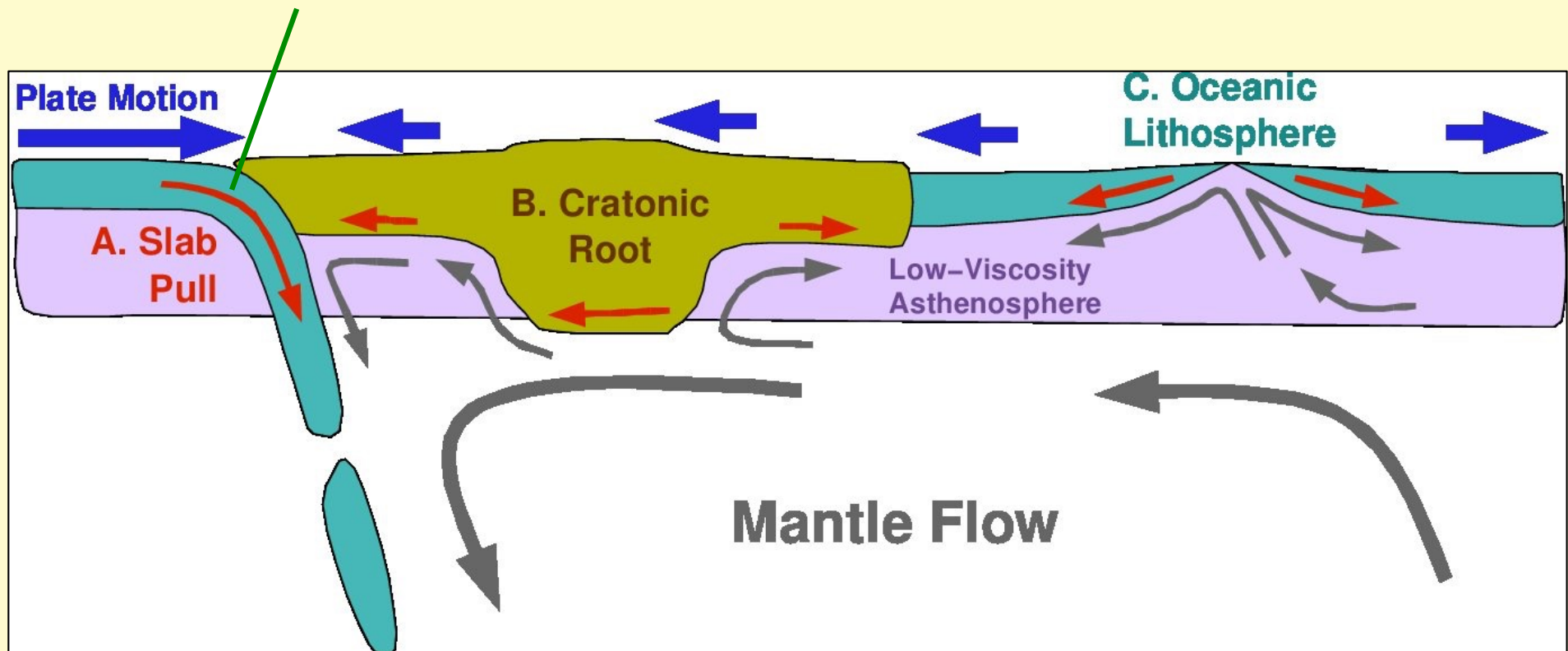


Lithosphere and Asthenosphere: Composition and Evolution

GEO-DEEP9300

Elastic Lithosphere: Valerie Maupin
Plate Flexure Clint Conrad



Geodynamic Processes of the Lithosphere & Asthenosphere

All Geodynamic Processes (except earthquakes) involve a force balance related to: $\text{Force} = \text{Mass} * \text{Acceleration}$

(Density * acceleration) =

Acceleration is negligible

(body force) +

(gradient of stresses) +

(material deformation)

