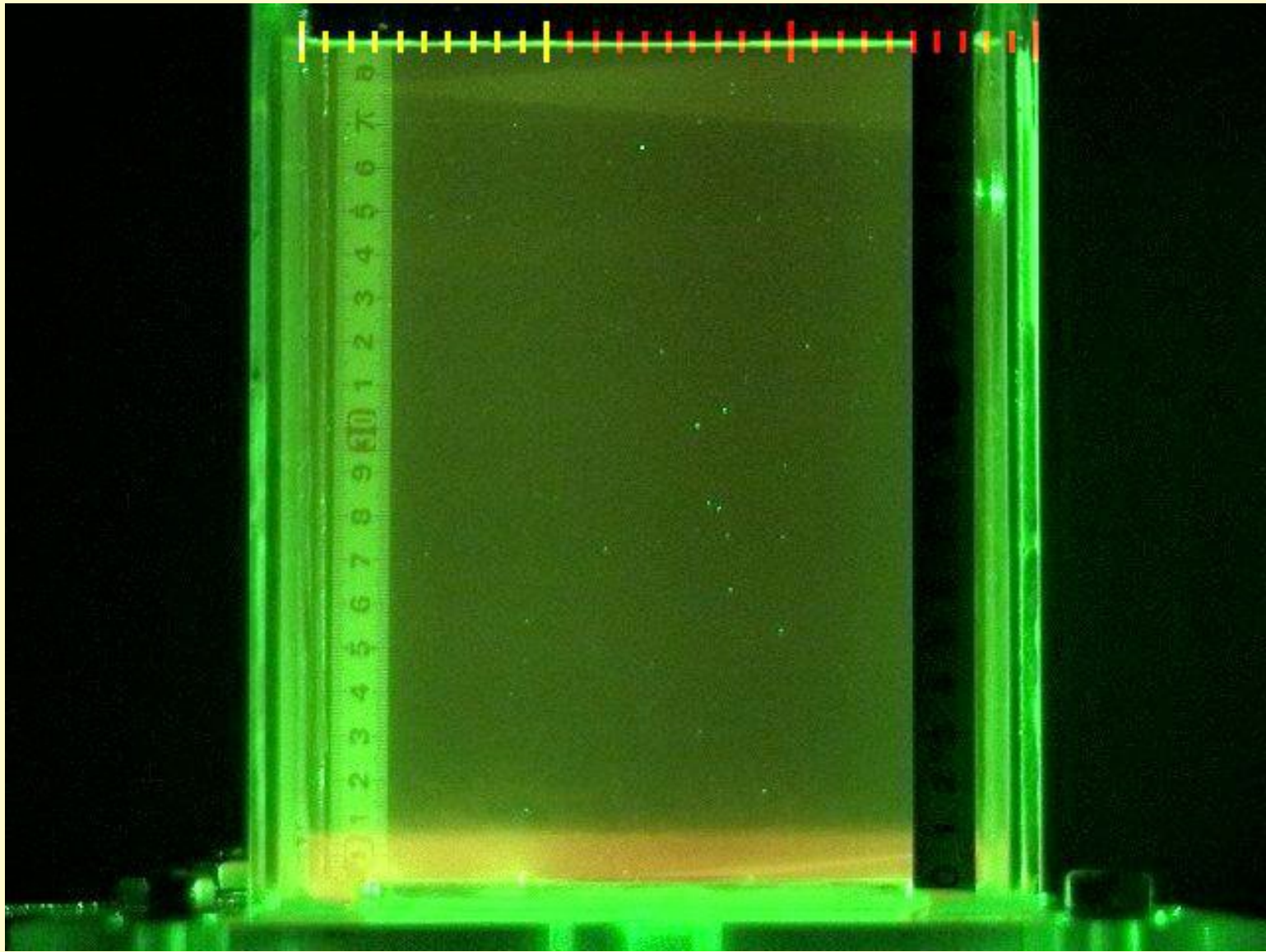
5 Mechanisms:

- Plumes
- Small-Scale Convection
- Shear-Driven Upwelling
- Petit-Spots
- Surface Loading



1. Mantle Plumes





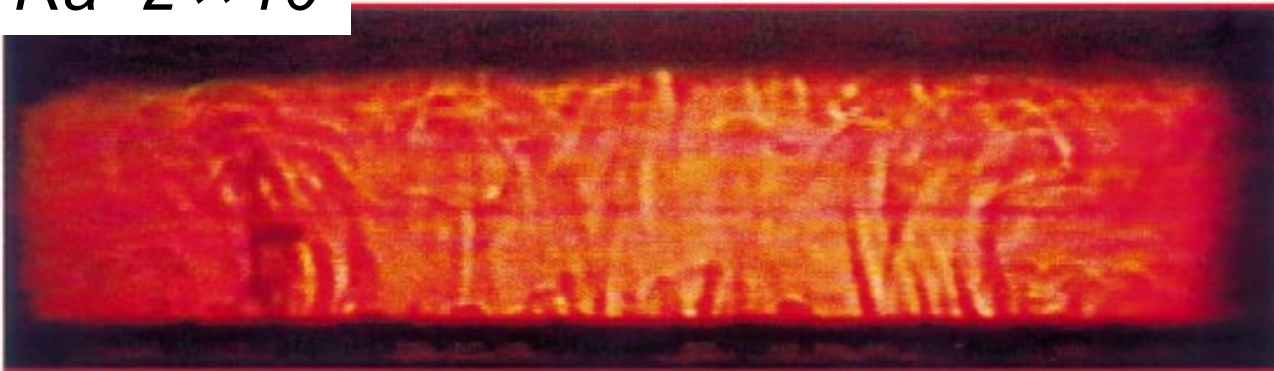
A Plume Experiment in Corn Syrup:
Plumes have heads and tails

$Ra \sim 4 \times 10^6$

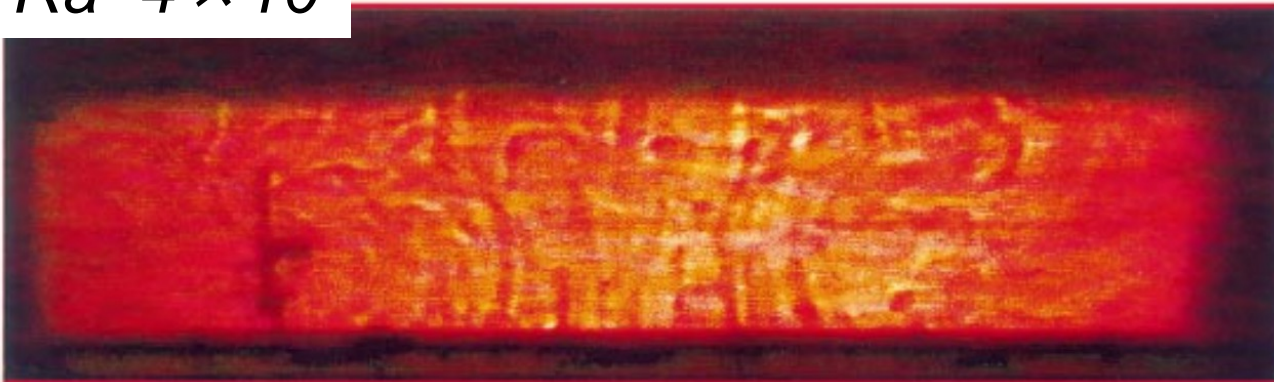


← **Base is Hot** →

$Ra \sim 2 \times 10^7$



$Ra \sim 4 \times 10^7$

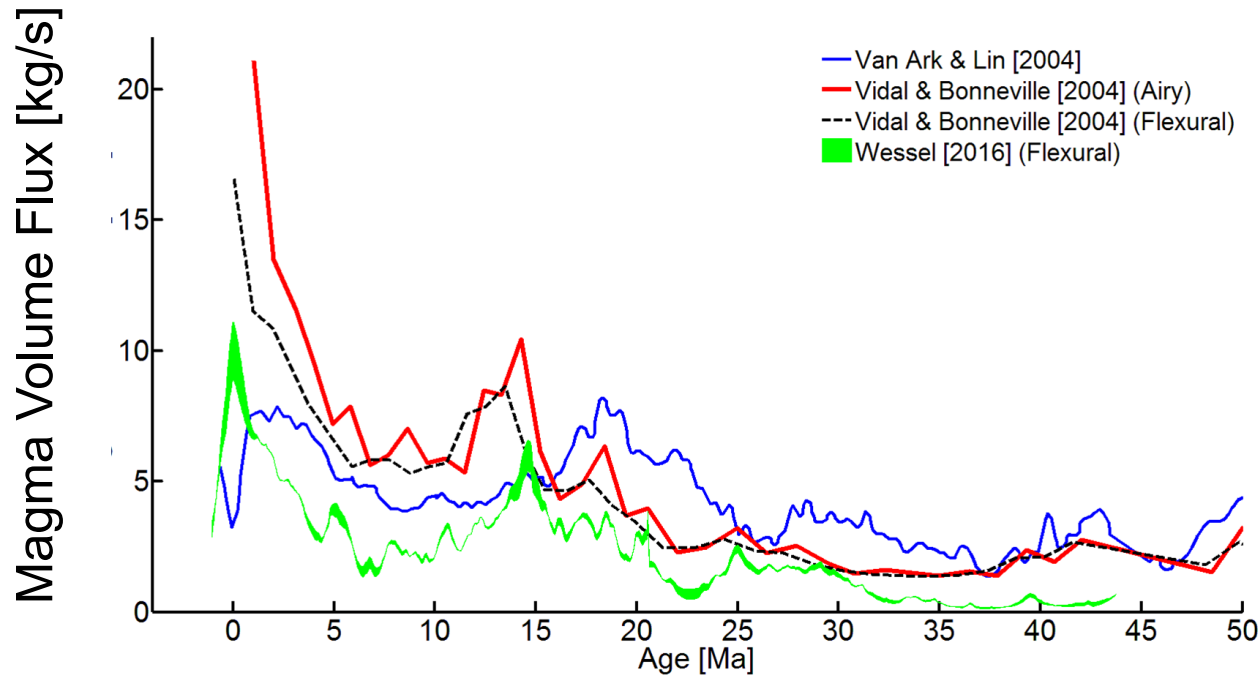
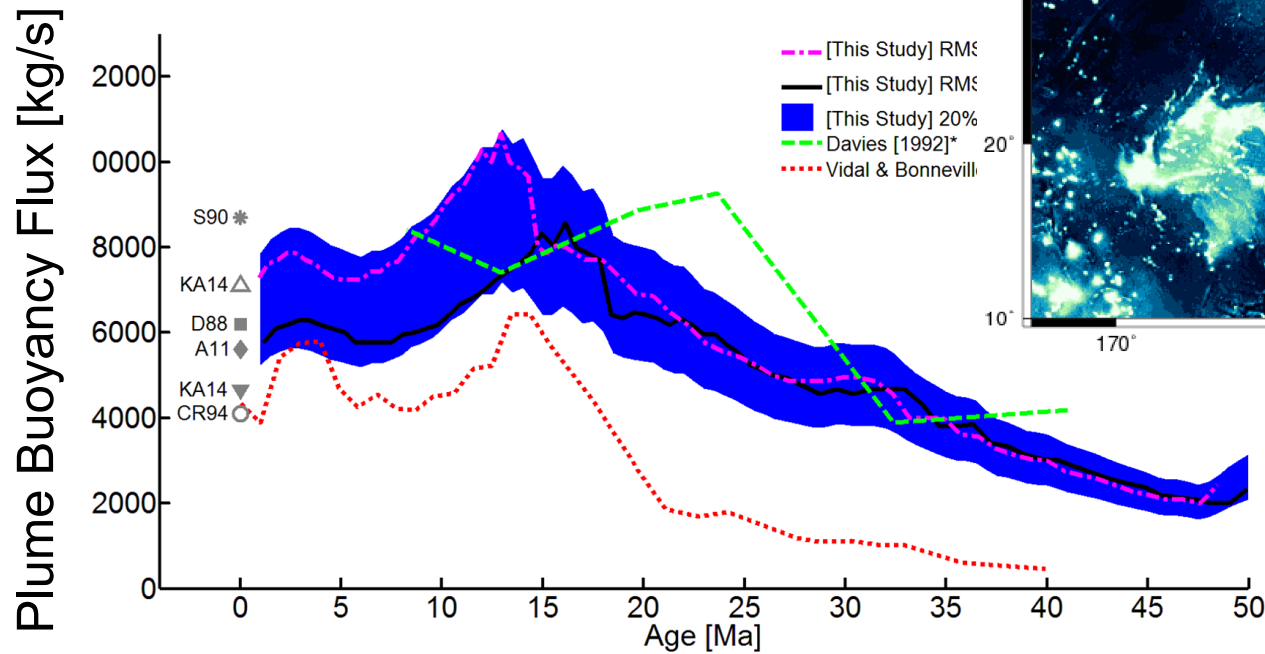
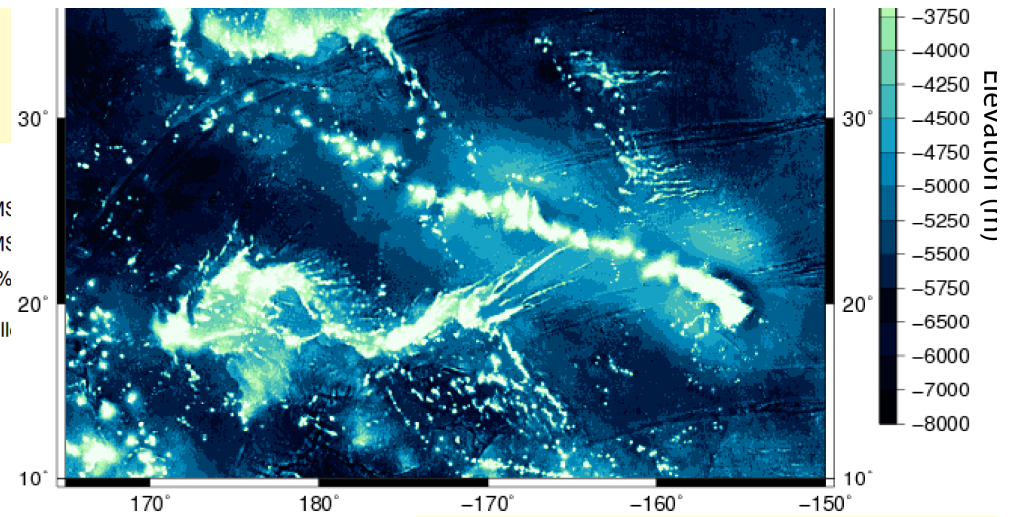


Laboratory experiment of convection in a tank of corn syrup.

Lithgow-Bertelloni et al. [2001]

More vigorous convection (hotter) makes for: smaller heads and thinner tails

Plumes uplift the lithosphere



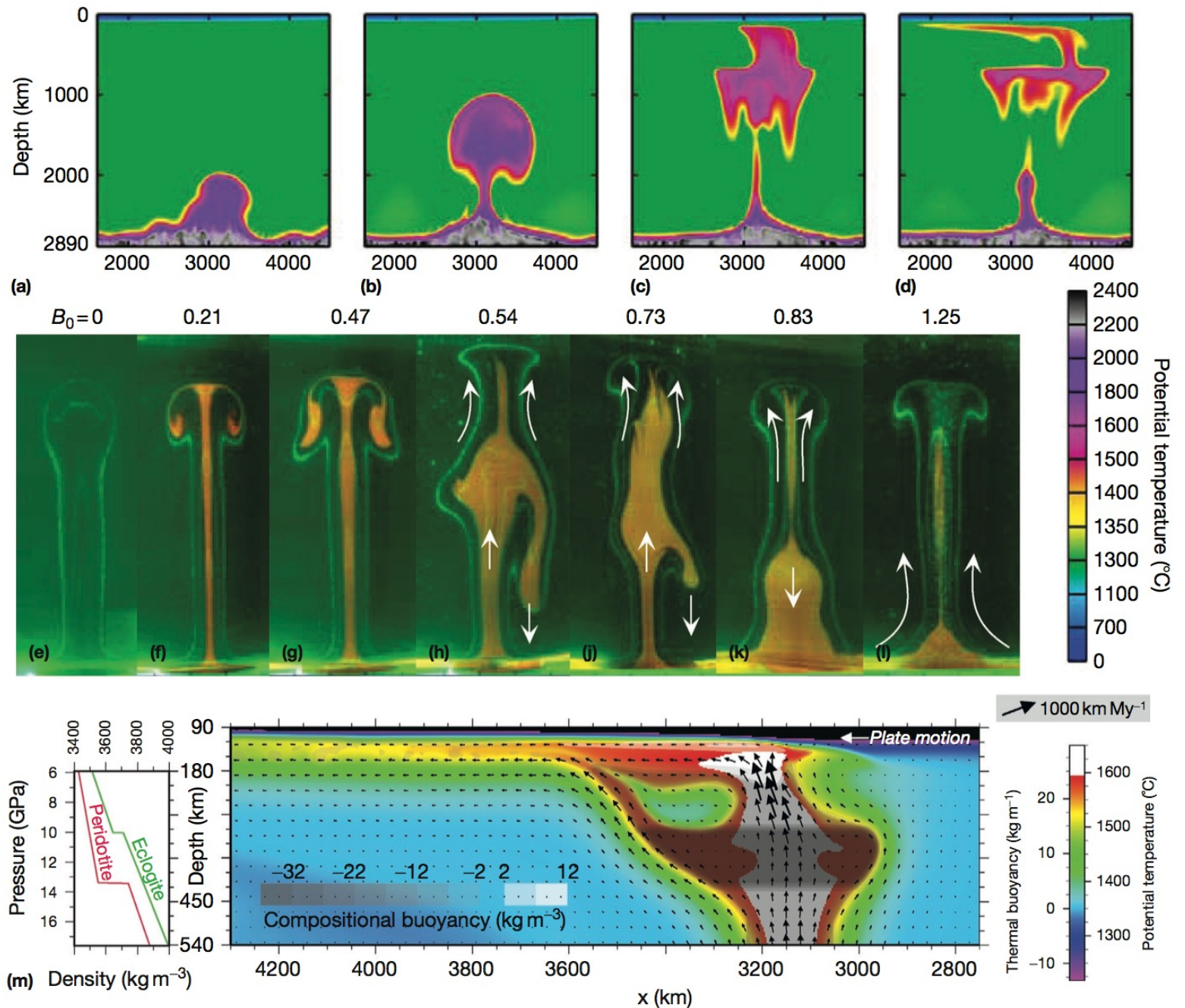
Hawaiian plume buoyancy flux has been growing with time

Togia, MS thesis [2015]

Plumes can exhibit interesting behavior!

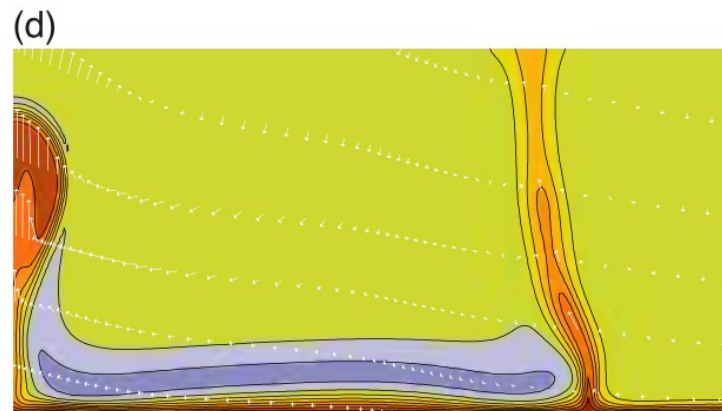
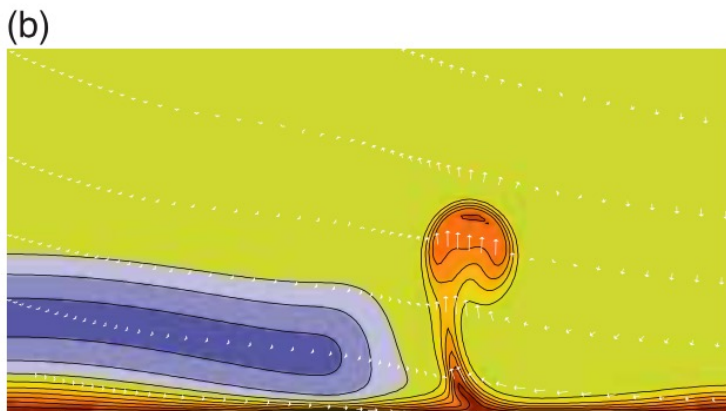
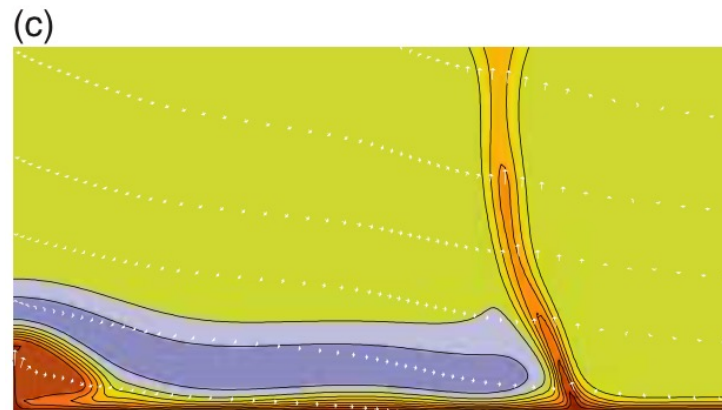
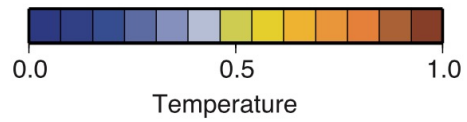
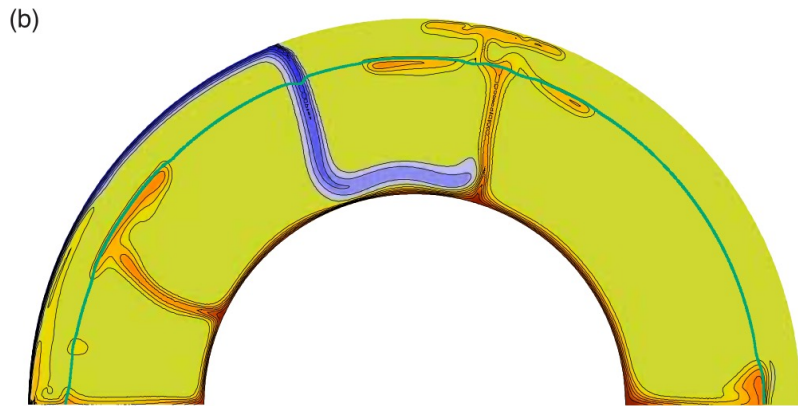
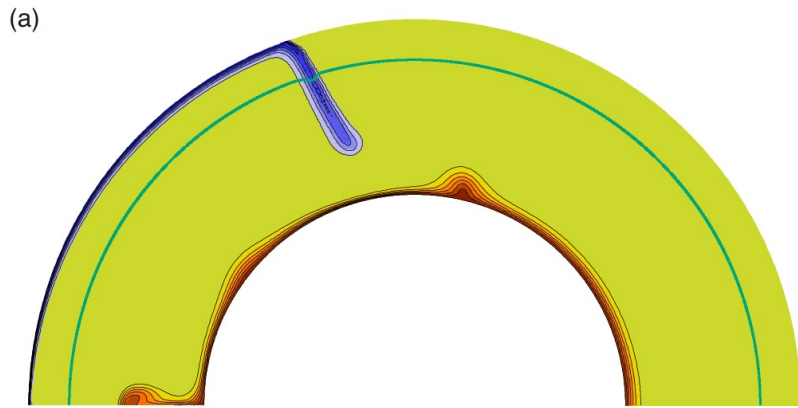
410 km:
Plume
Accelerated