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(material deformation)

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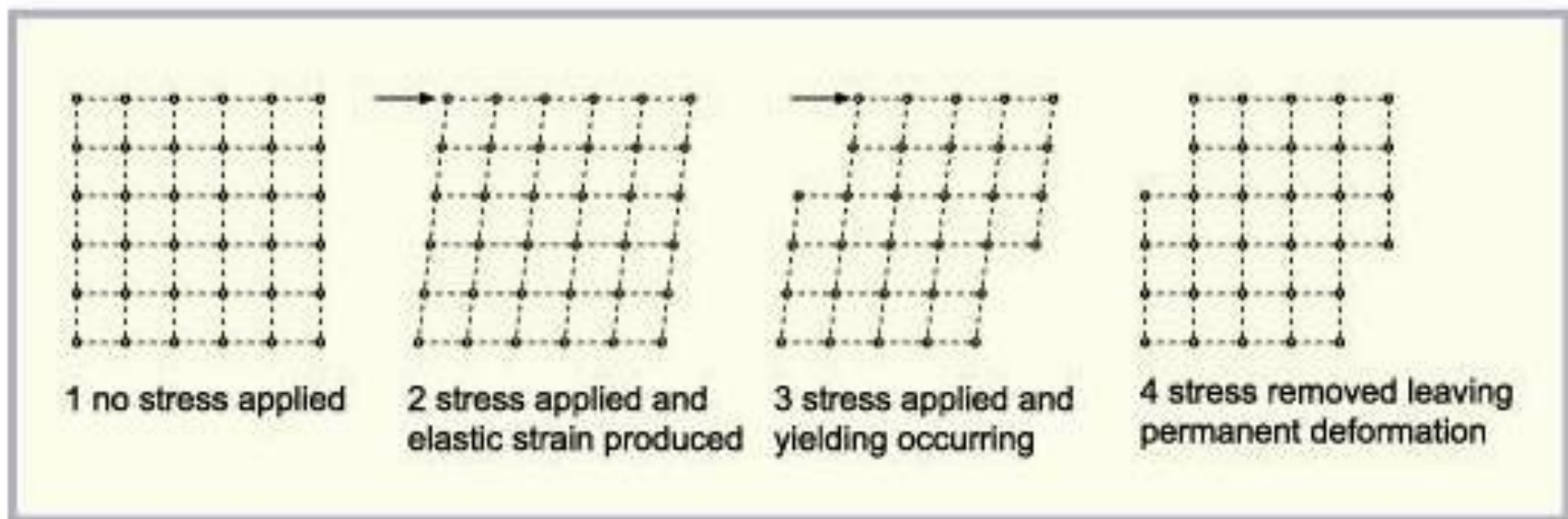
deformation depends on rheology

(material deformation)

→ Body forces drive geodynamic processes

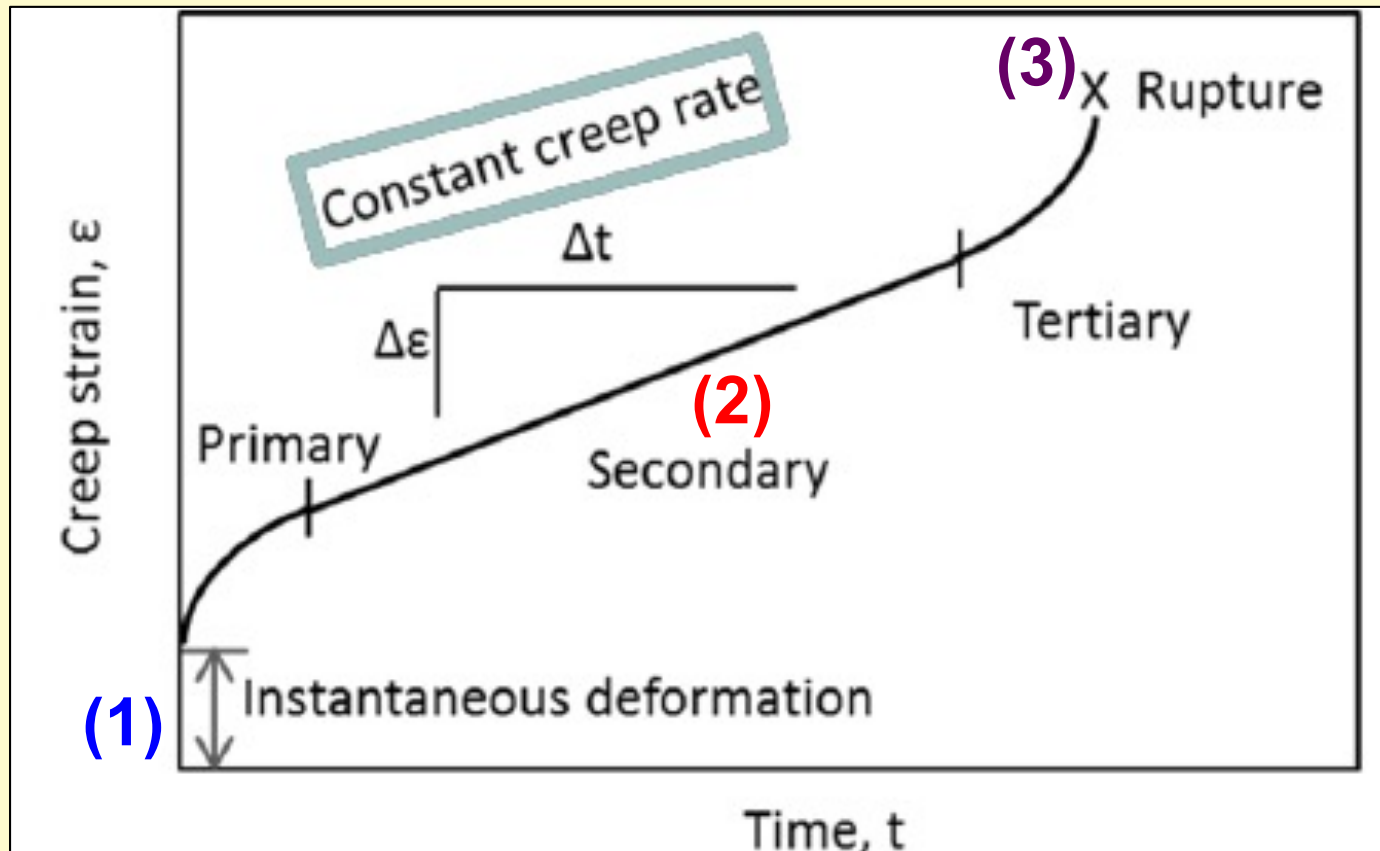
→ Material deformation resists the body forces

Apply a constant stress to a material: How does it deform?



Stages of Plastic Deformation

Apply a constant stress to a material: How does it deform?



Types of rheology that are important for the lithosphere:

1. Elastic Deformation: Stress \sim Strain
2. Viscous Deformation: Stress \sim Rate of Strain
3. Brittle Fracture: Strain \rightarrow infinity (discontinuity)

For a viscoelastic material:

Elastic Deformation:

$$(\text{stress}) = E (\text{strain})$$

E = Young's Modulus

$E = 70 \text{ GPa}$ (typical rock)

Viscous Deformation:

$$(\text{stress}) = \eta (\text{strain-rate})$$

η = Newtonian Viscosity

$\eta = 10^{20} \text{ Pa s}$ (typical mantle)