610 km:
Plume
Impeded



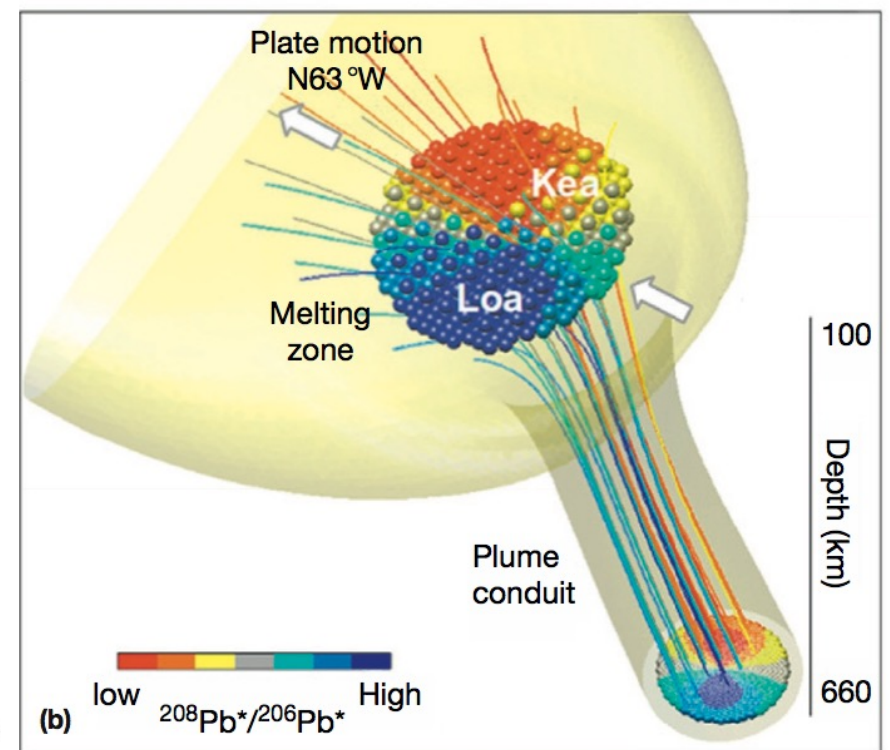
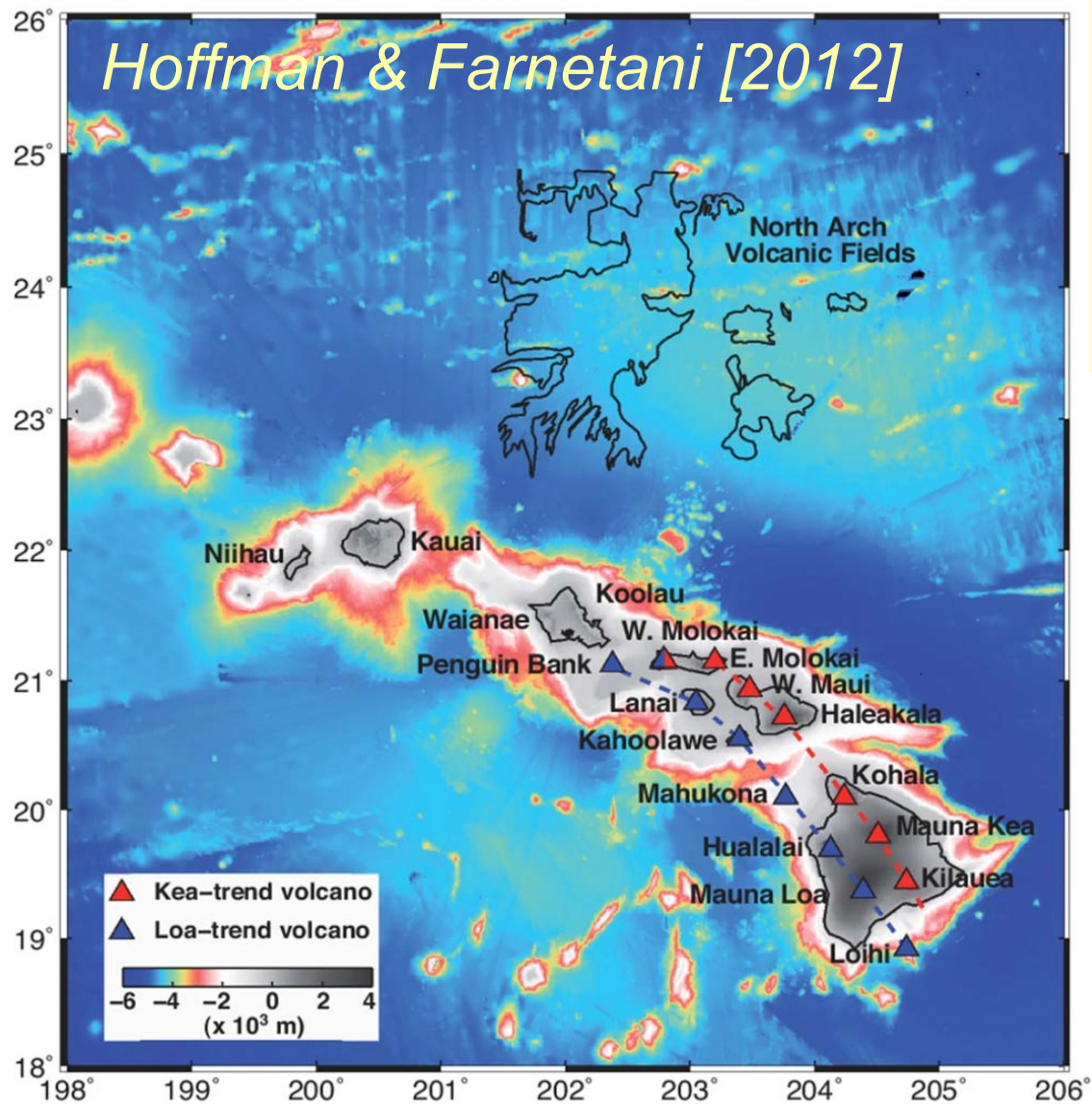
Ballmer et al. [2015]

Interaction of plumes and slabs

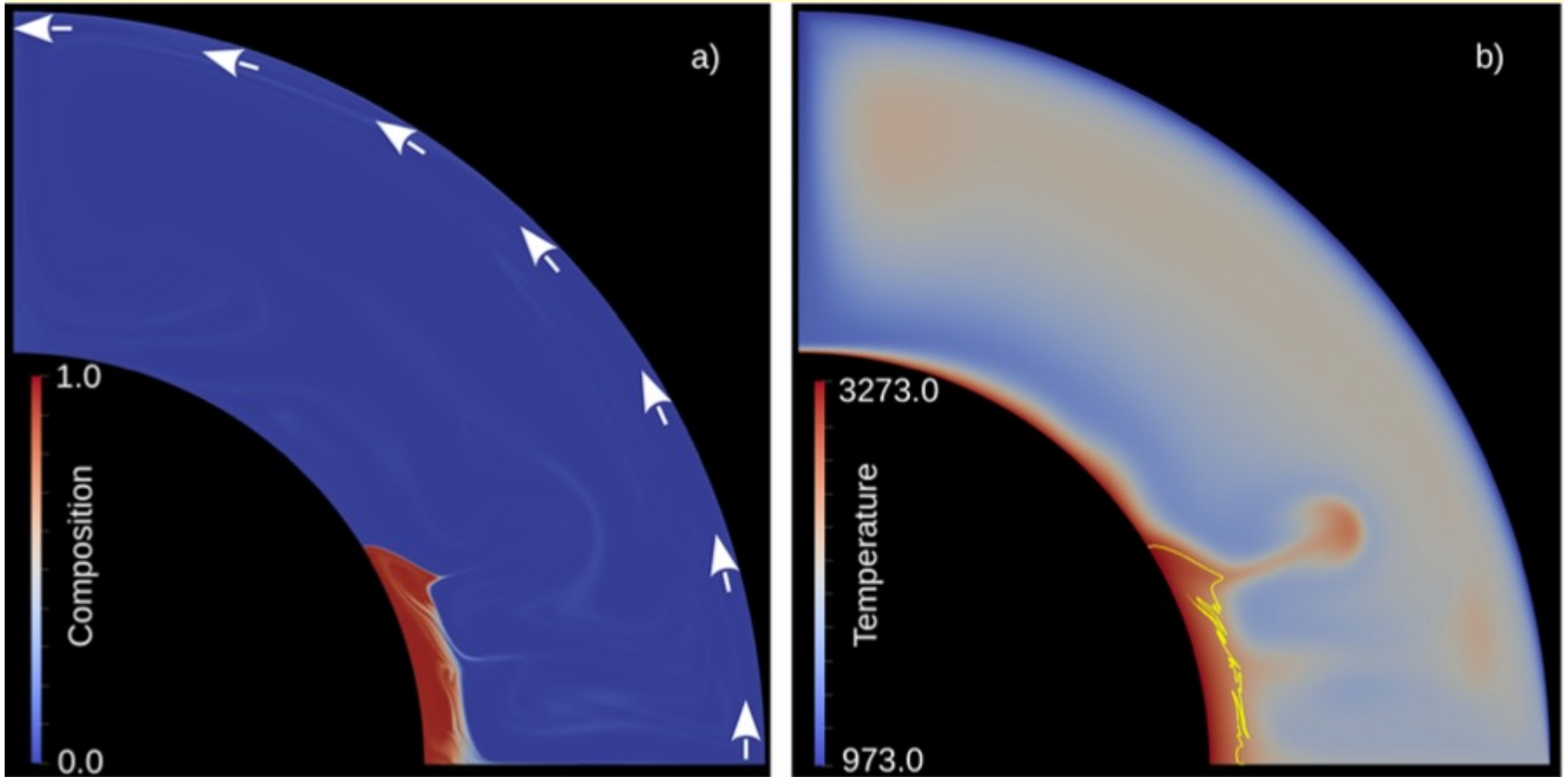


Tan & Gurnis
[2002]

Bisected plumes may sample preserved geochemical gradients from the deep mantle



Bisected plumes may sample preserved geochemical gradients from the deep mantle

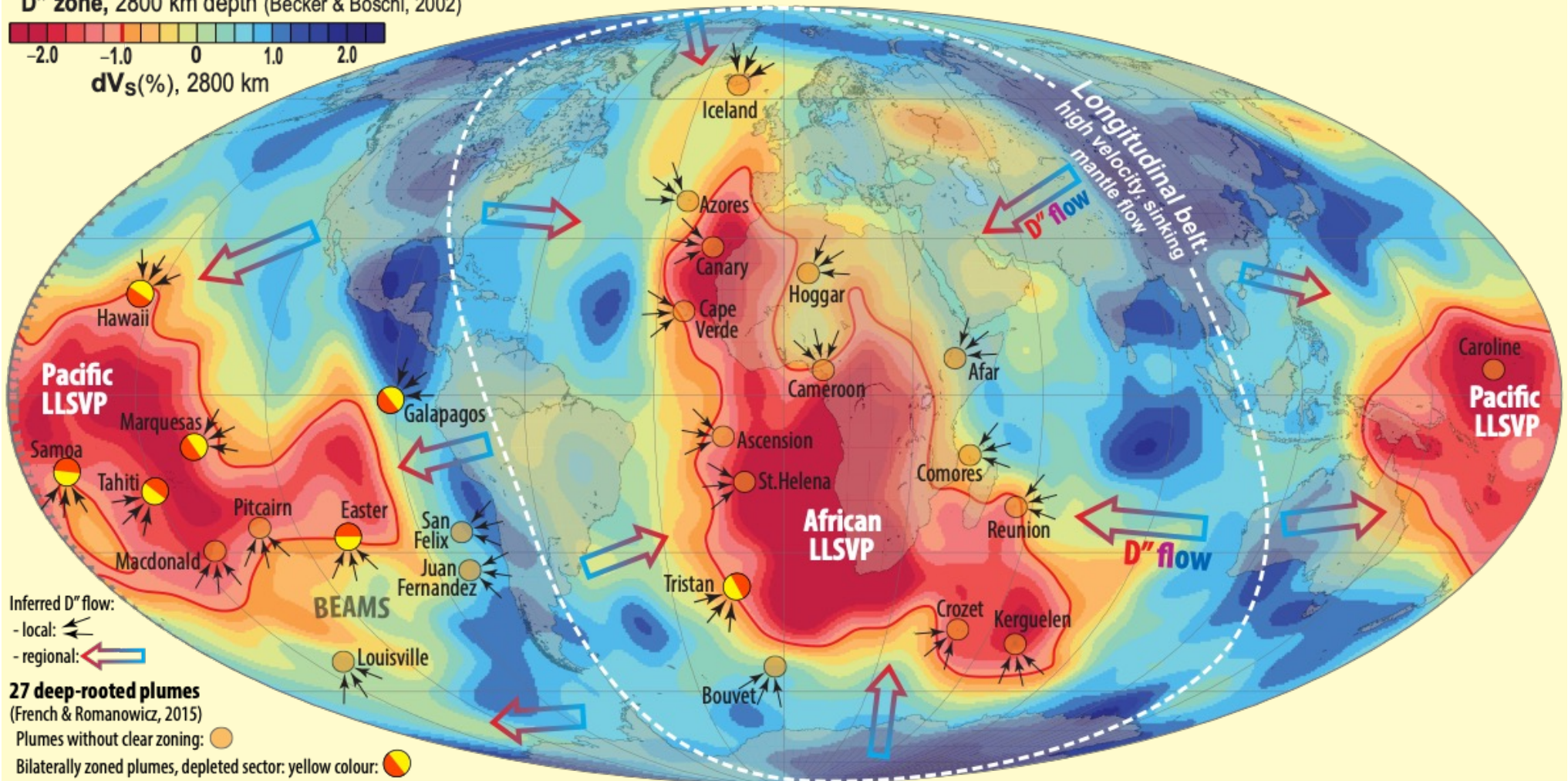
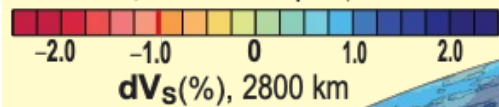


Plumes rise from the edges of the LLSVPs in the lower mantle
[Heyn *et al.*, 2020]

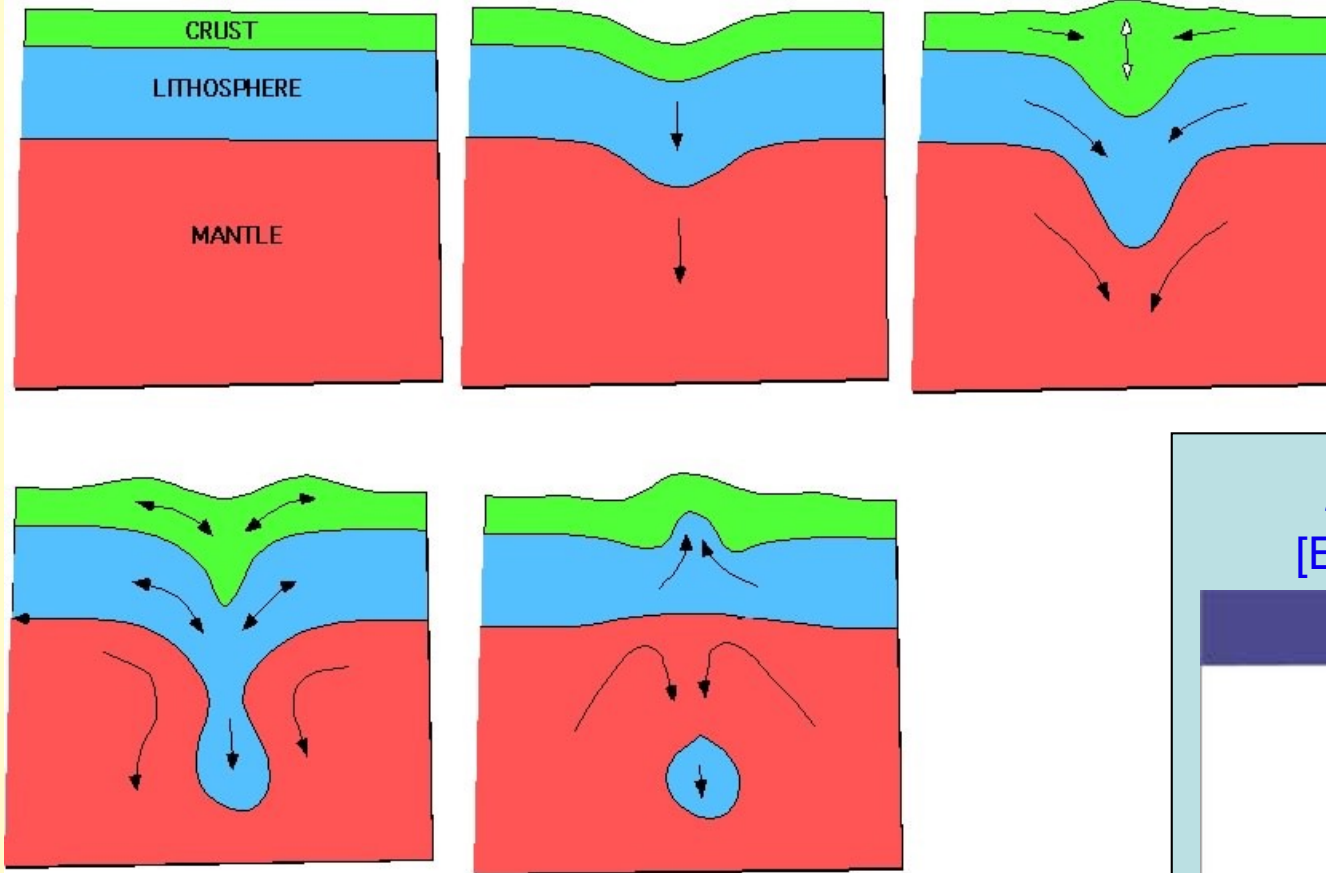
Bisected plumes may sample preserved geochemical gradients from the deep mantle

Base map: **SMEAN** S-wave tomography model

D'' zone, 2800 km depth (Becker & Boschi, 2002)

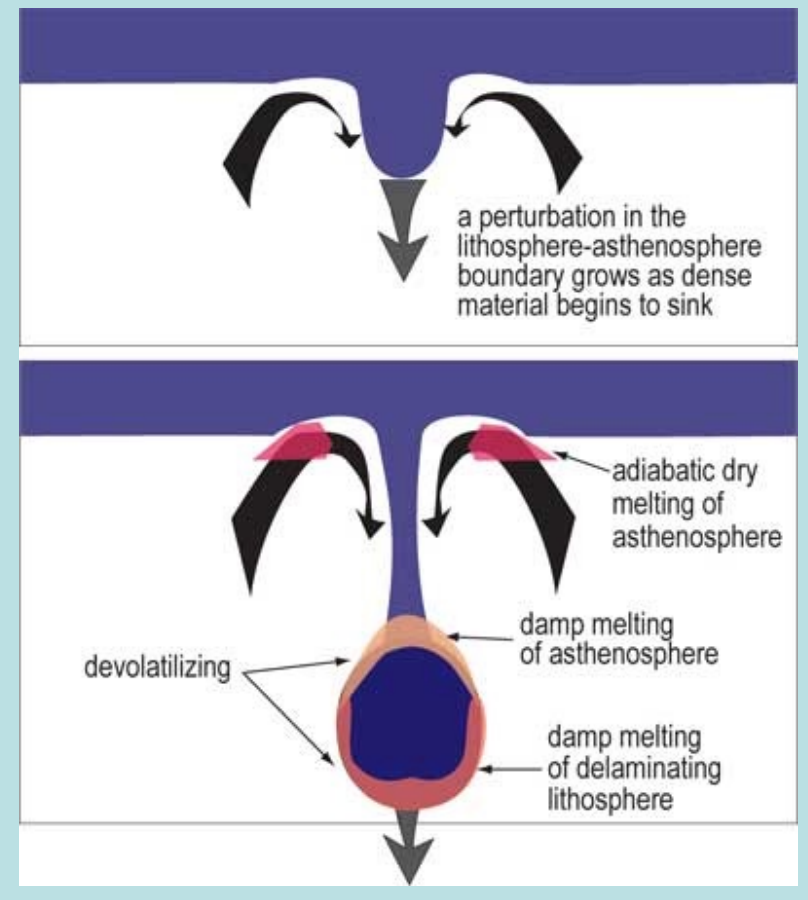


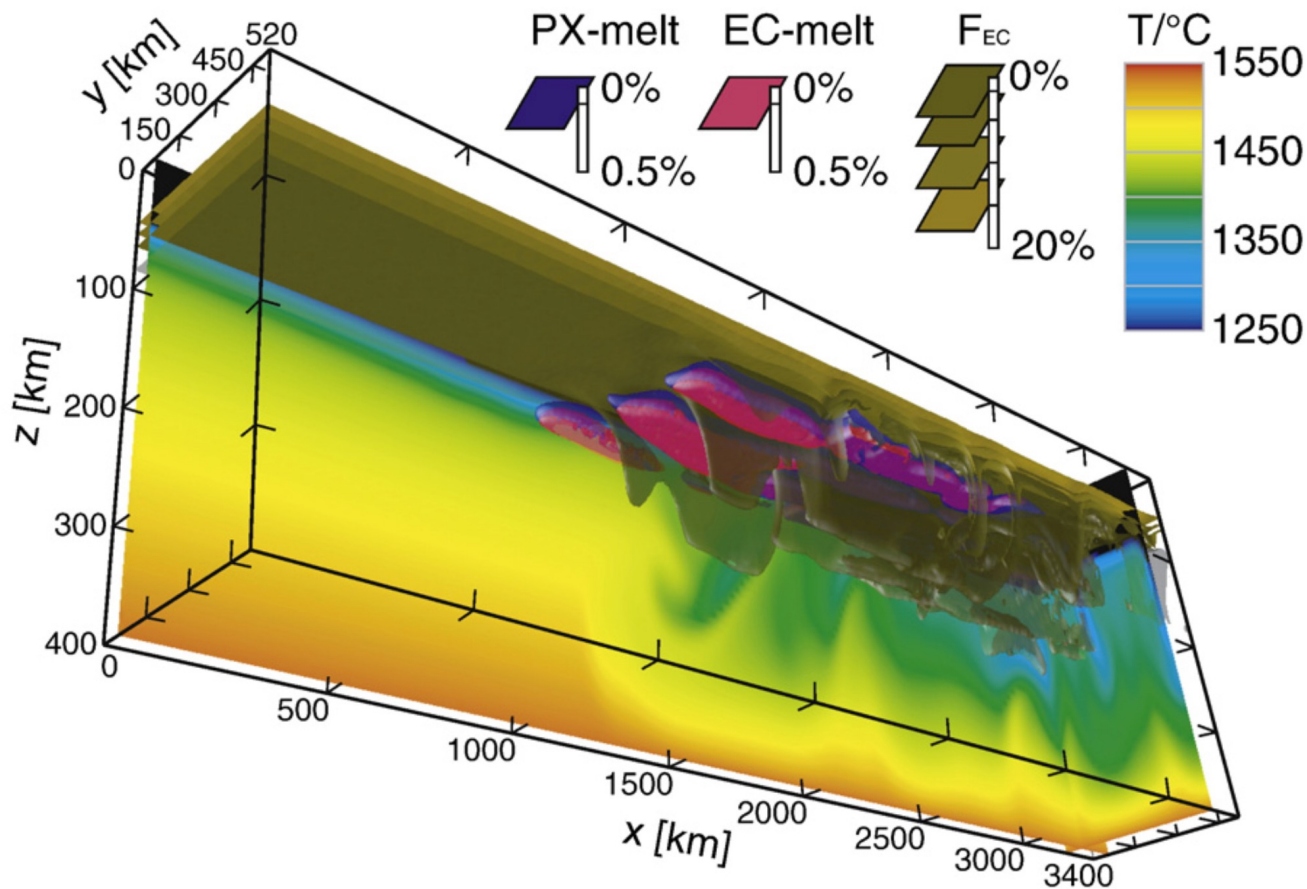
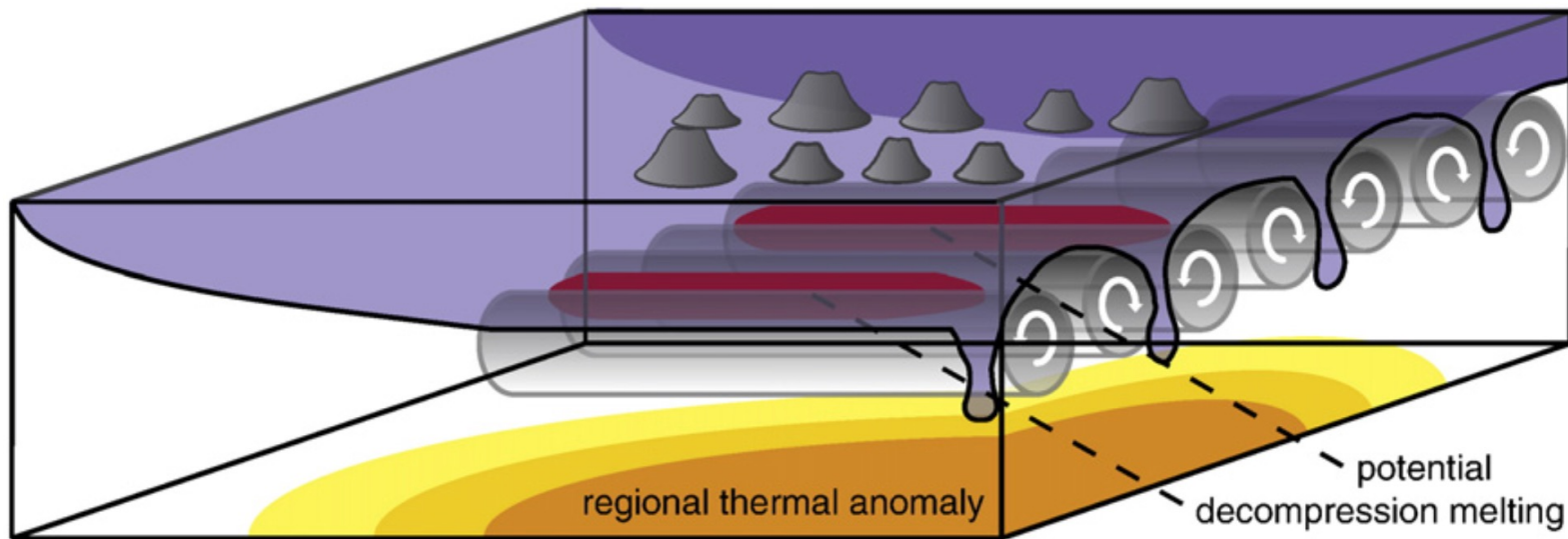
2. Small-Scale Convection



- Like an upside-down plume:
 - Cold “drips” sink into the mantle.
 - Return flow involves upwelling
→ produces melting.
- Explains some intraplate volcanism (continents & oceans)

Lithospheric Drips [Elkins-Tanton & Hager, 2000]

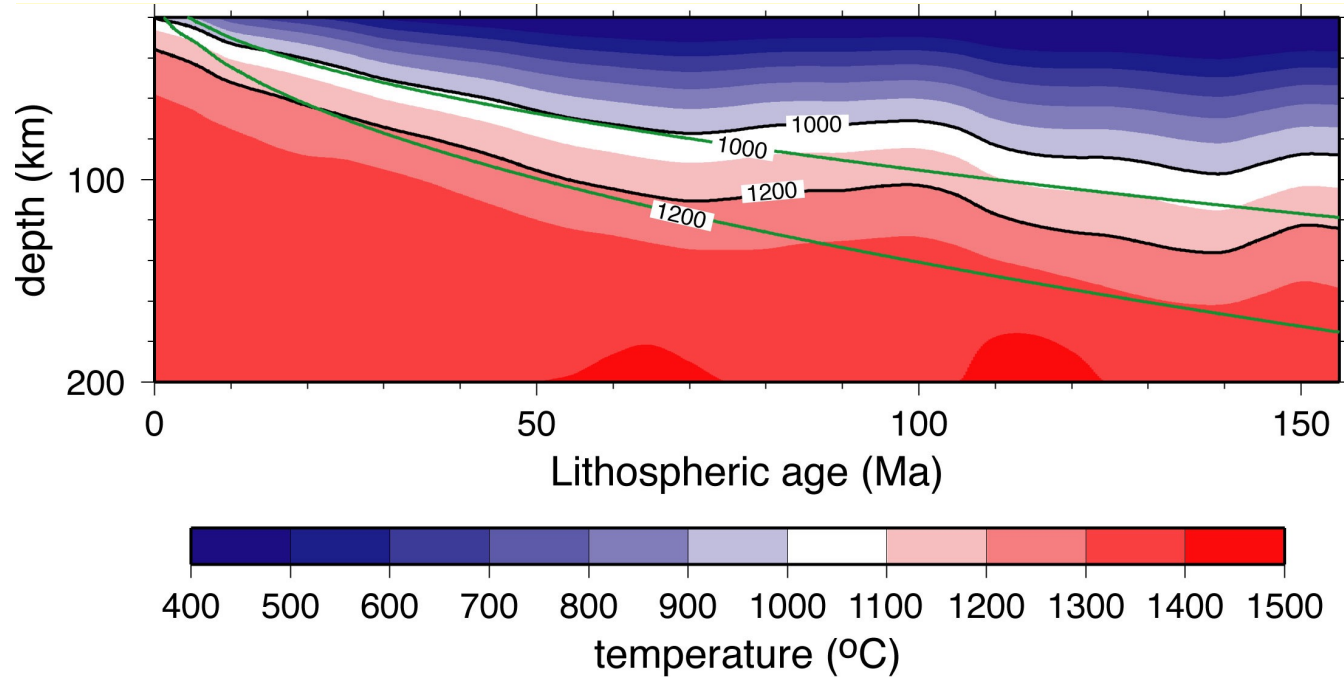




Small-Scale Convection (SSC) beneath the oceanic plates.

[Ballmer *et al.*, 2010]

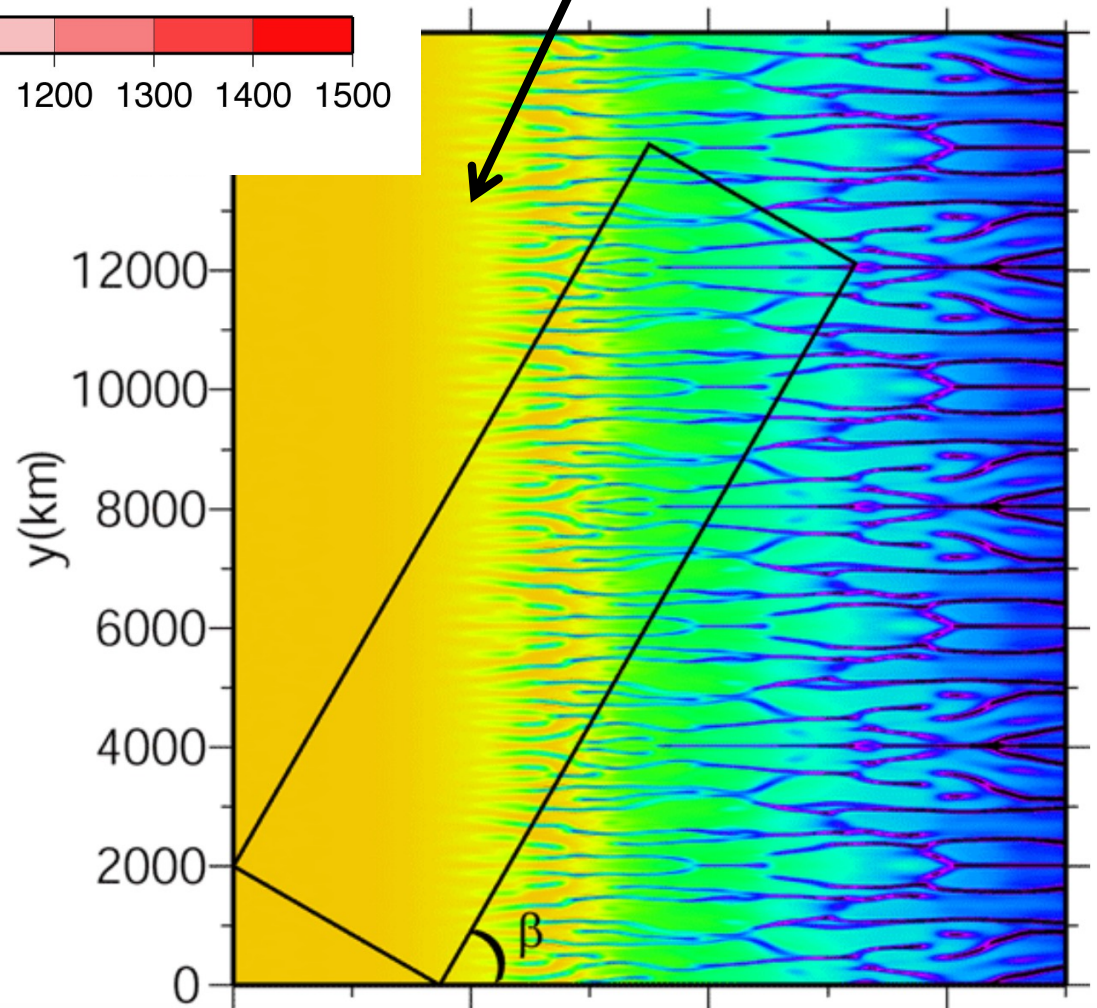
“Richter Rolls” organize the volcanism into linear trends

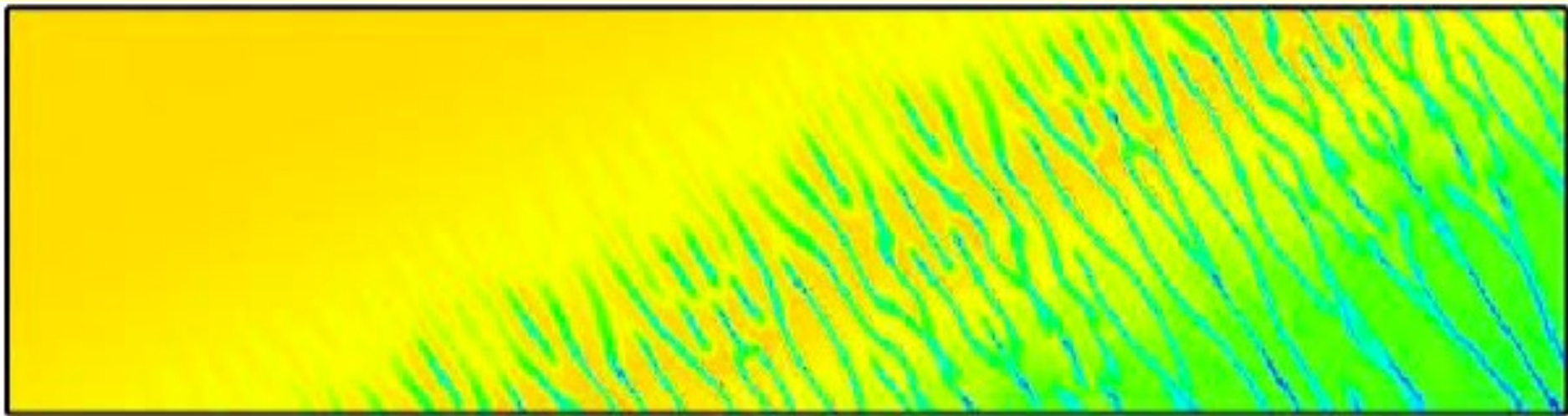


Onset of small-scale convection
 → Occurs at a critical plate age

Thermal Profile of the Pacific Lithosphere
 → Onset at ~70 Myr (?)
 [Ritzwoller et al., 2004]

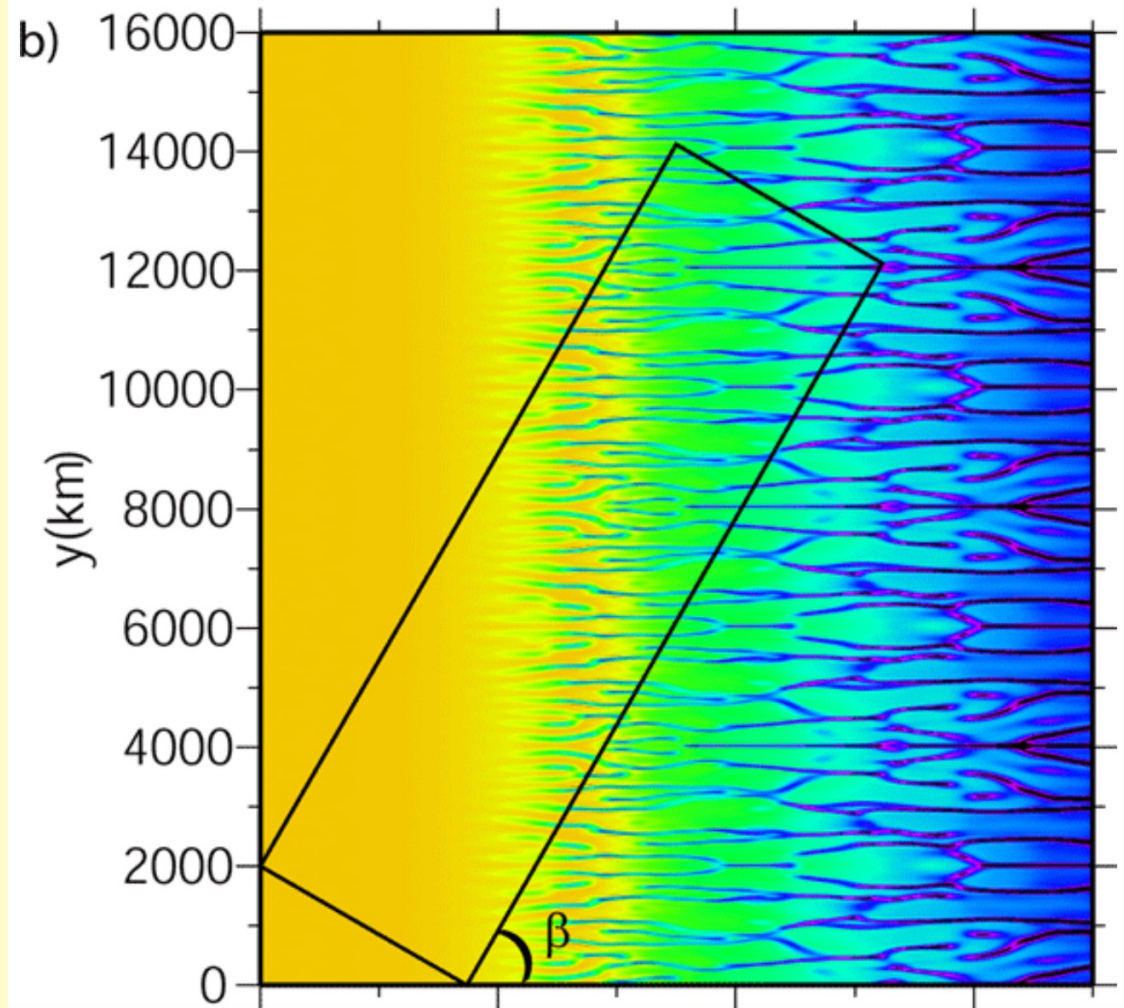
van Hunen & Zhong [2006]



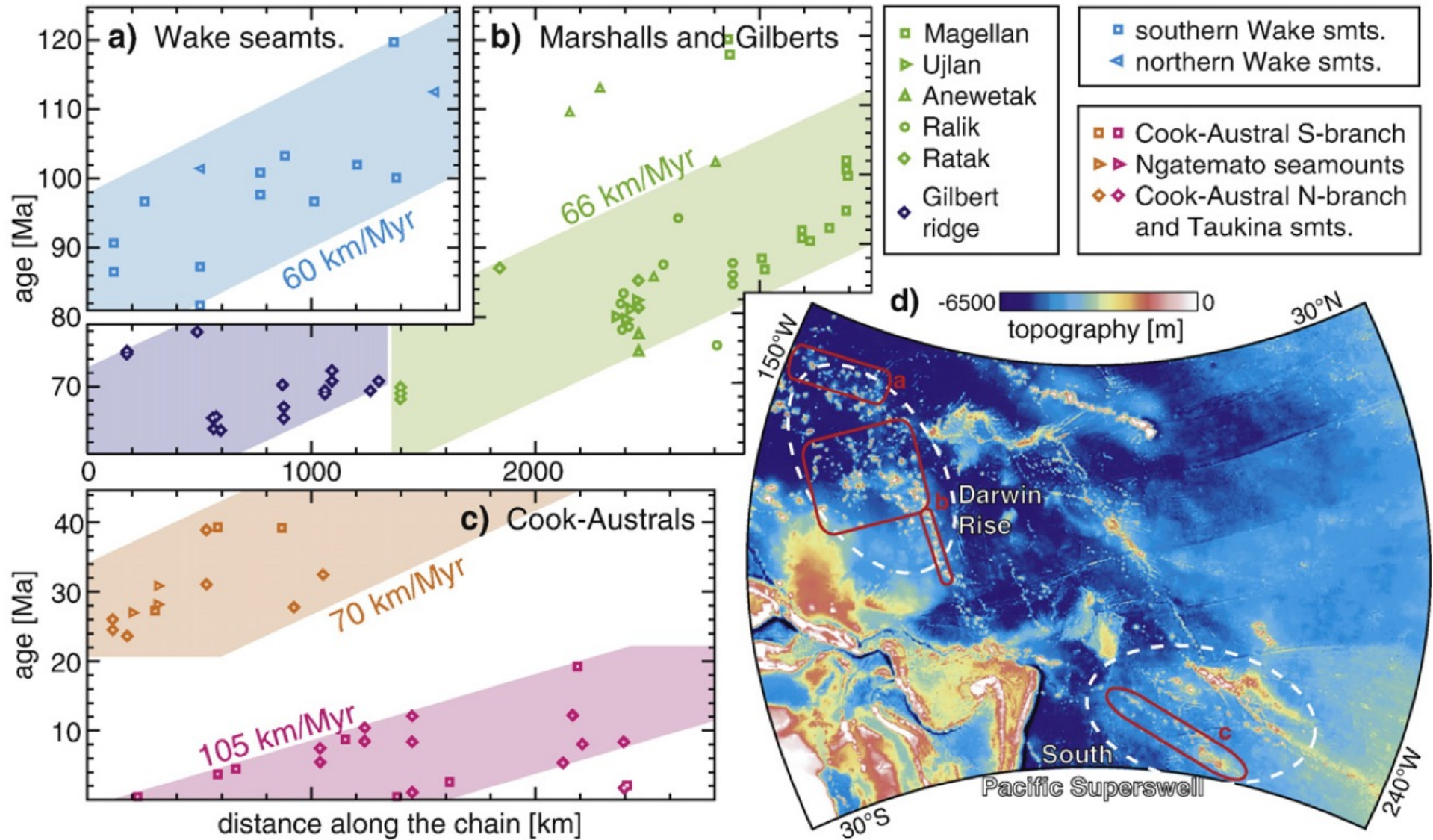


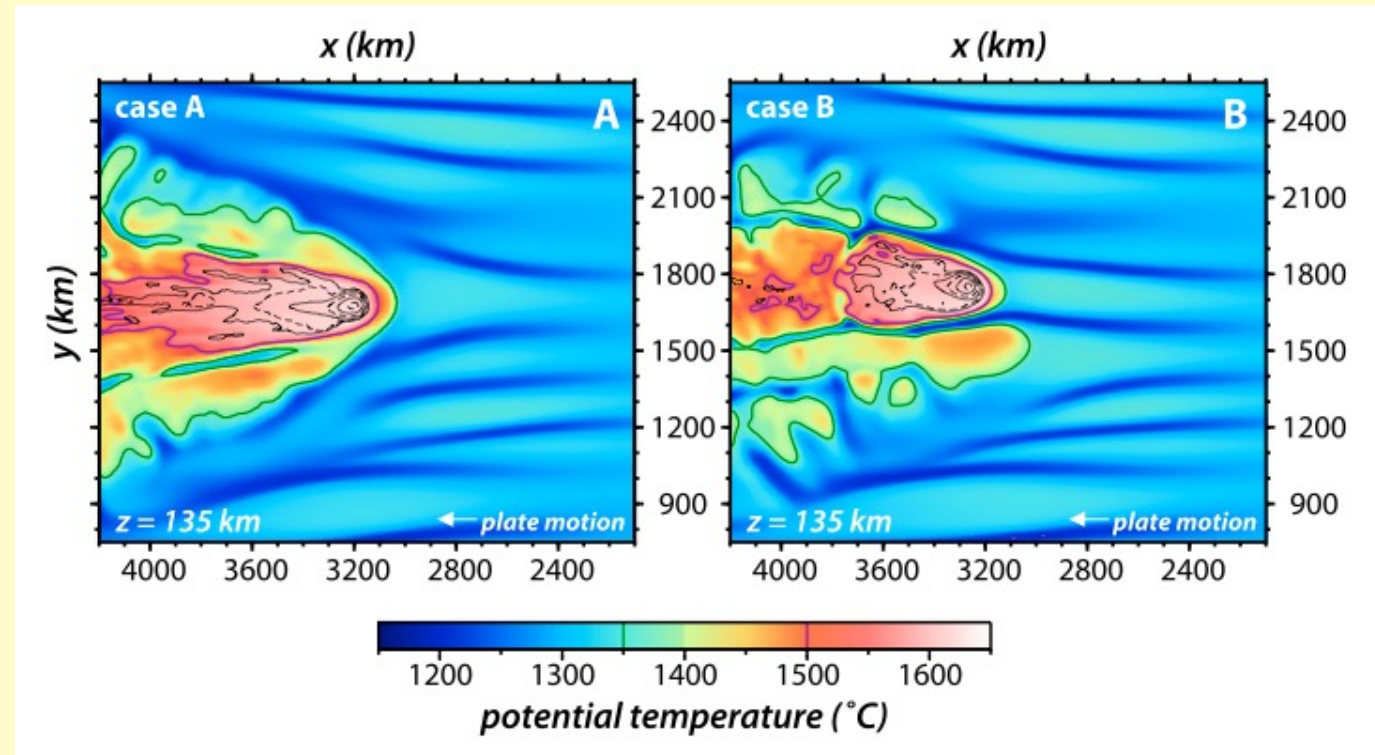
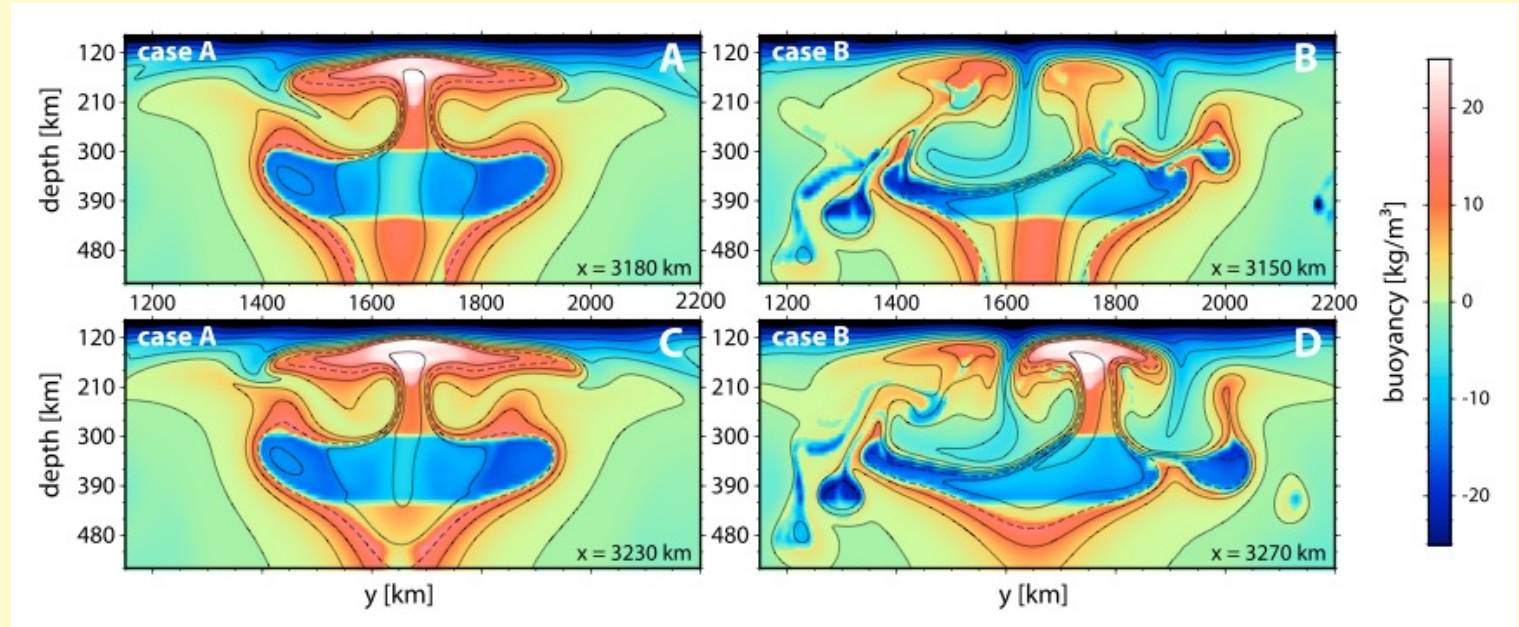
*Richter Rolls organize into
the direction of plate motion*

van Hunen & Zhong [2006]



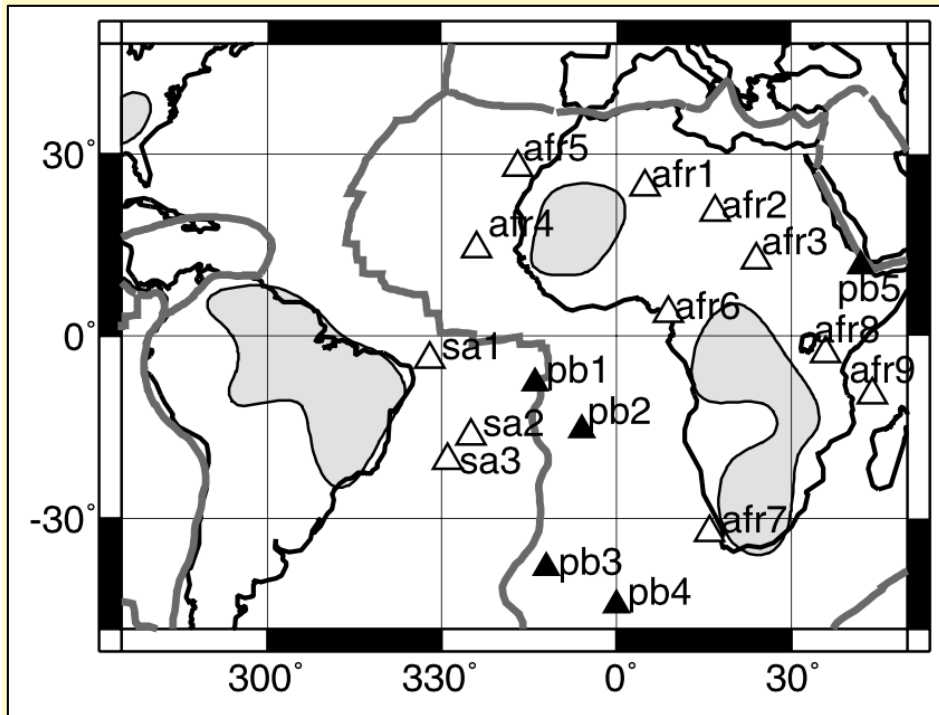
SSC may explain some mountains and minor volcanism.





Interaction of a plume with small-scale convection

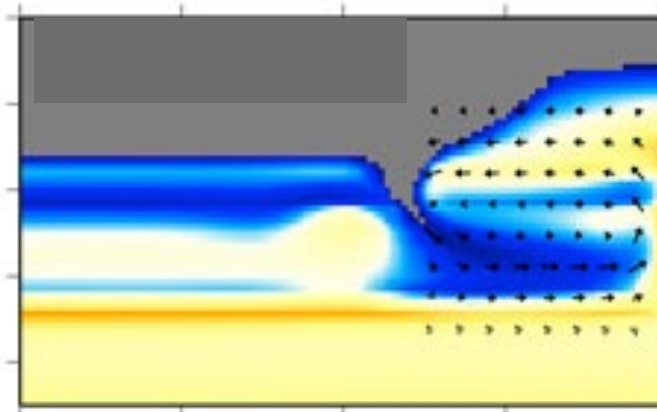
Ballmer et al., 2013



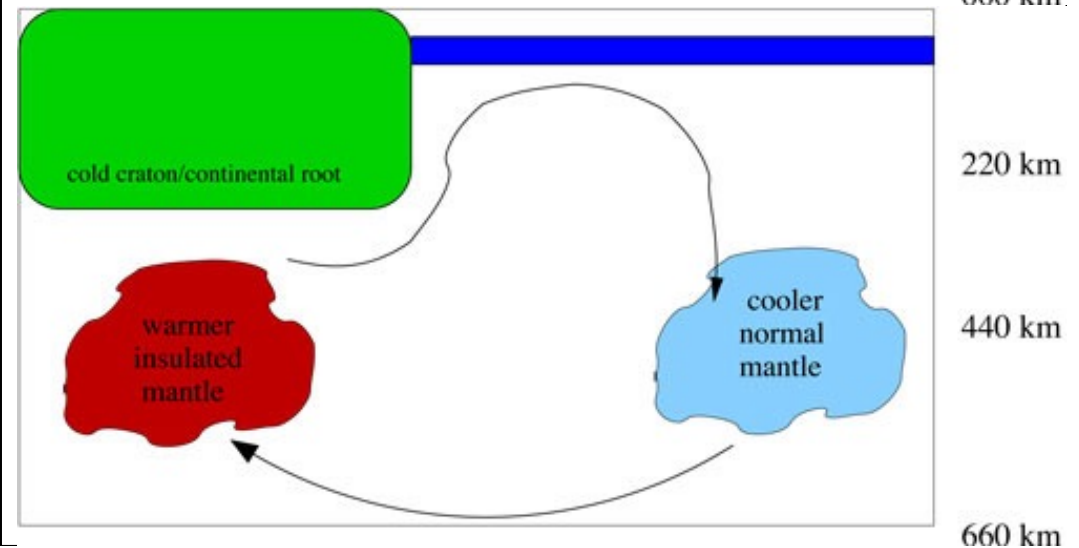
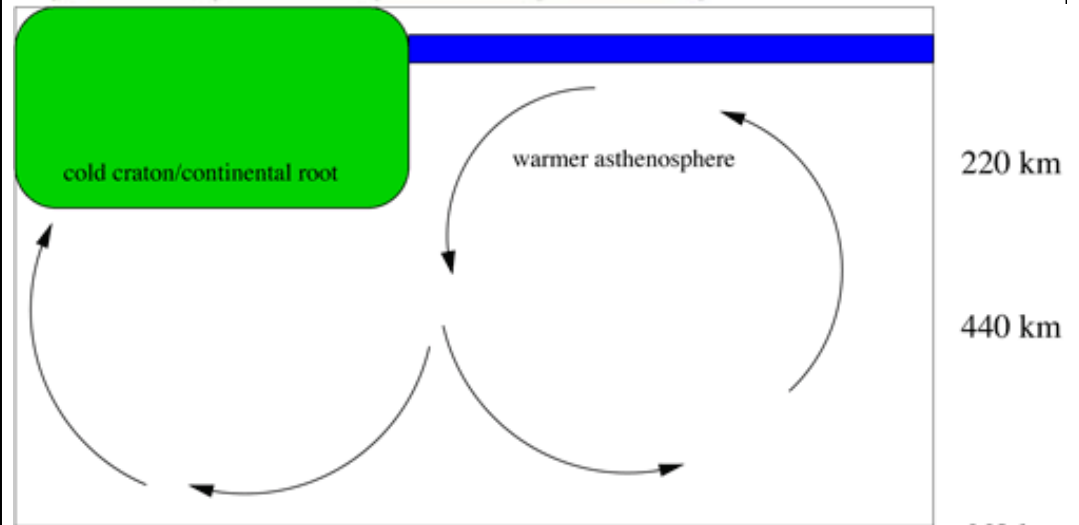
King & Ritsema [2000]

Edge-Driven Convection

The edge of a craton focuses upwelling flow that leads to volcanism.



King & Anderson [1998]



3. Shear-Driven Upwelling (SDU)

Ingredients:

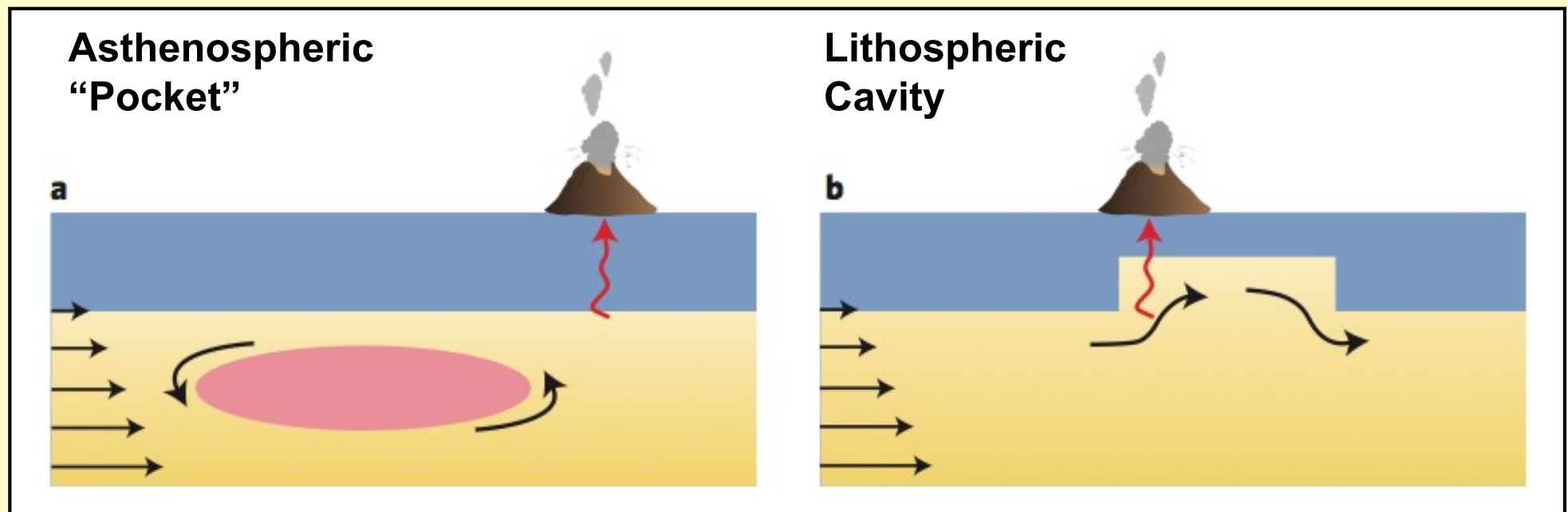
- Near-Solidus Asthenosphere
- Viscosity Heterogeneity
- Rapid Asthenospheric Shear

Conrad et al. [PEPI, 2010]

Conrad et al. [Nat. Geosci., 2011]

Bianco et al. [JGR, 2011]

Density Heterogeneity not required



King [News & Views, Nature Geoscience, 2011]

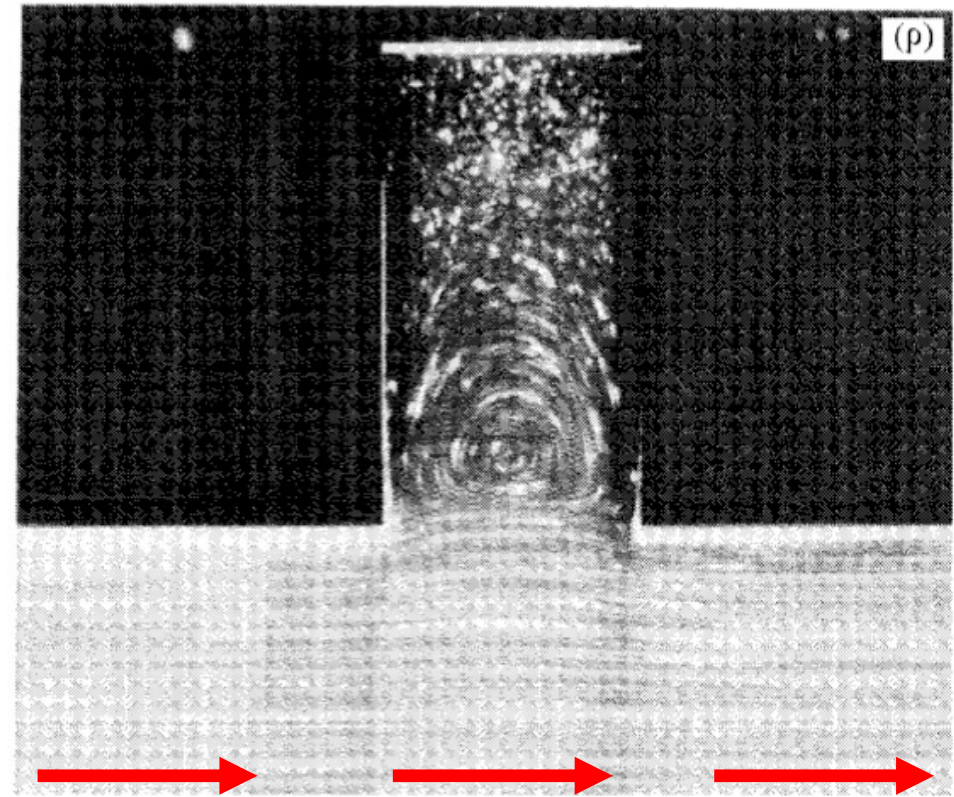
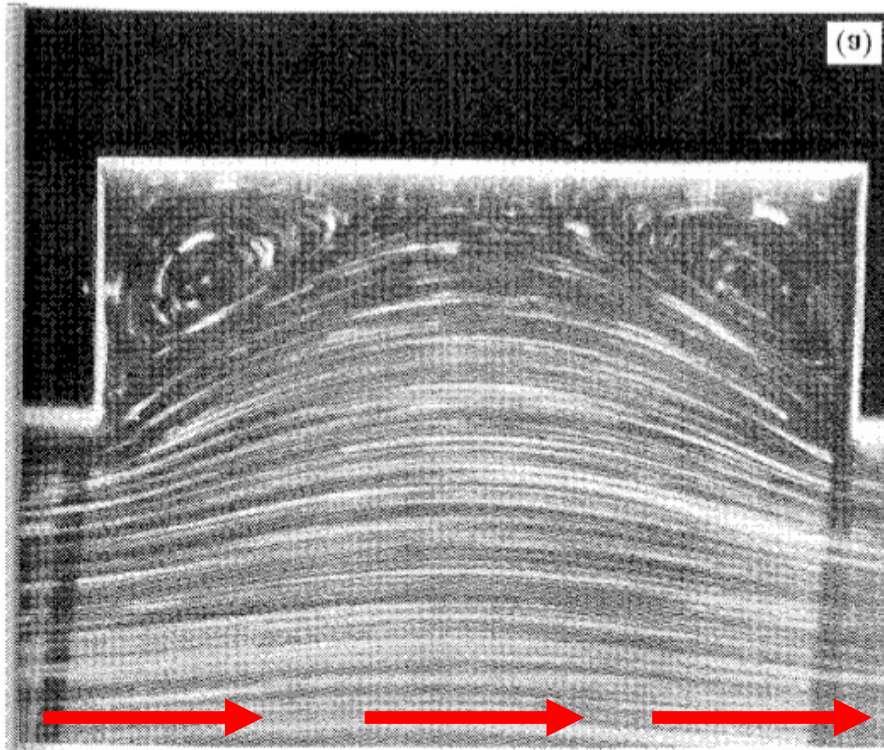
Shear-Driven Cavity Flow

A classic engineering problem

Industrial Applications:

- Spin Coating
- Mixing Processes
- Liquid cooling by melt spinning
- Benchmark for computational schemes

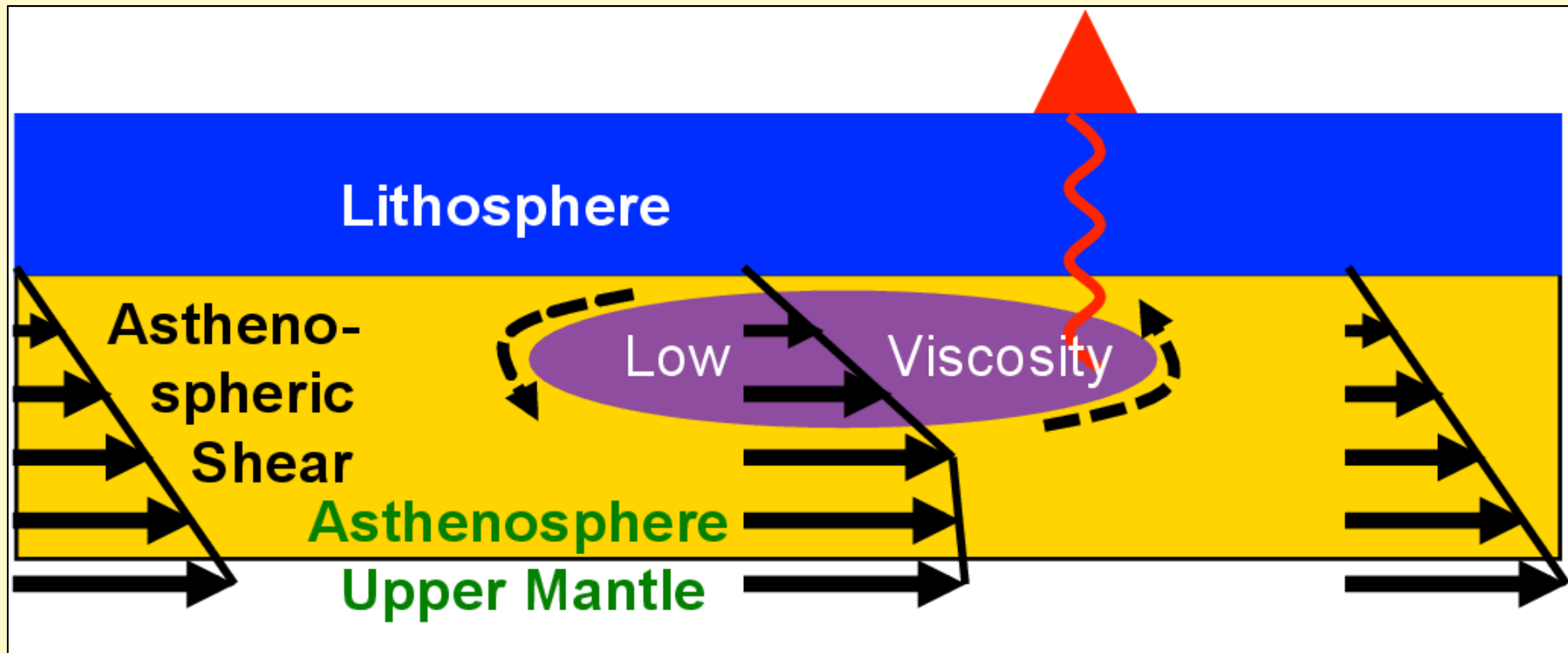
Shear-Driven Cavity Flow



[Taneda, 1979]

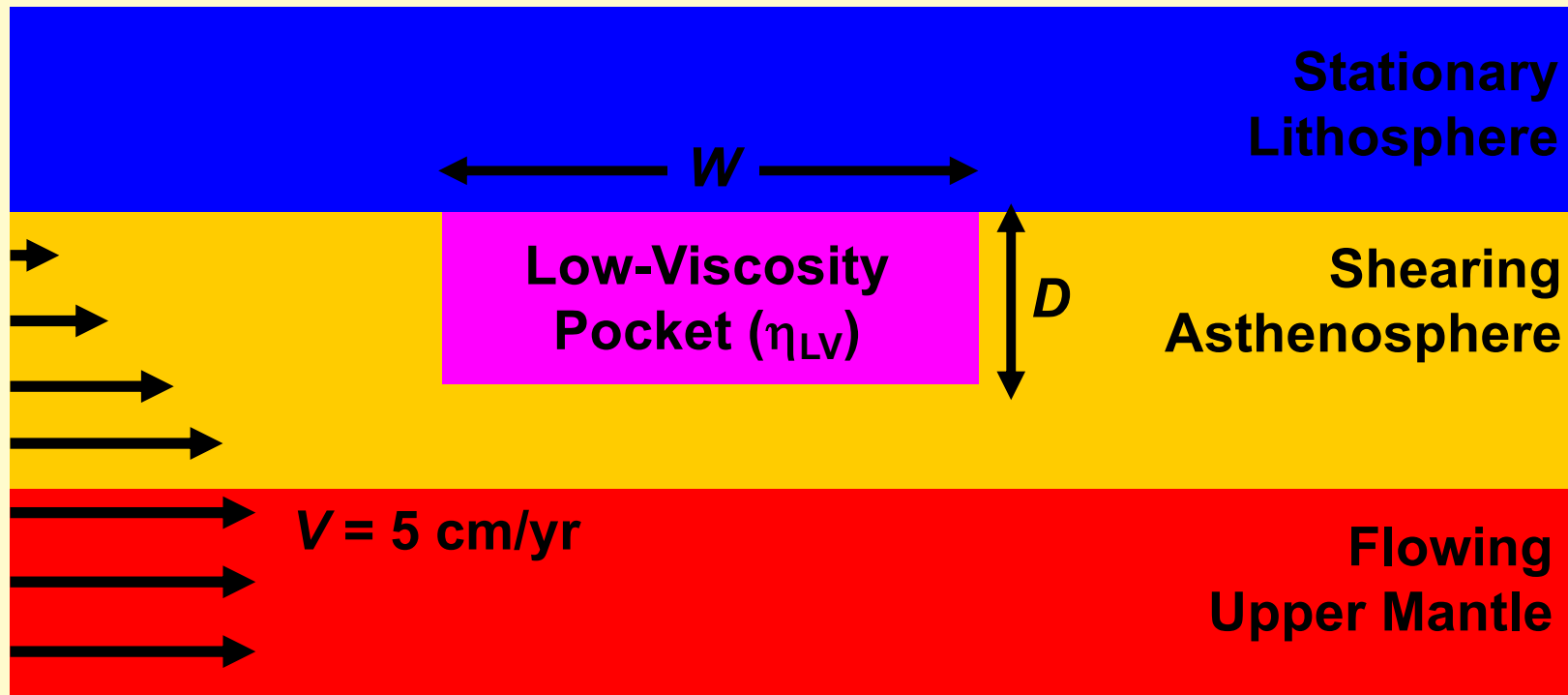
Shear-Driven Upwelling within an Low-Viscosity Pocket

- A previously unstudied variation of shear-driven cavity flow:
Flow within a low-viscosity “pocket”



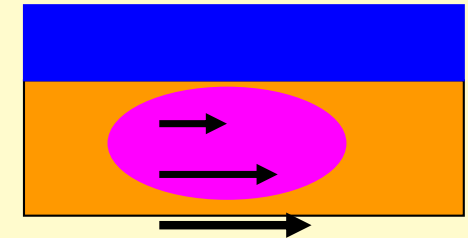
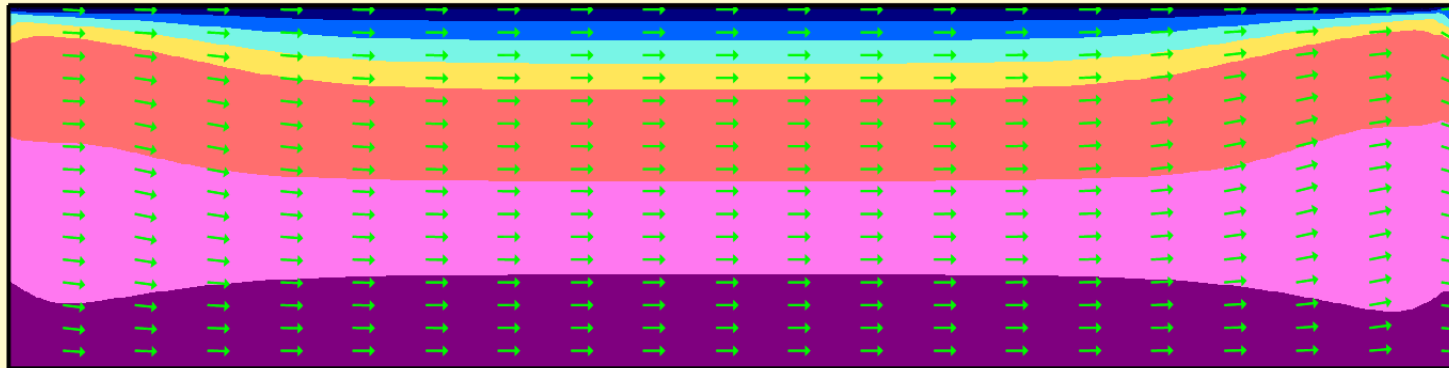
Measuring SDU within an Low-Viscosity Pocket

- Flow simulation of viscous shear in 2D
- Vary the dimensions and viscosity of the pocket
- Measure maximum upwelling

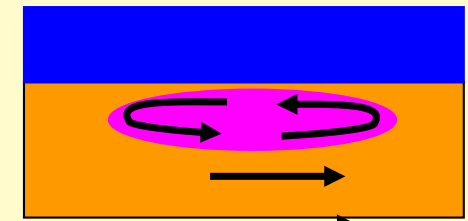
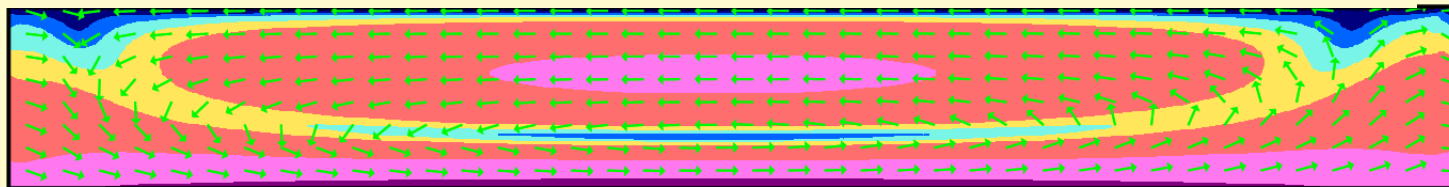


Flow patterns within a low-viscosity “pocket”

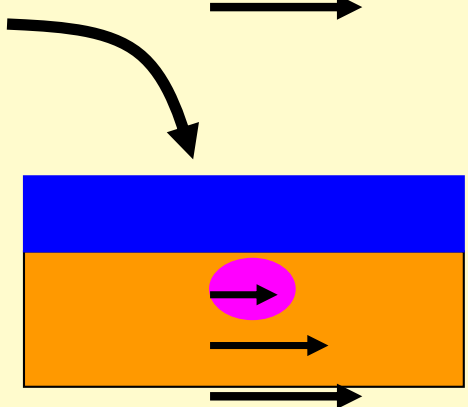
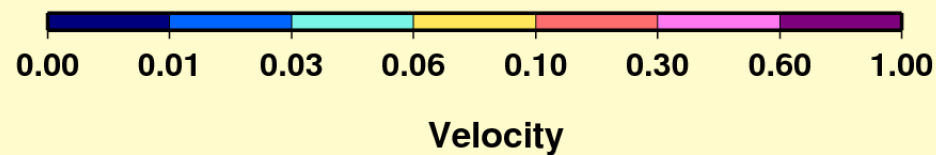
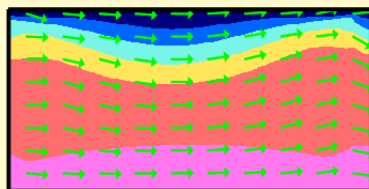
A) Thick pocket: Shear develops within the pocket

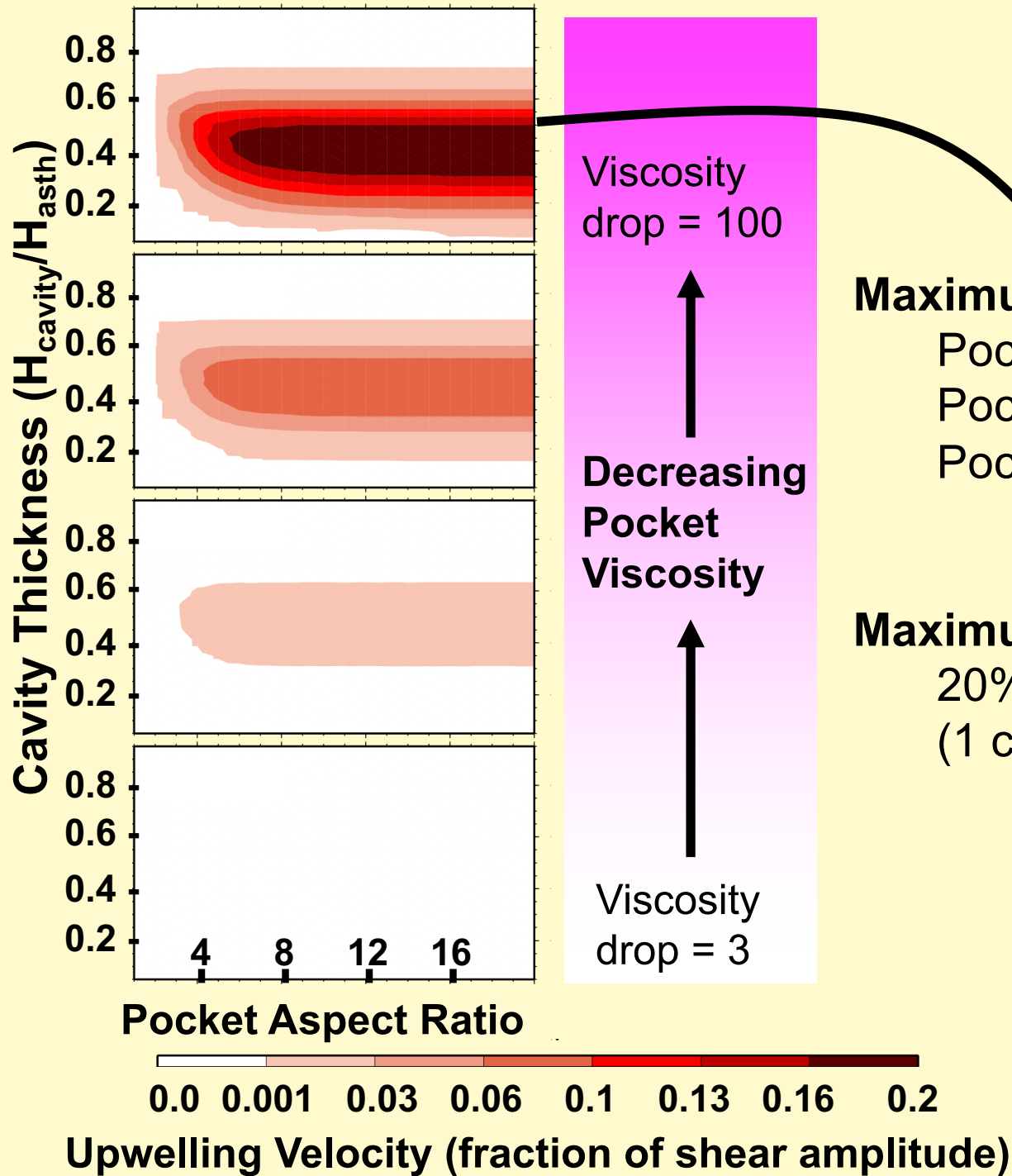


B) Shallow & wide pocket: Circulation within the pocket



C) Shallow & narrow pocket: Pocket rides along with flow





Viscosity drop = 100

Decreasing Pocket Viscosity

Viscosity drop = 3

Maximum upwelling occurs for:

- Pocket viscosity drop of 100
- Pocket aspect ratio > 5
- Pocket occupies the upper 20-60% of asthenosphere

Maximum rate of upwelling:

- 20% of V
- (1 cm/yr)

Where is asthenospheric shear largest?

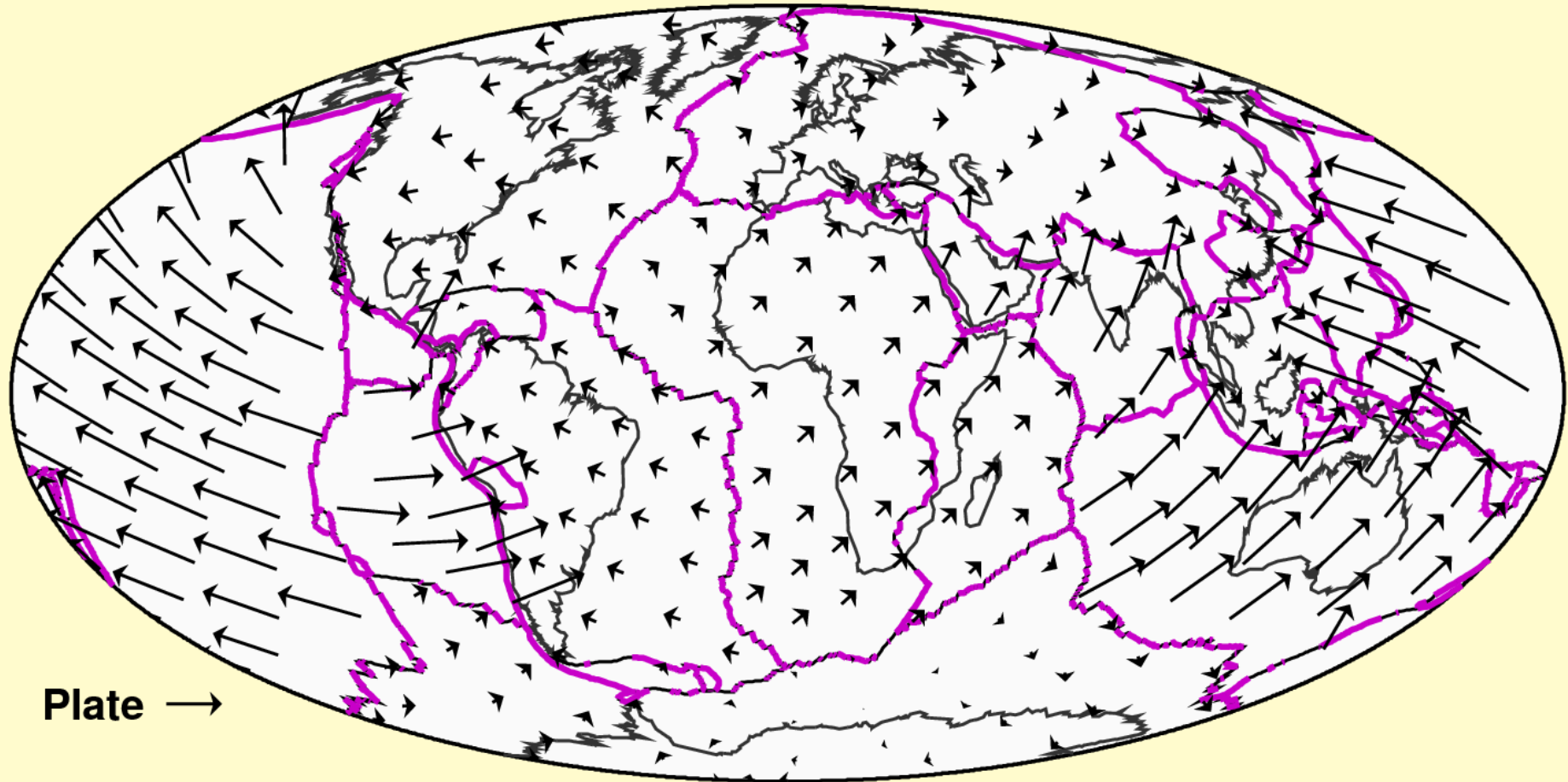
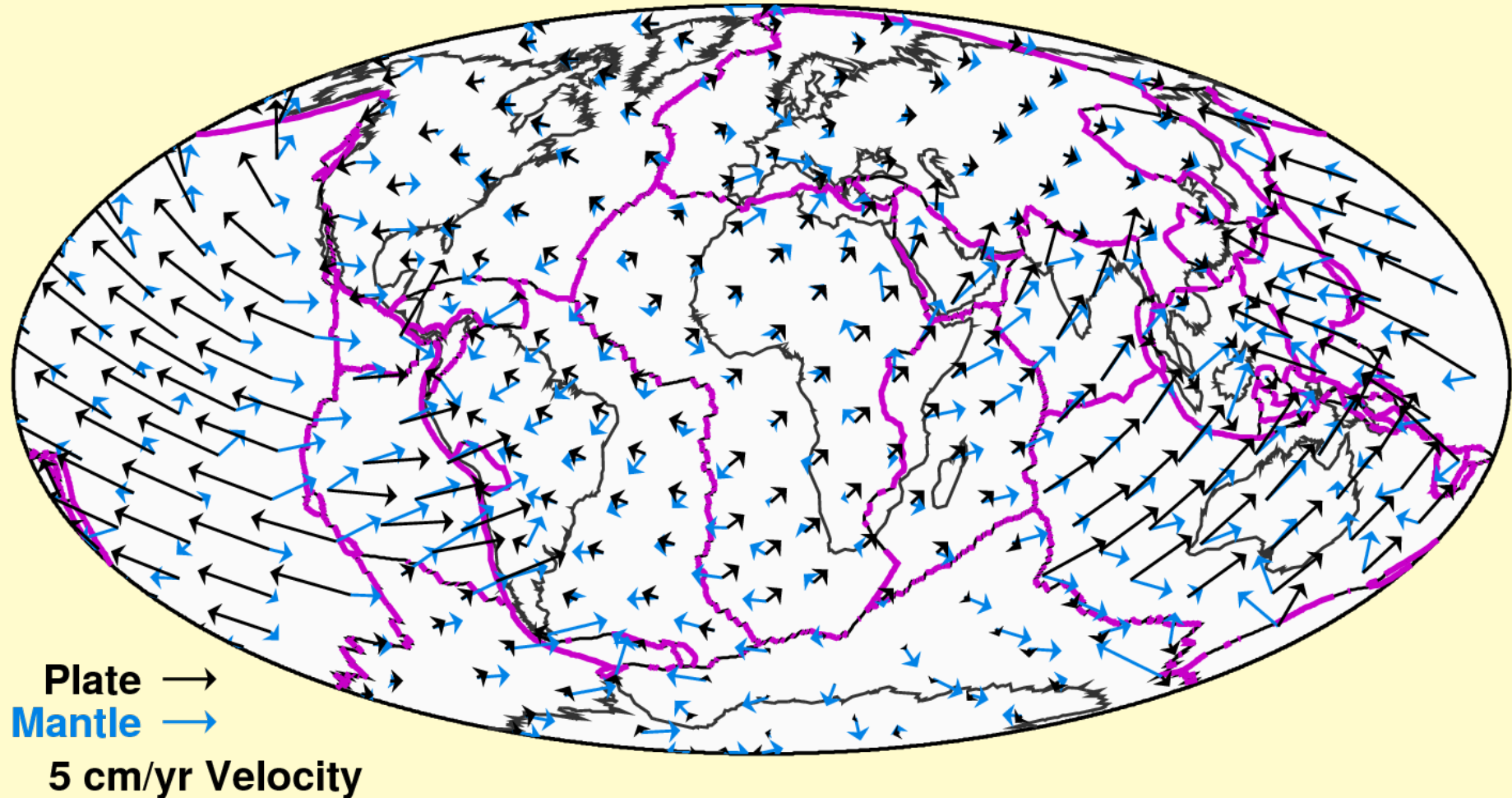


Plate →

5 cm/yr Velocity

Surface Plate Motions

Where is asthenospheric shear largest?



Surface Plate Motions and Mantle Flow at 300 km

Mantle flow field from *Conrad and Behn* [2010]

Where is asthenospheric shear largest?

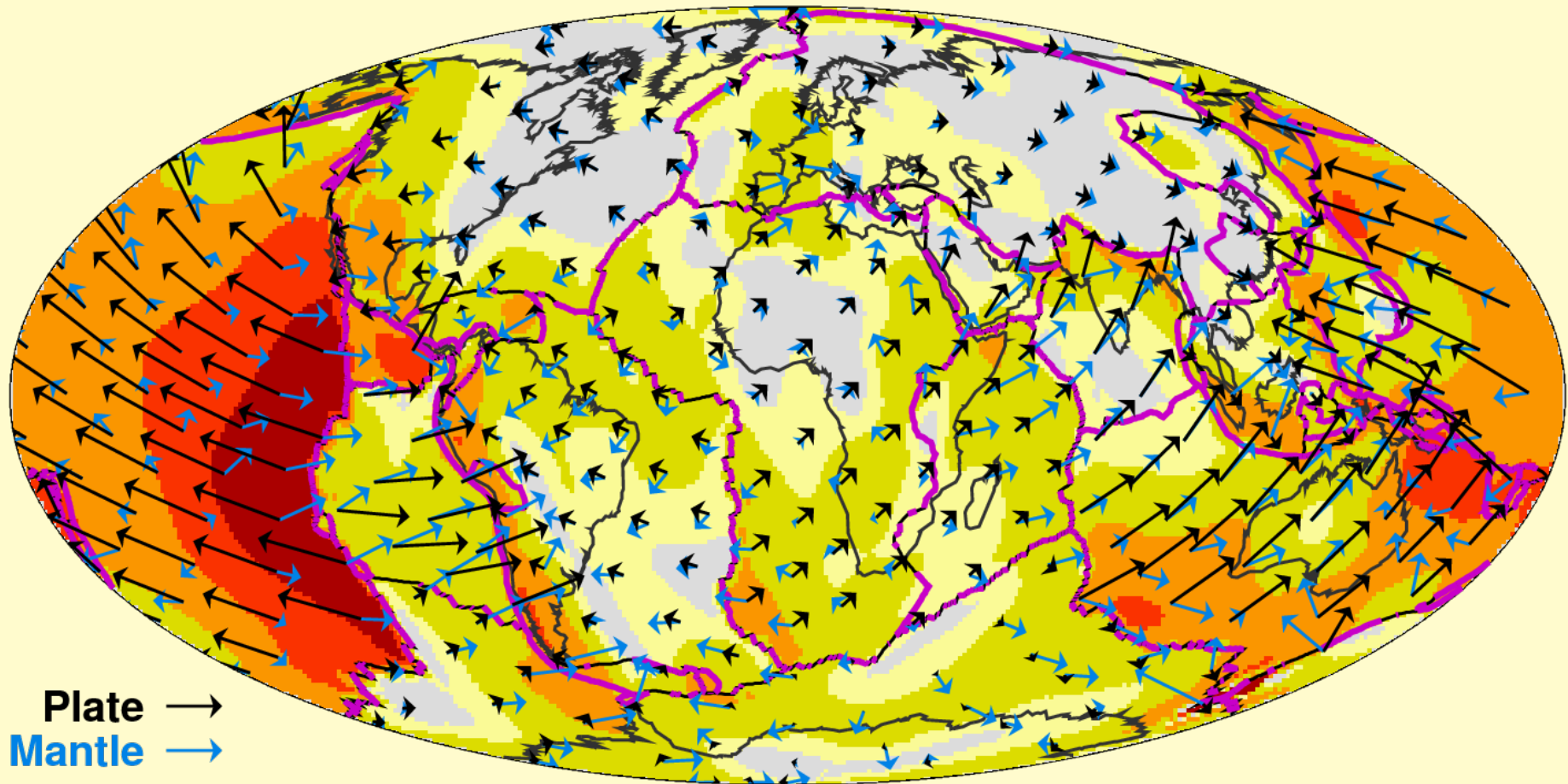
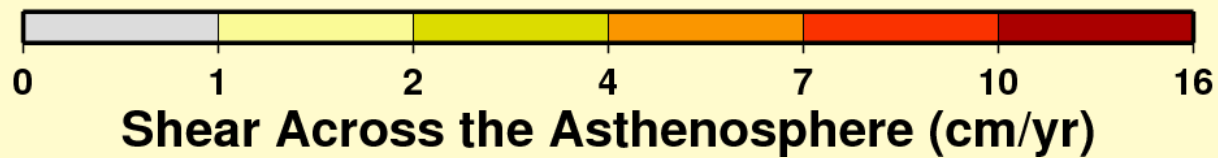
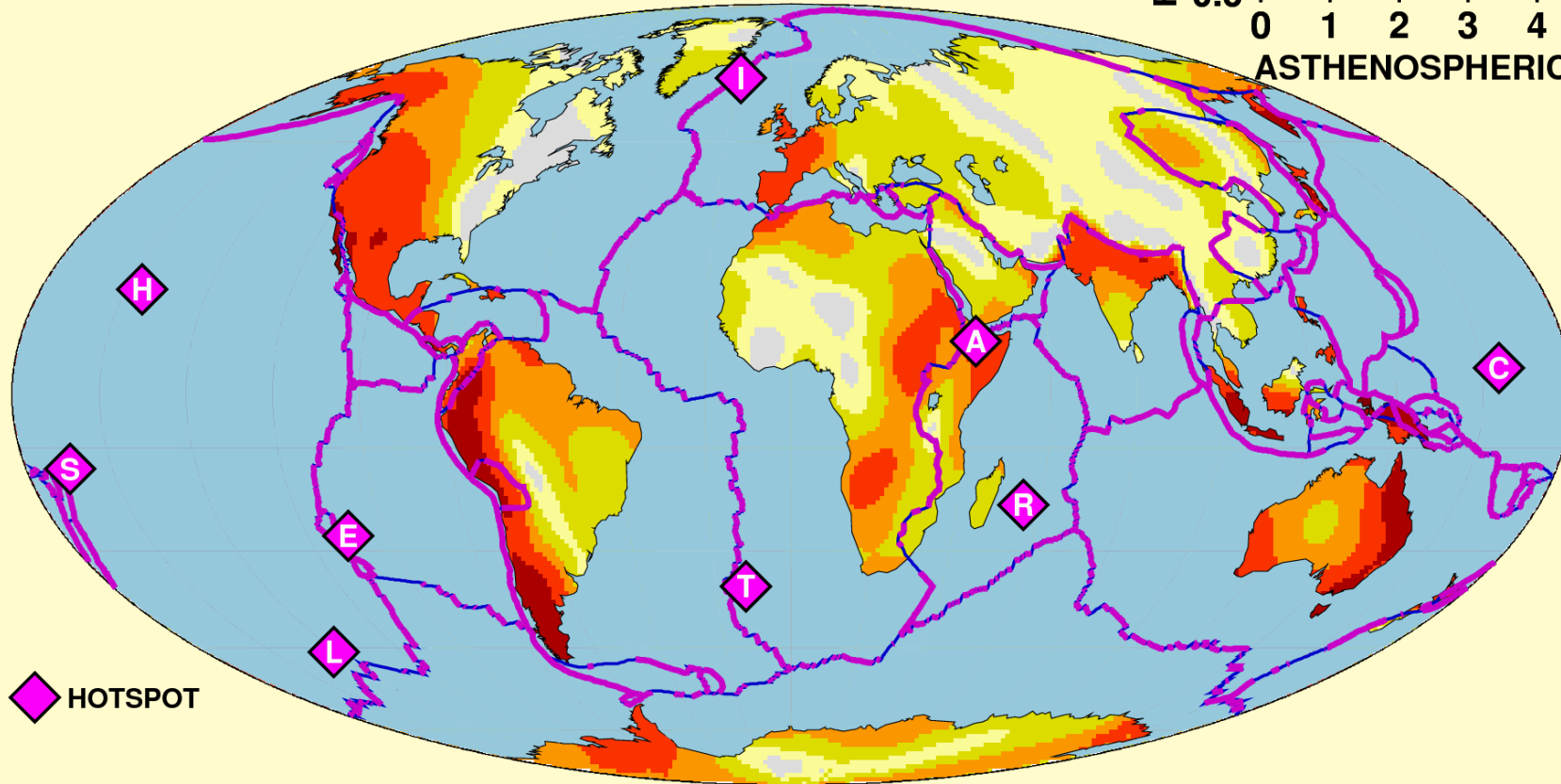
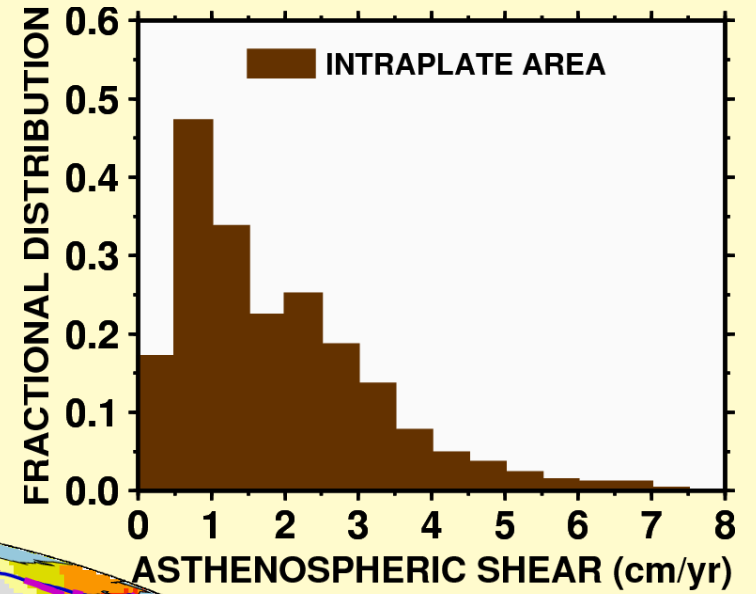


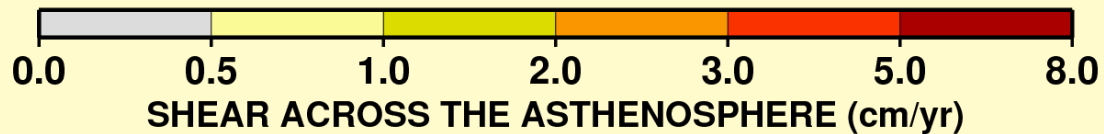
Plate →
Mantle →
5 cm/yr Velocity



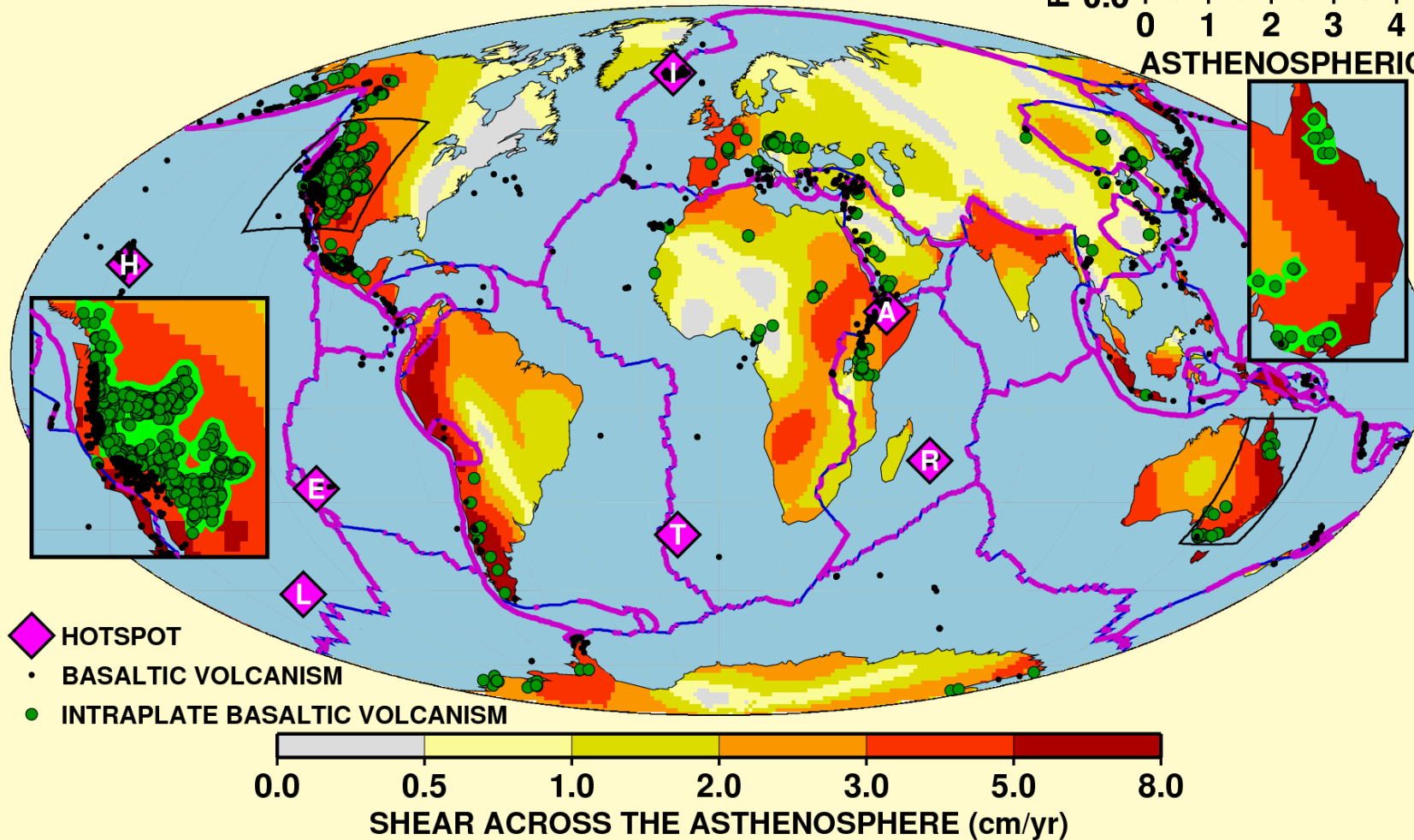
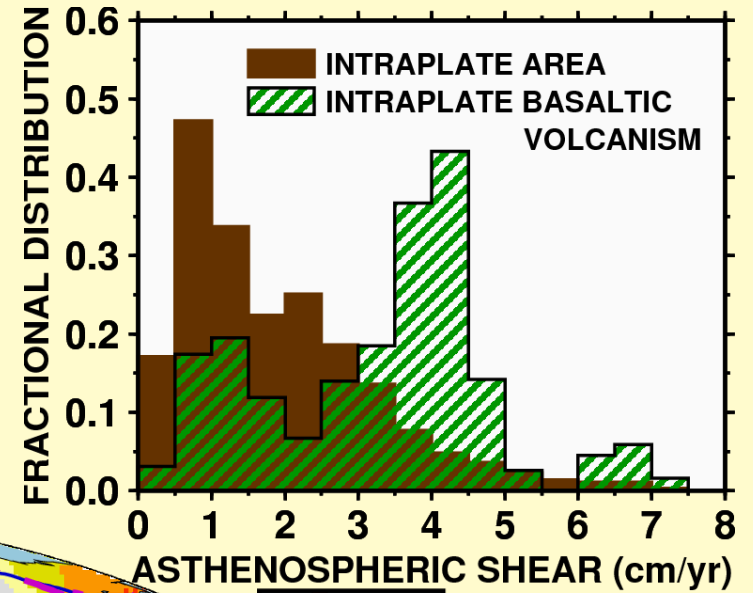
Asthenospheric Shear Beneath Continents



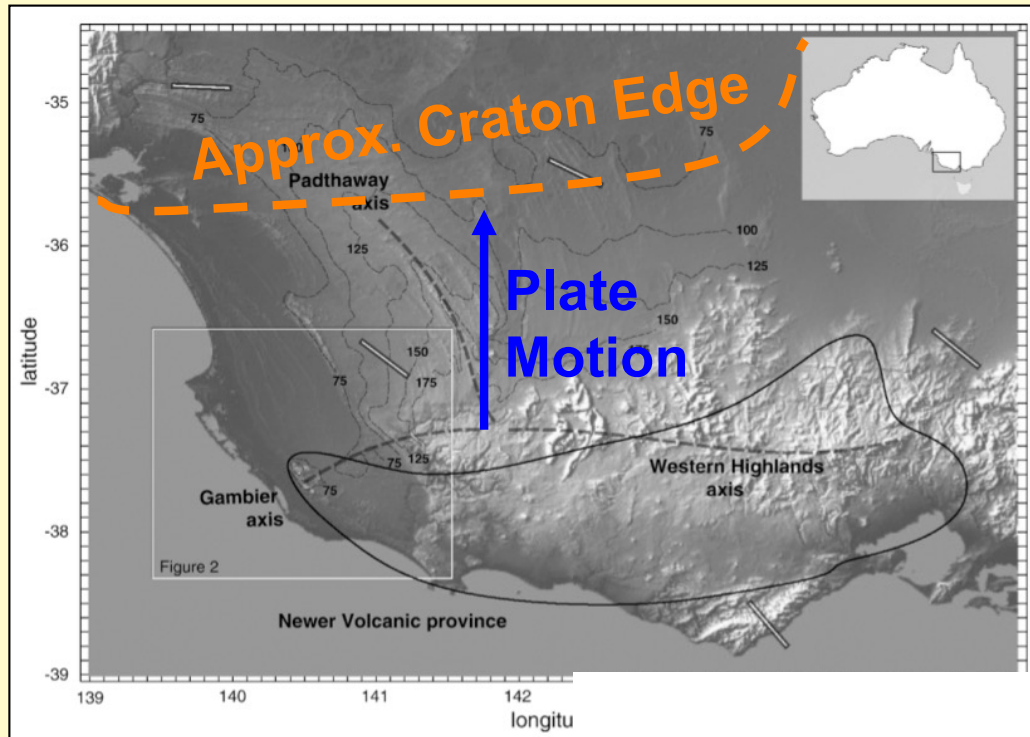
 HOTSPOT



Asthenospheric Shear and Continental Volcanism

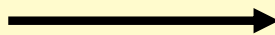


The Newer Volcanic Province (South Australia)

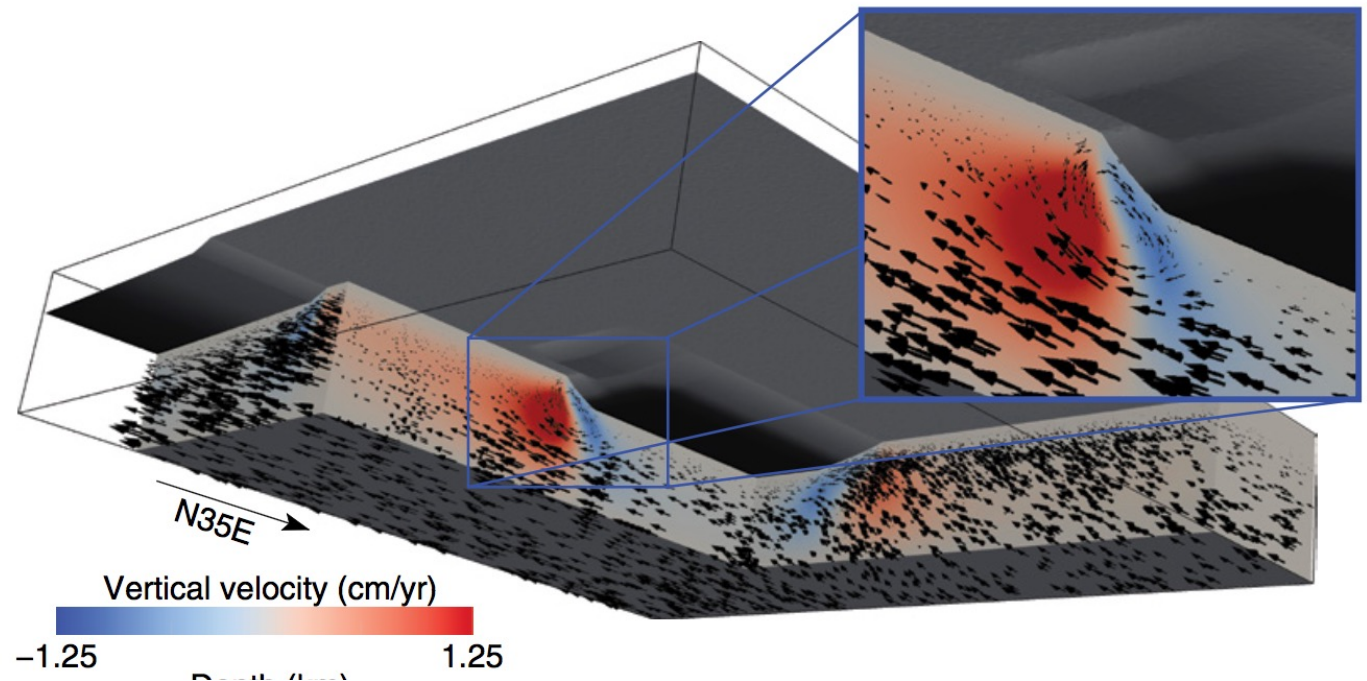


Mt. Gambier, Erupted 4500 yrs ago

SDU-like model

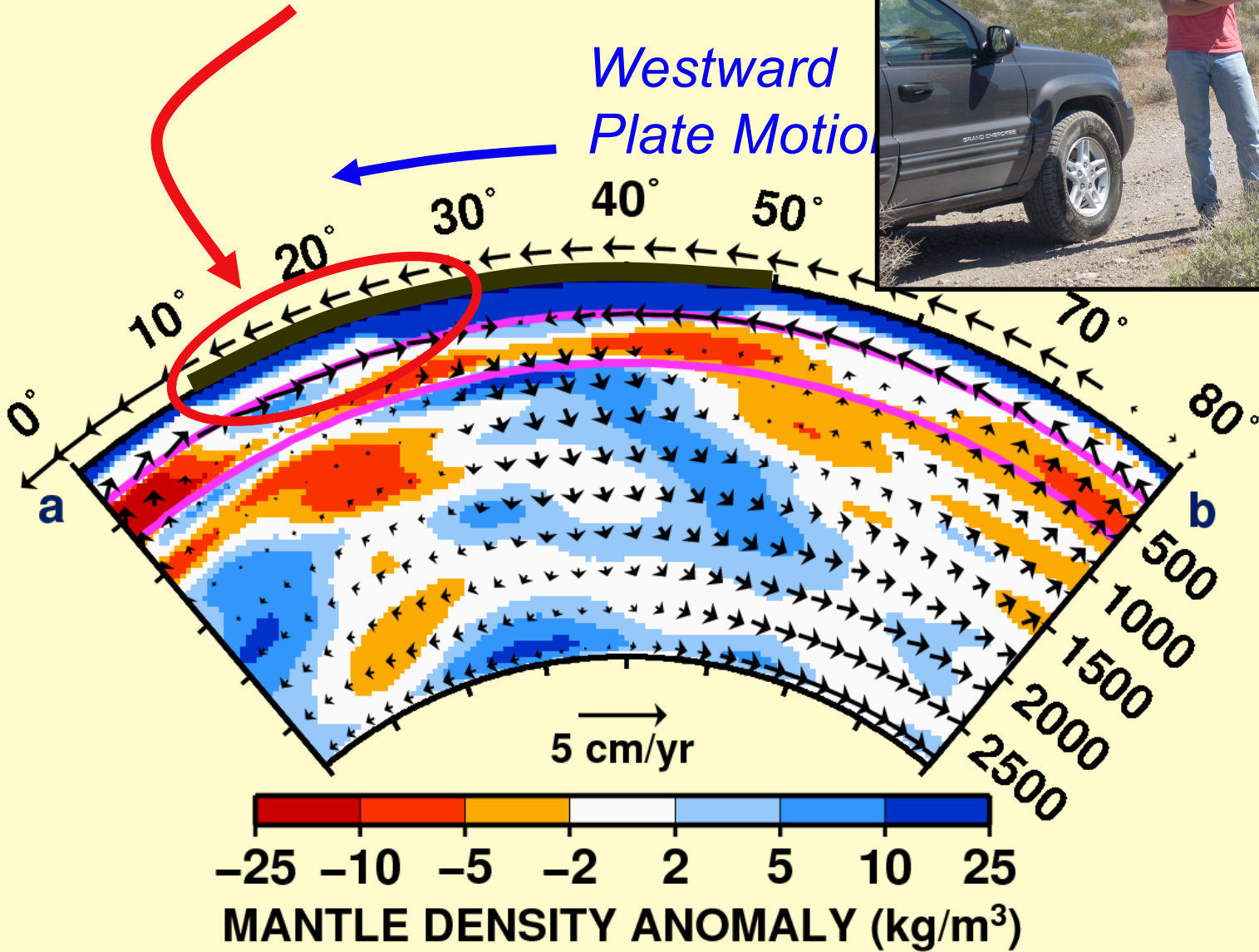


Davies & Rawlinson [2014]



Asthenospheric Shear beneath the Western US

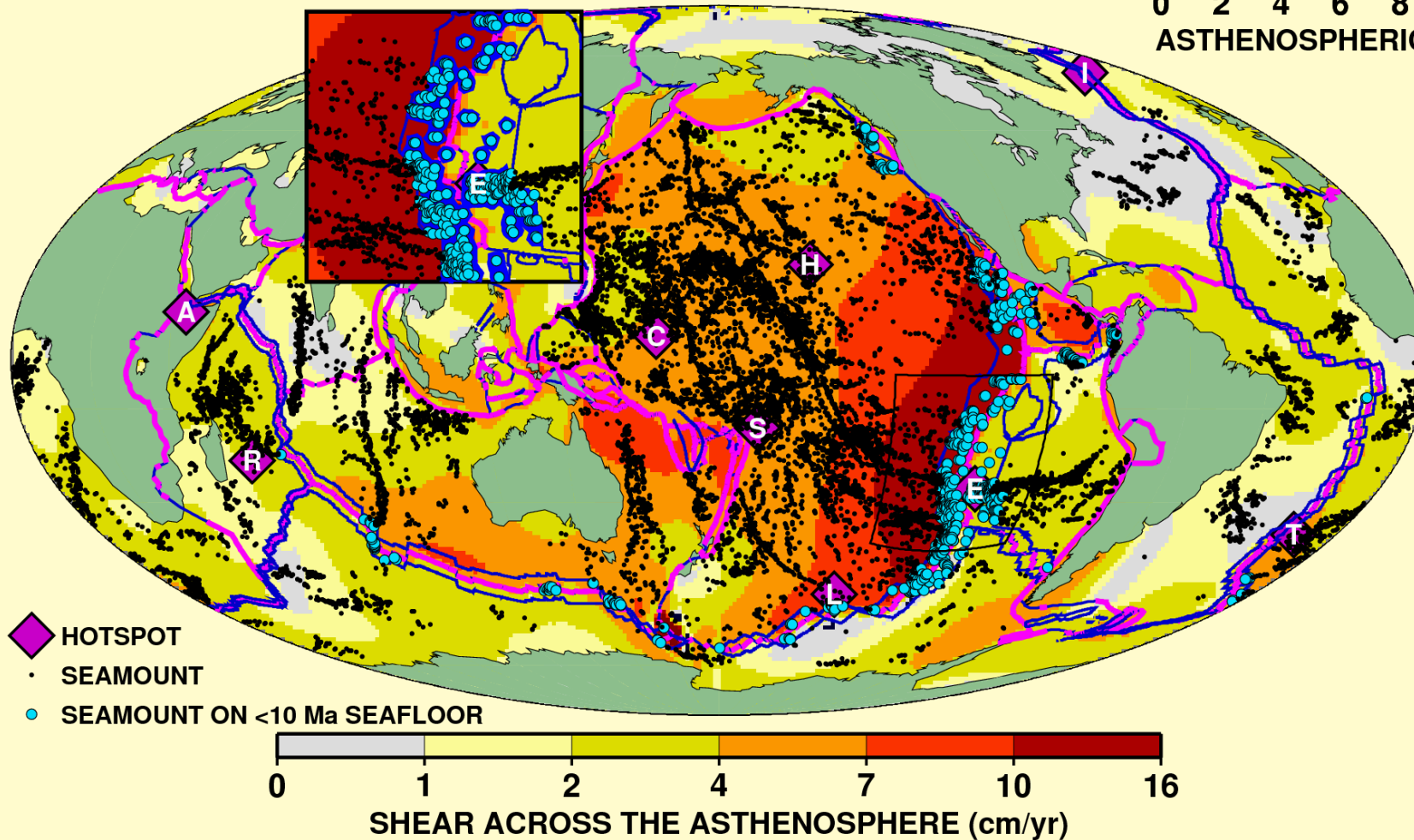
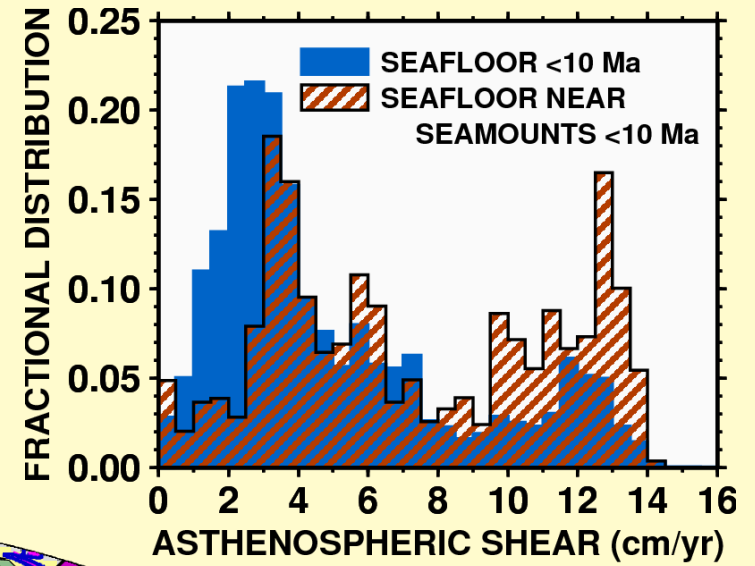
About 5 cm/yr shear

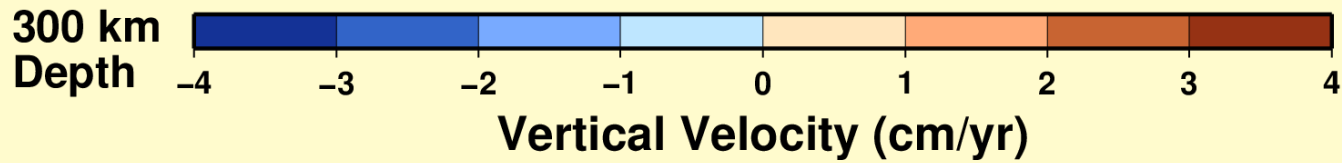
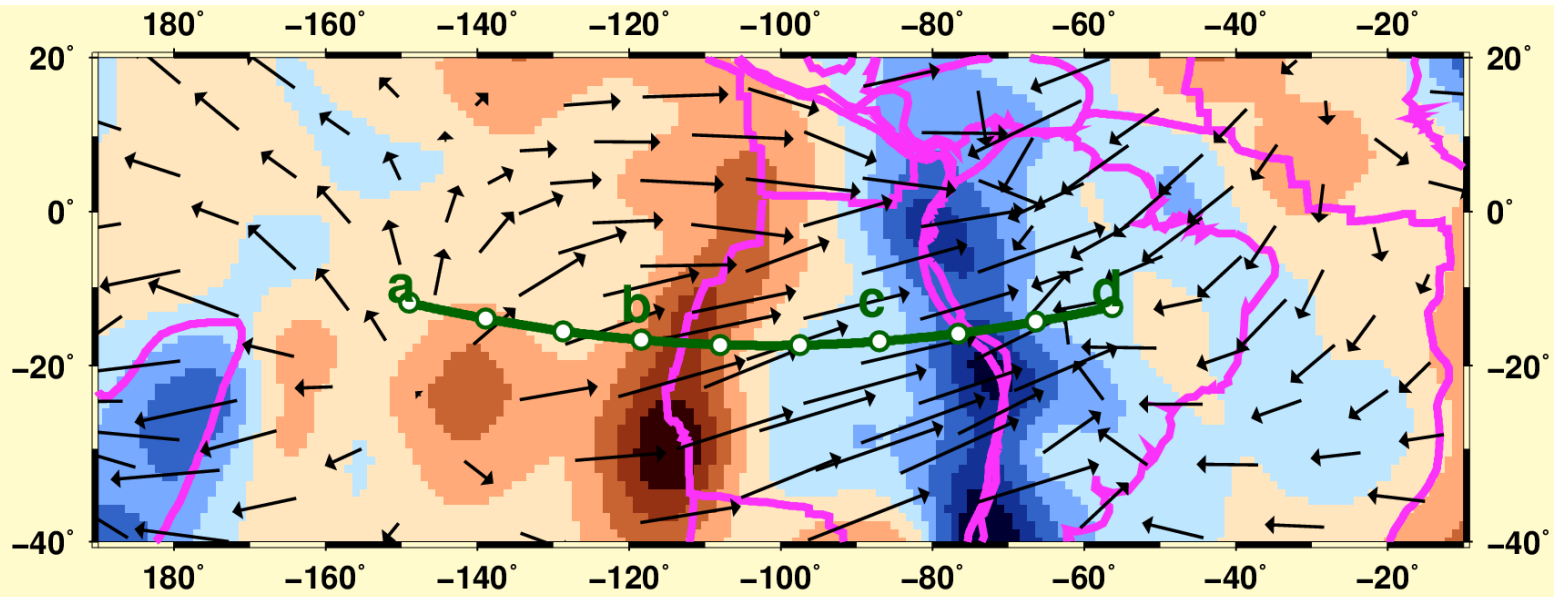


Crater Flat, Southern Nevada

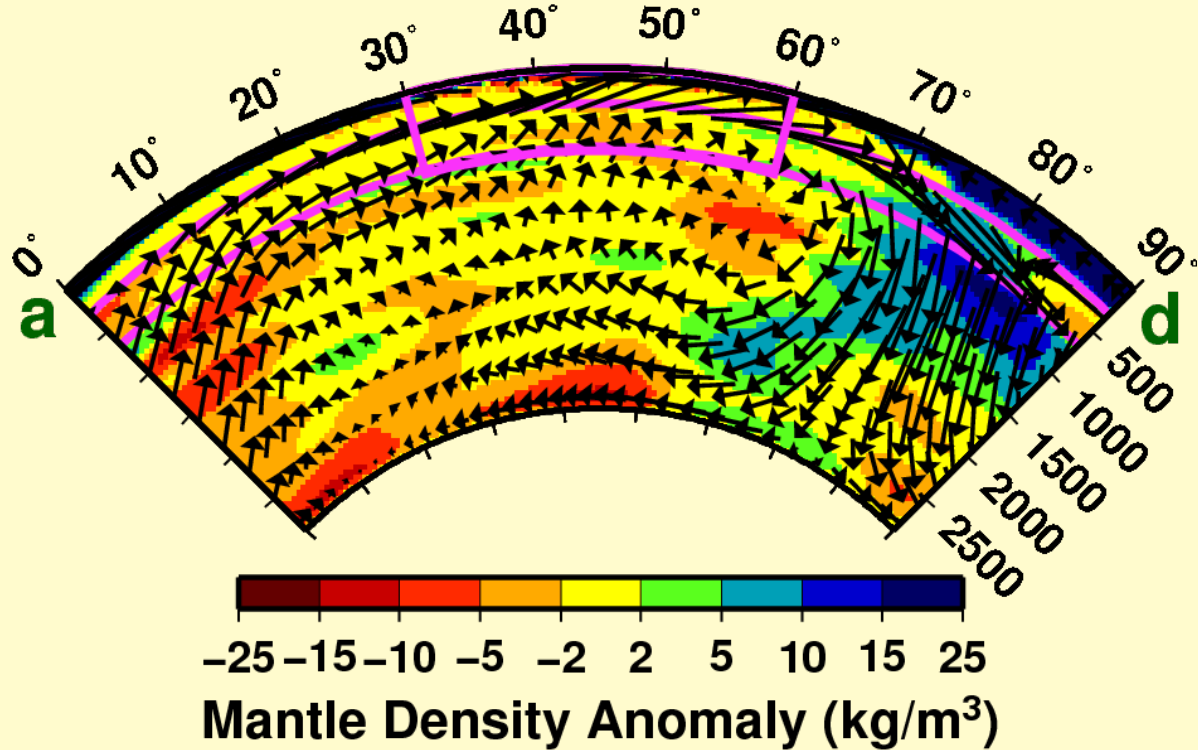


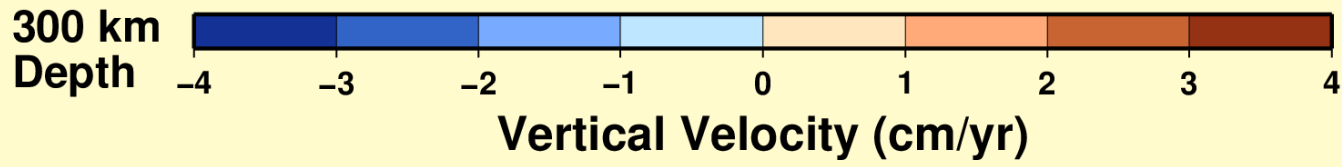
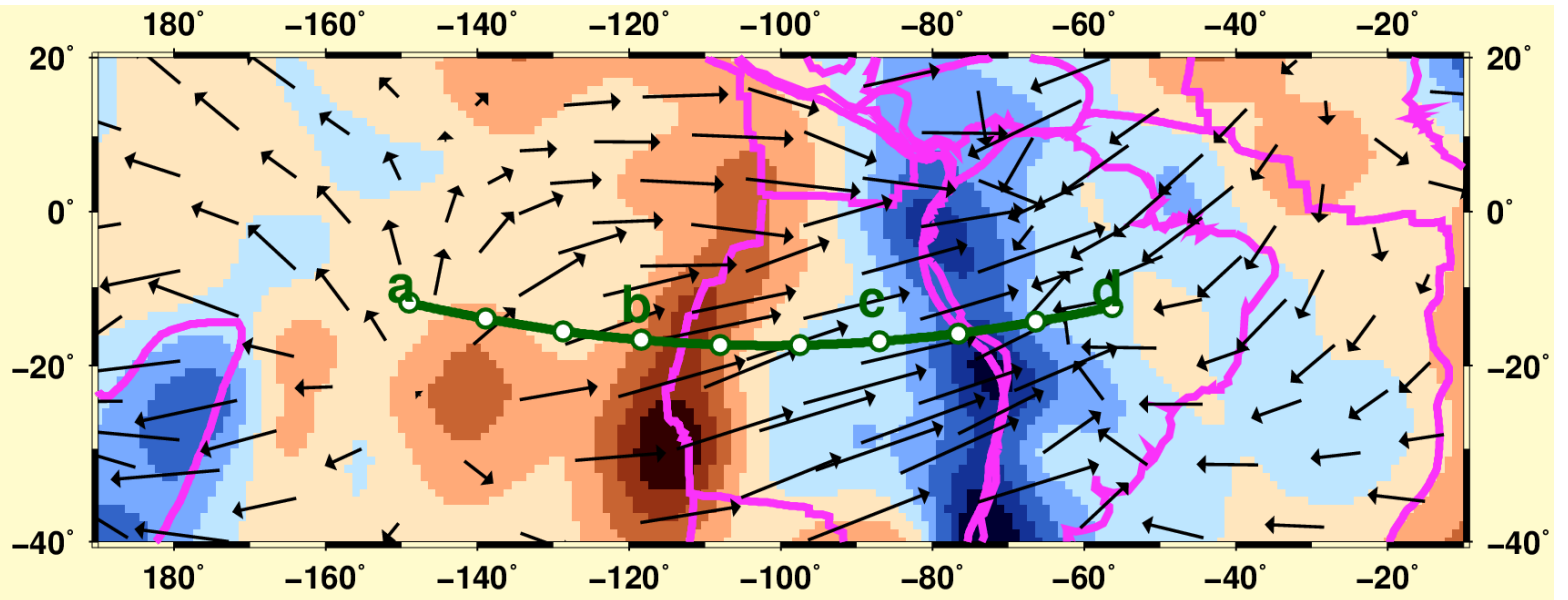
Seamount Locations and Asthenospheric Shear Beneath Seafloor



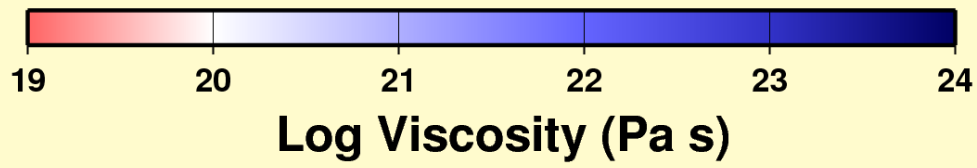
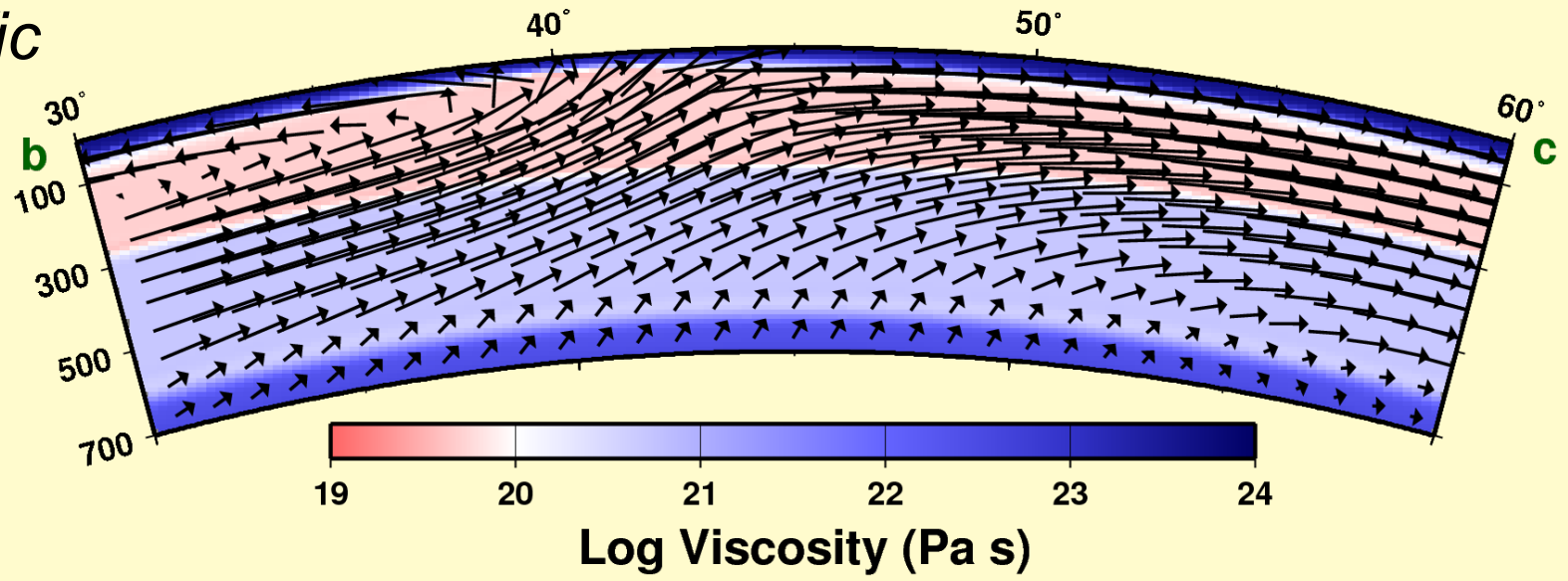


*Mantle Flow
 beneath the
 South Pacific*





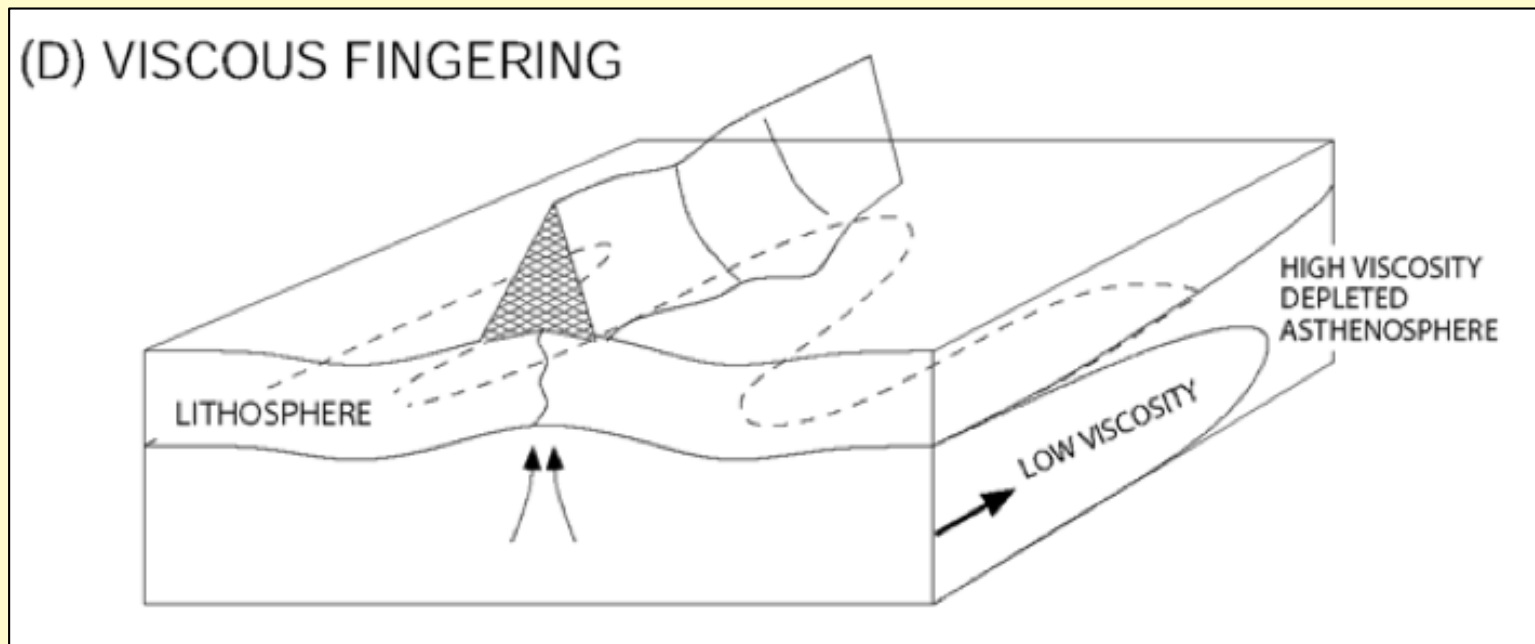
*Mantle Flow
 beneath the
 South Pacific*





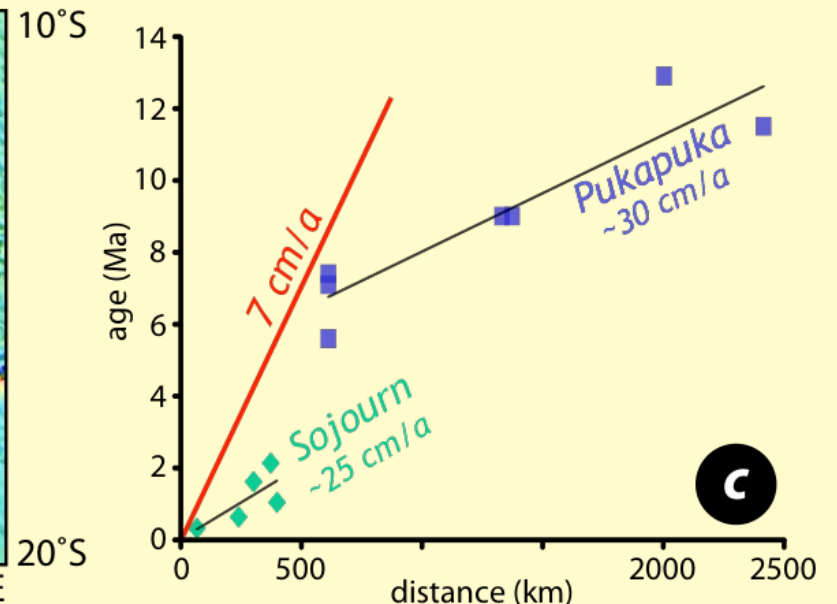
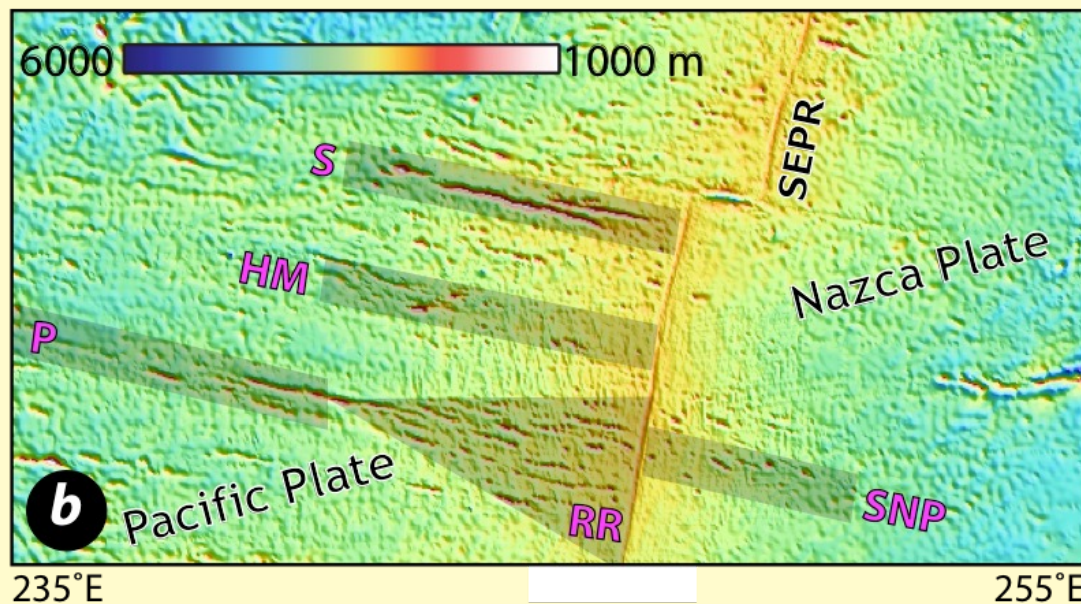
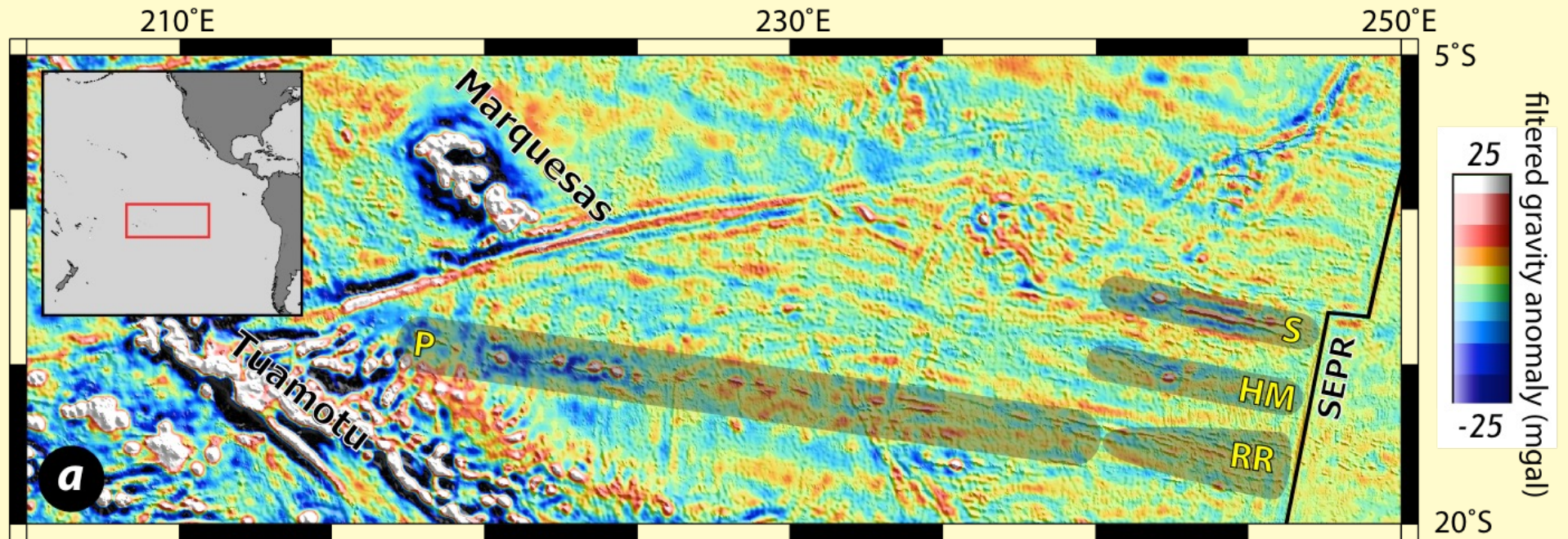
*Viscous Fingering:
Injection of a less-viscous
fluid into a more viscous one*

**Is viscous fingering important
in the South Pacific?**



Weerarante et al. [JGR, 2007]

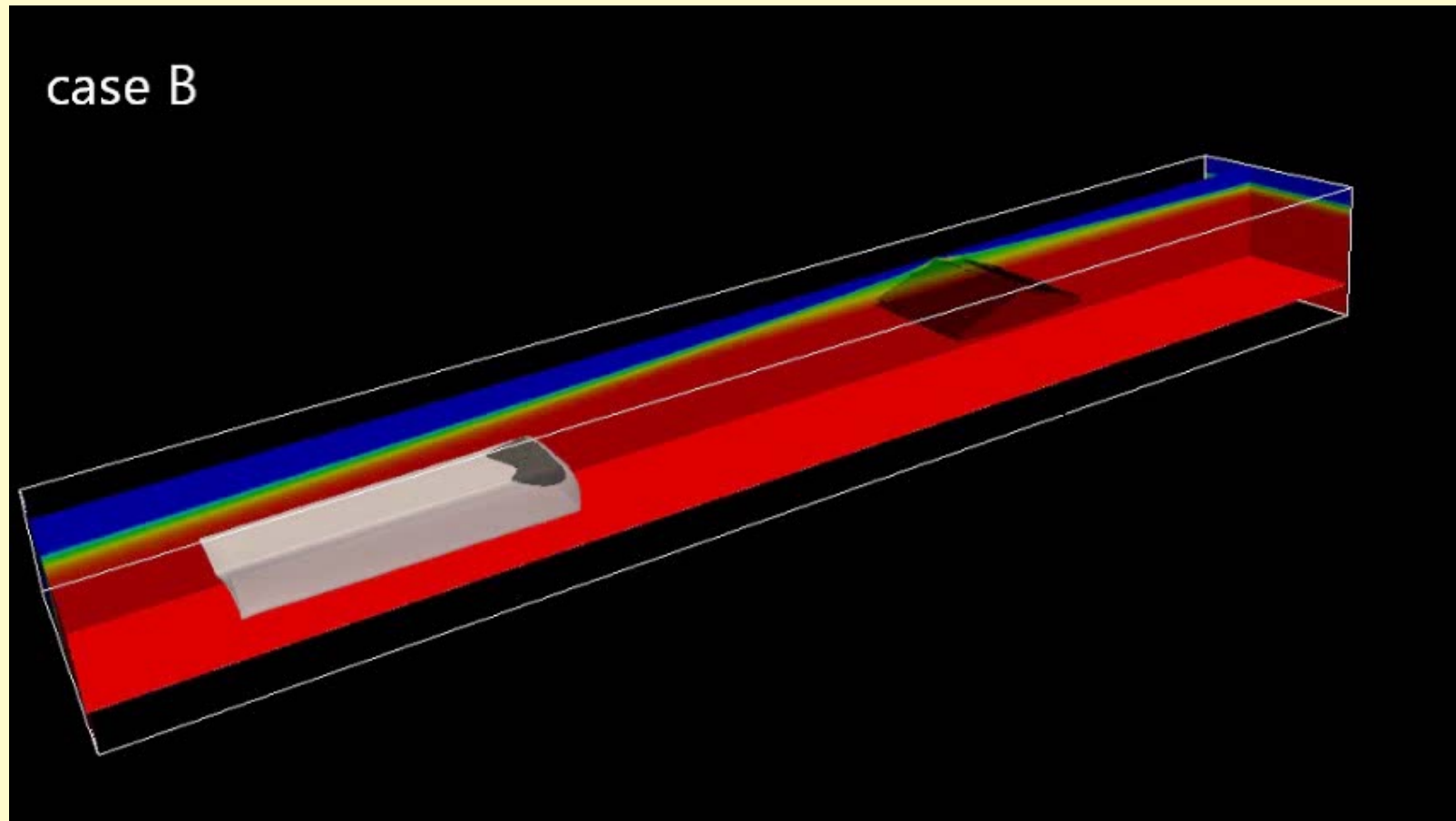
Volcano chains in the South Pacific



S = Sojourn RR = Rano Rahi SNP = Seamounts
 HM = Hoto-Matua P = Pukapuka on Nazca Plate

Ballmer et al. [Geology, 2013]

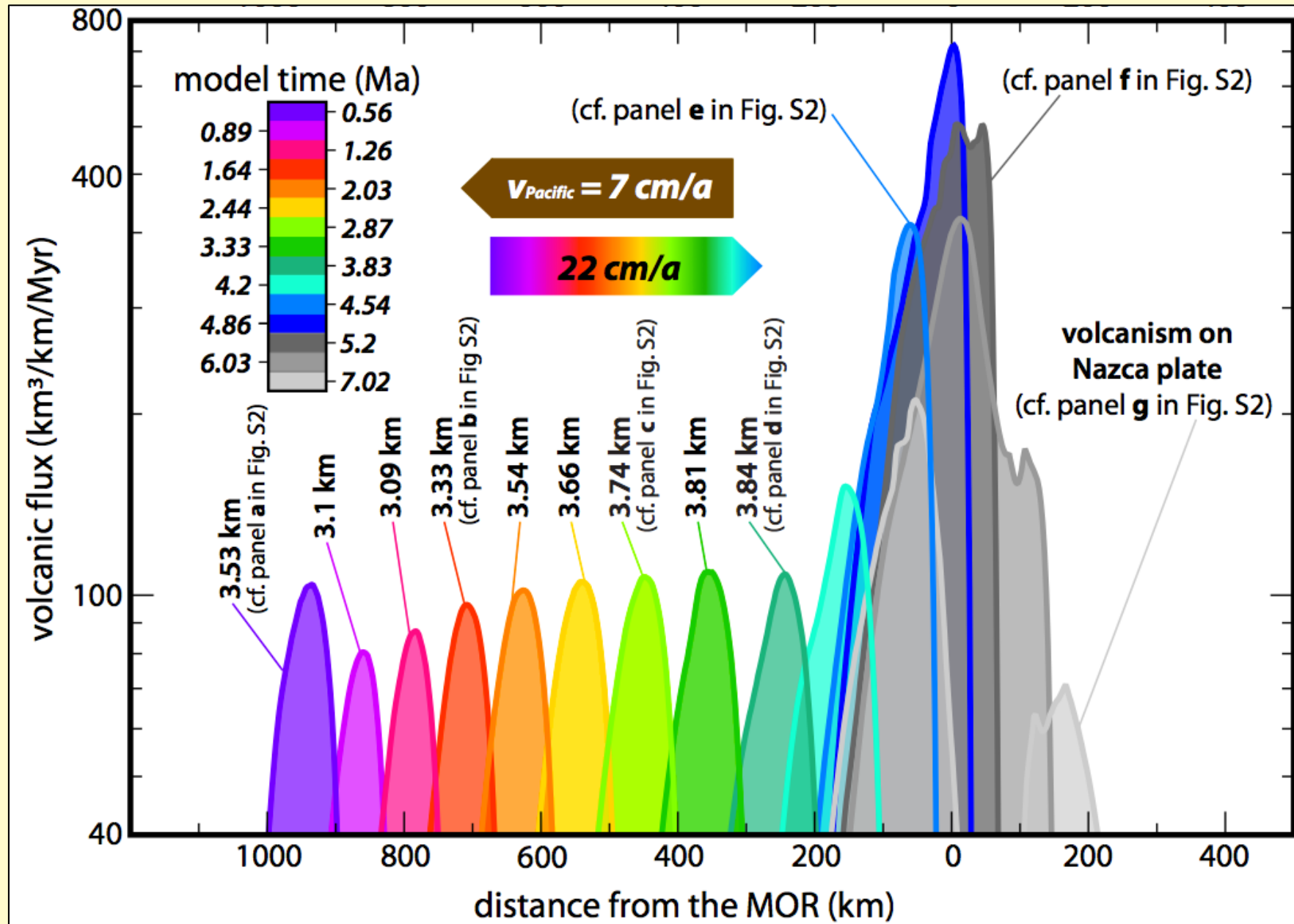
Hot and Wet Anomaly being pushed toward the East Pacific Rise



Like a “sideways” plume

Ballmer et al. [2013]

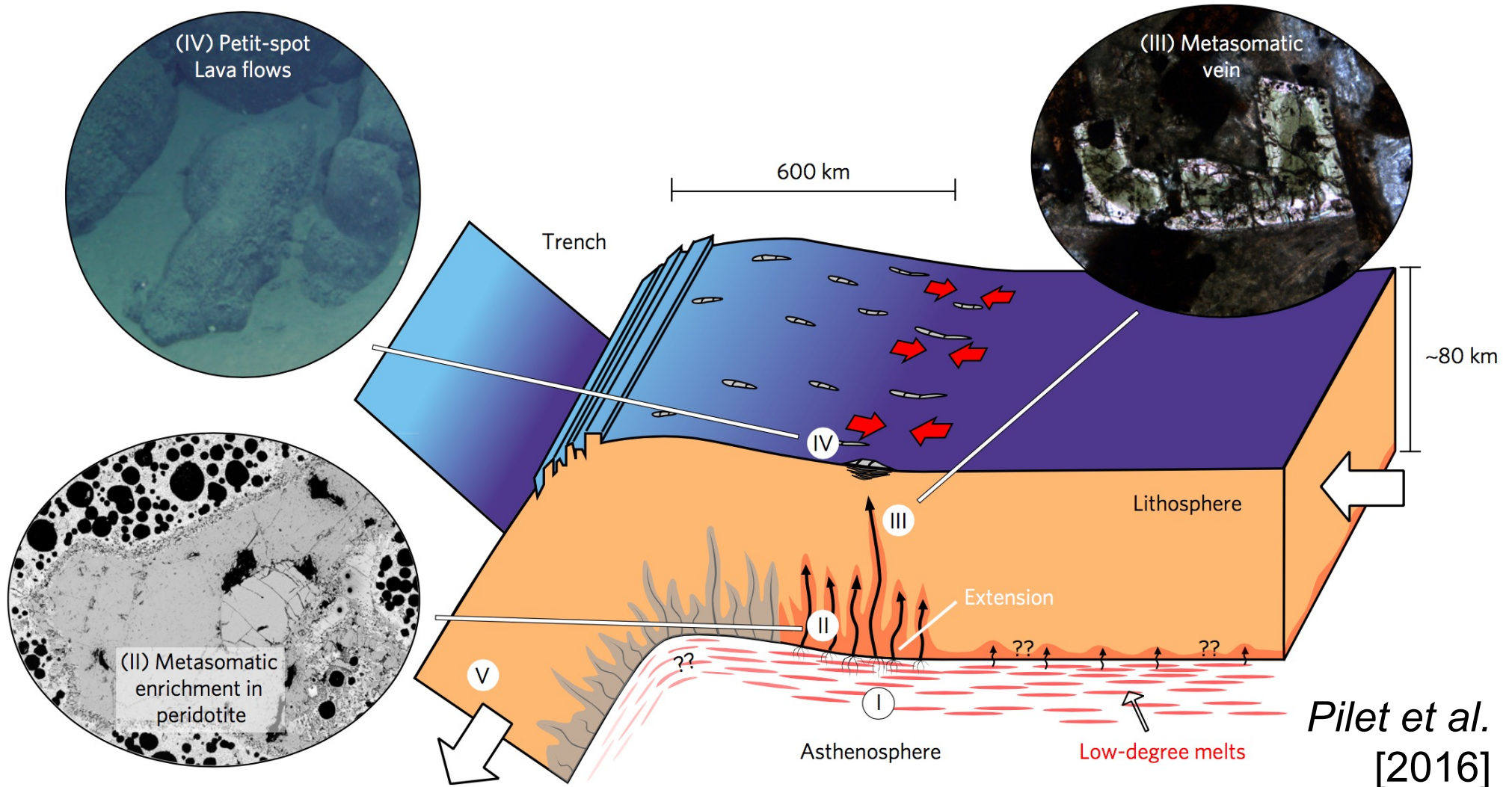
Case B: Hot and Wet Anomaly (Pukapuka)



4. Petit Spots

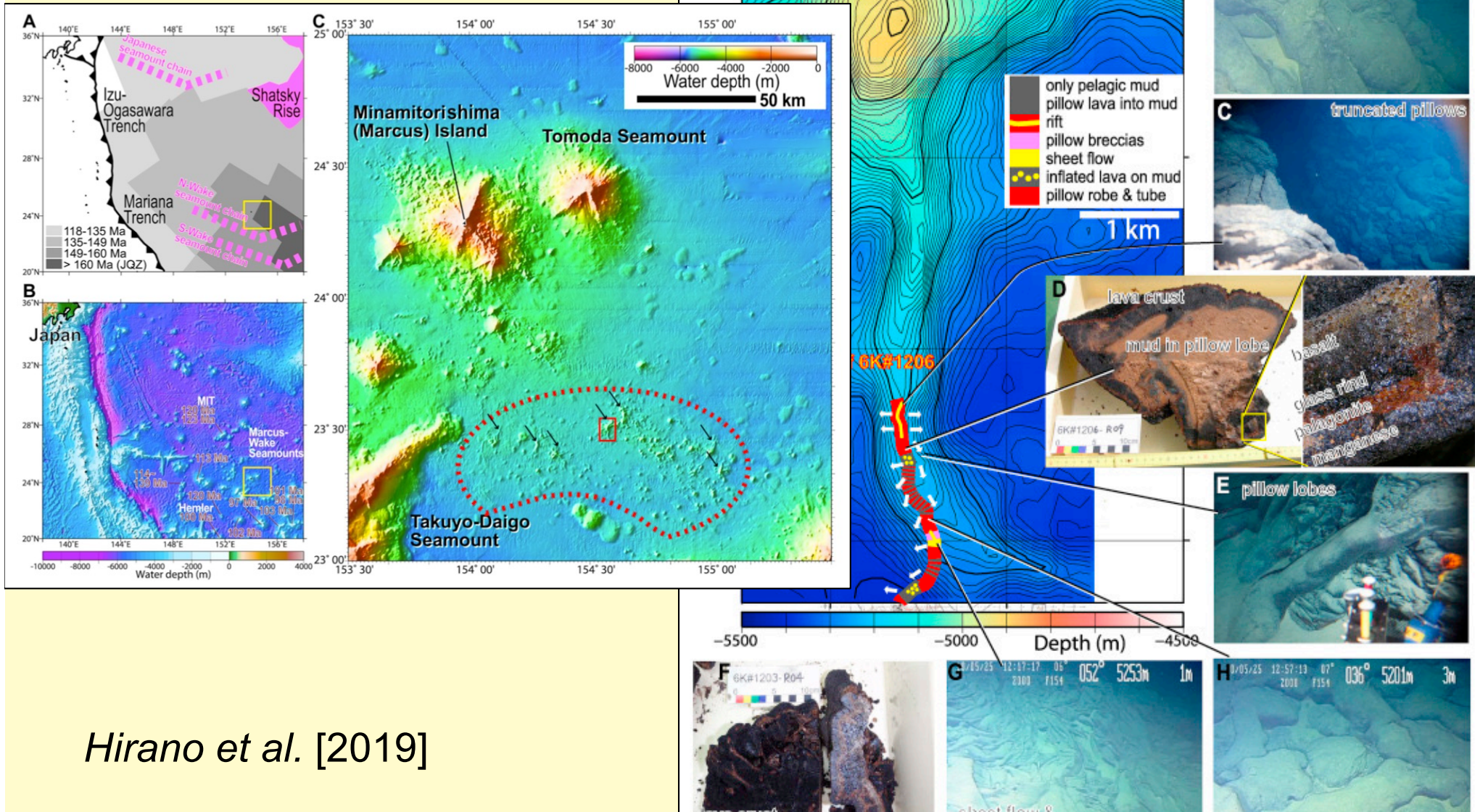
- Very minor volcanism
- Happens near trenches
- Due to plate flexure

Implies that low-degree melt is prevalent in the upper asthenosphere



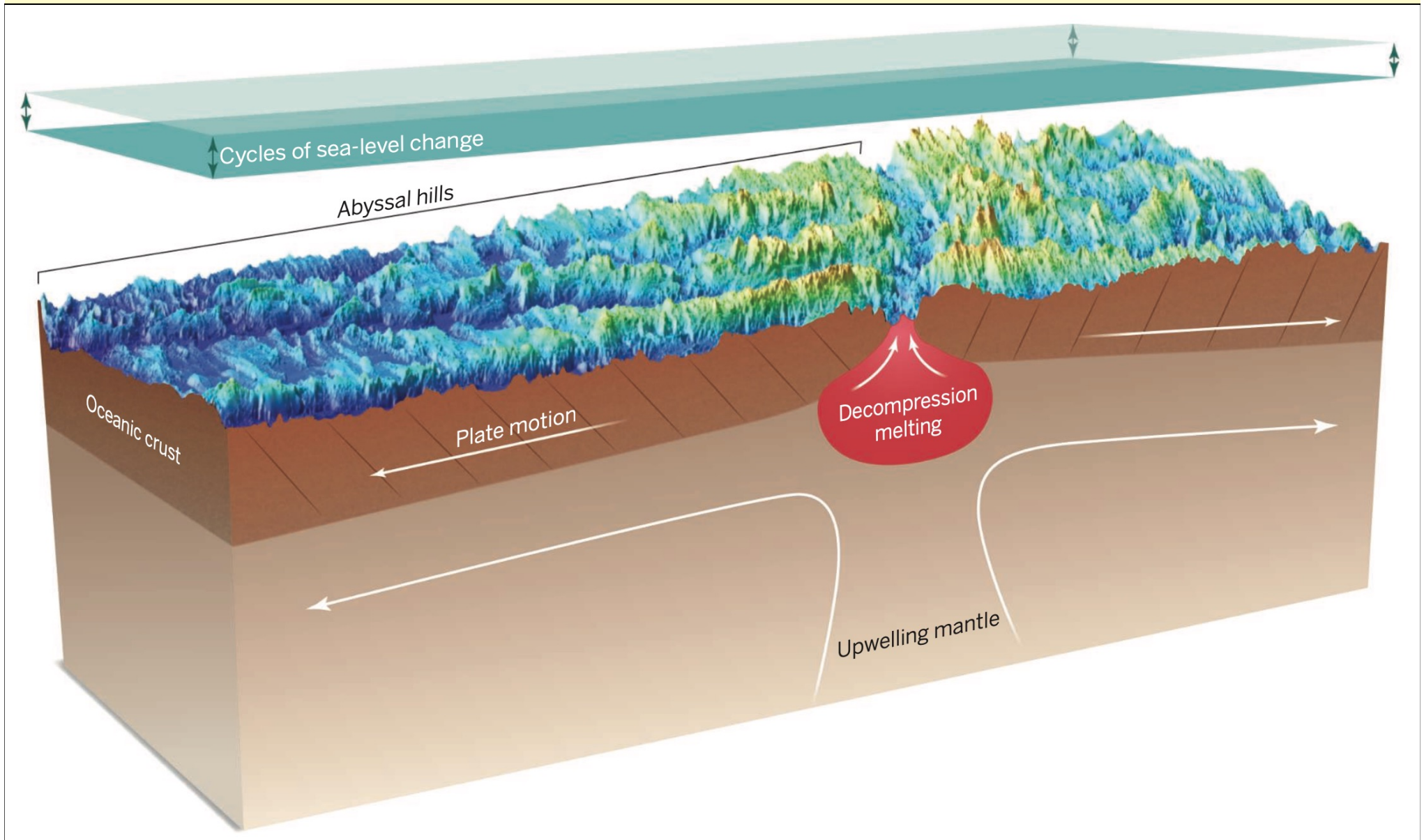
4. Petit Spots

- Very minor volcanism
- Happens near trenches
- Due to plate flexure



Hirano et al. [2019]

5. Surface Loading Affects on Volcanism



Are the abyssal hills caused by sea level changes?

Rising Sea Level:

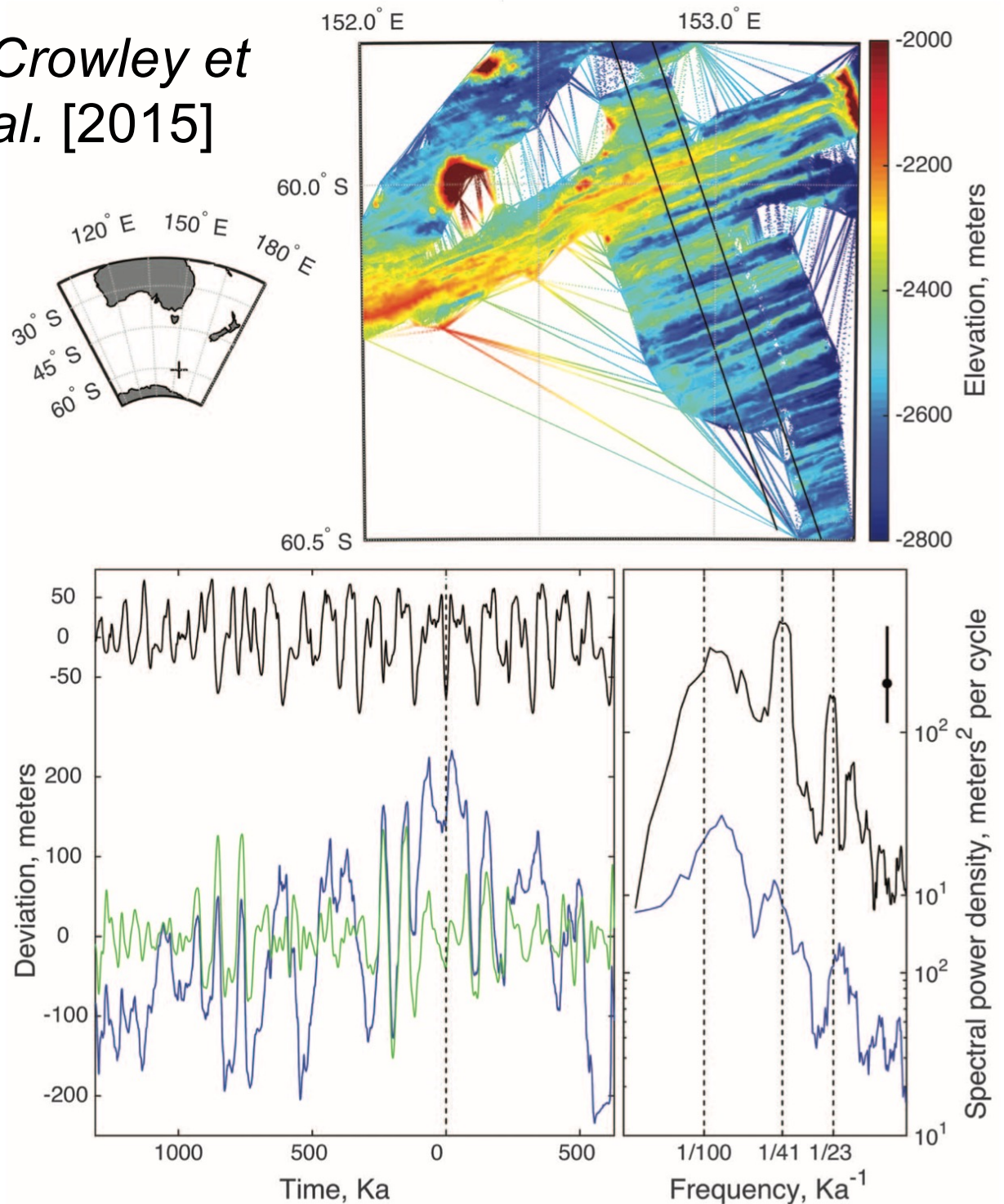
- Increasing pressure
- Melting is suppressed
- Less volcanic output
- A “valley” in the abyssal hills

Falling Sea Level:

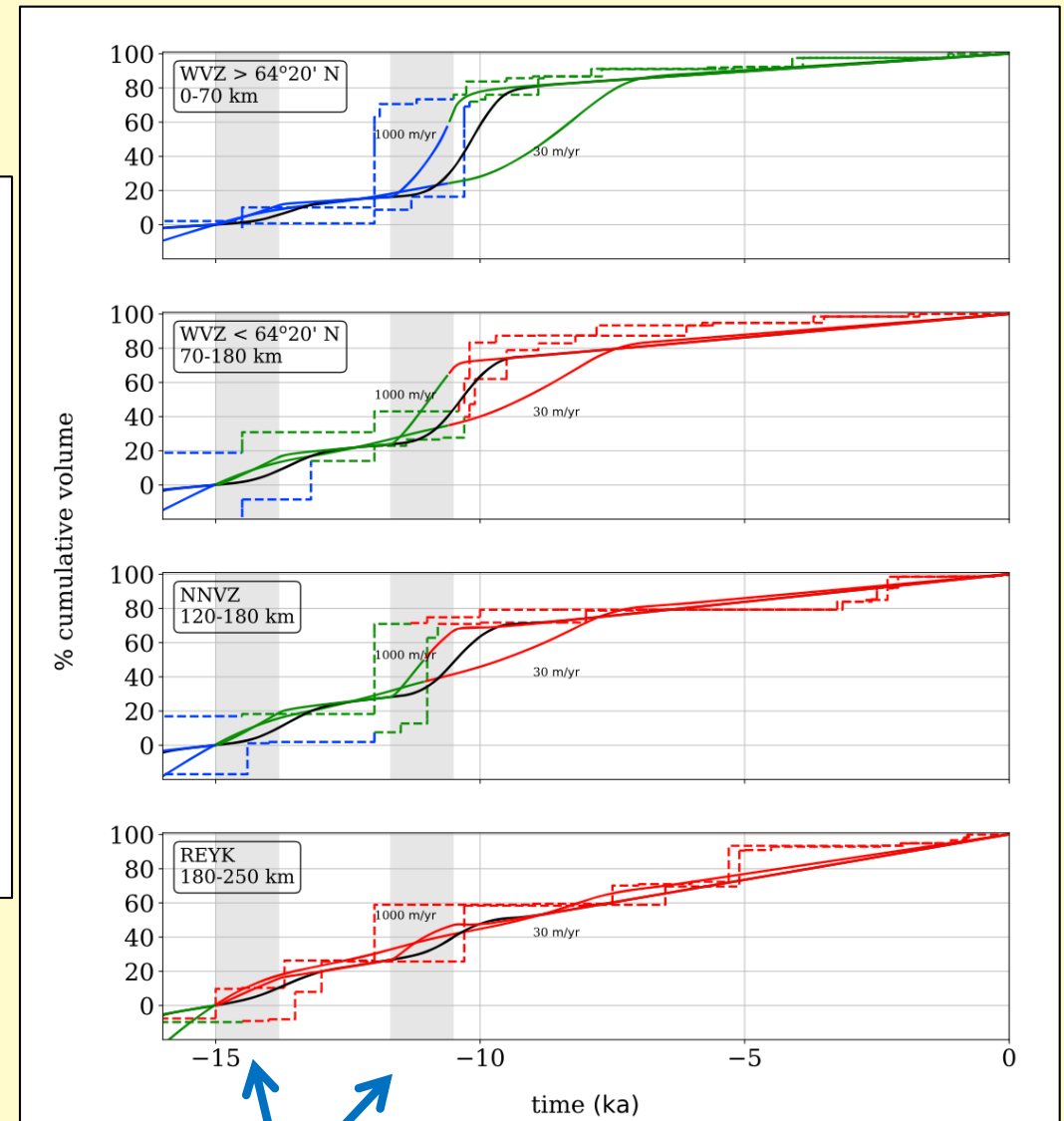
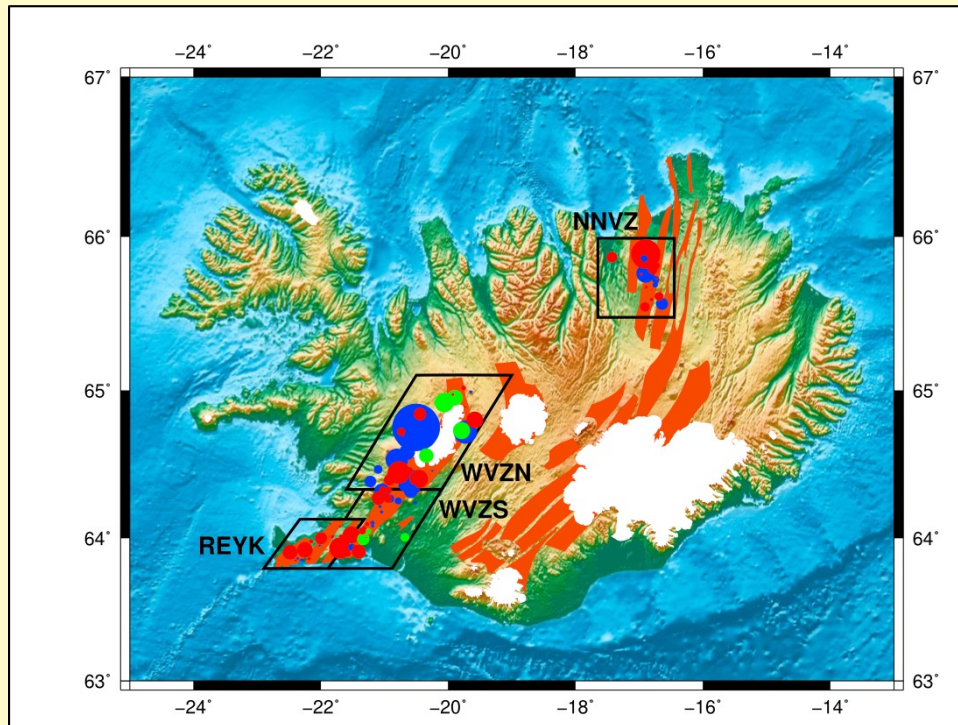
- Decreasing pressure
- Melting is enhanced
- More volcanic output
- A “peak” in the abyssal hills

Observation of Milankovitch timescales in the abyssal hill fabric →

Crowley et al. [2015]



Amplification of Icelandic volcanism during deglaciation



Eksinchol et al. [2019]

Deglaciation periods

Intraplate Volcanism - in the oceans and on land

- More than 24,000 seamounts on the seafloor
- Many still undiscovered and we don't know their ages
- Many fields of minor volcanism on continental areas

Melting can be caused by:

- **Plumes** rising from the LLSVP edges
- **Small-scale convection** on older lithosphere
- **Shear-driven upwelling** (especially in the Pacific)
- **Petit-spot volcanism** near trenches
- Removal of surface loads

Implications:

→ Mantle beneath the plates is close to melting or is already partially melted

→ Climate change can cause volcanism