***Maxwell Time* $\sim 2\eta / E \sim 100 \text{ years}$**

The stresses relax over this timescale

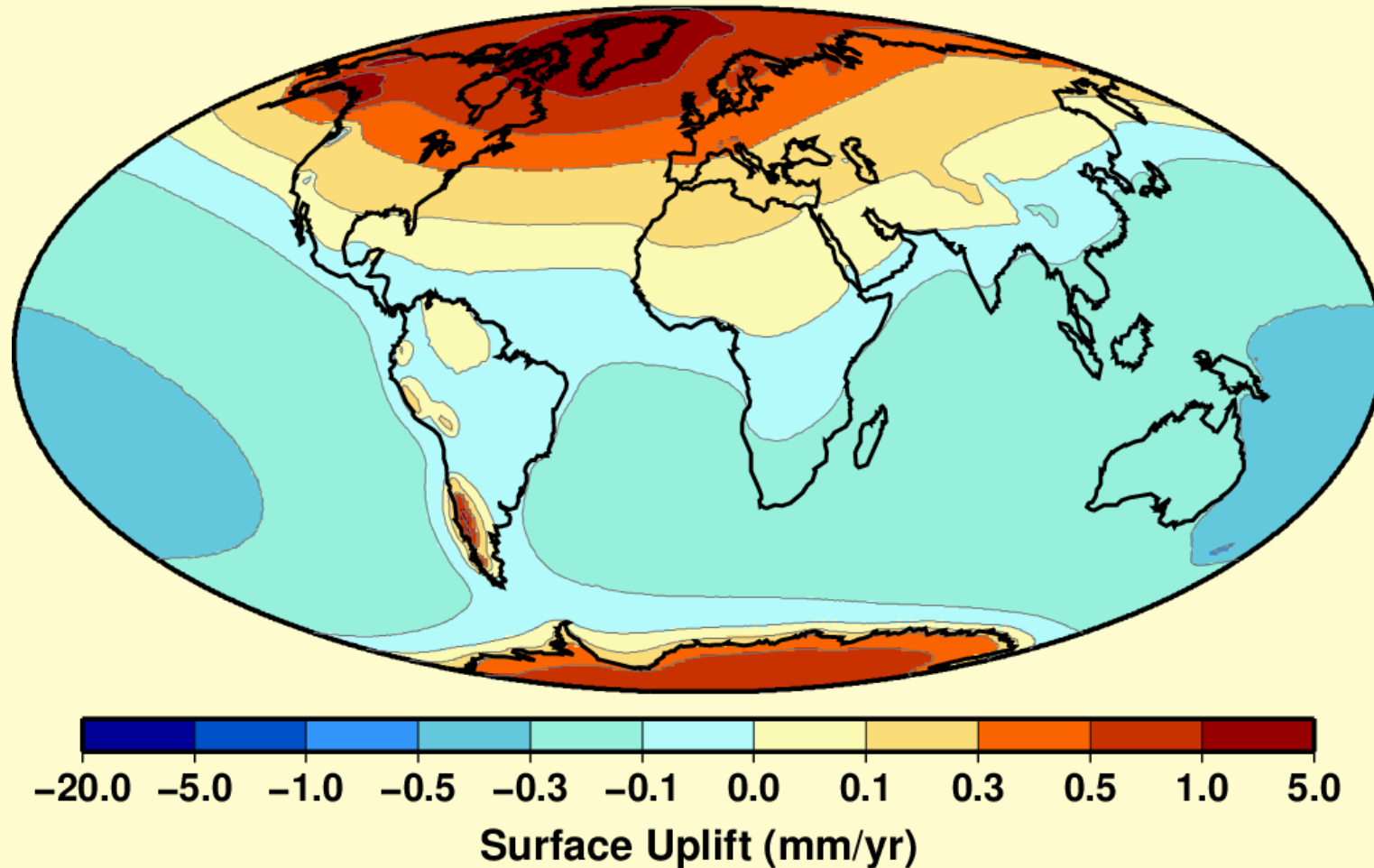
Shorter than 100 years:

Elastic deformation

Longer than 100 years:

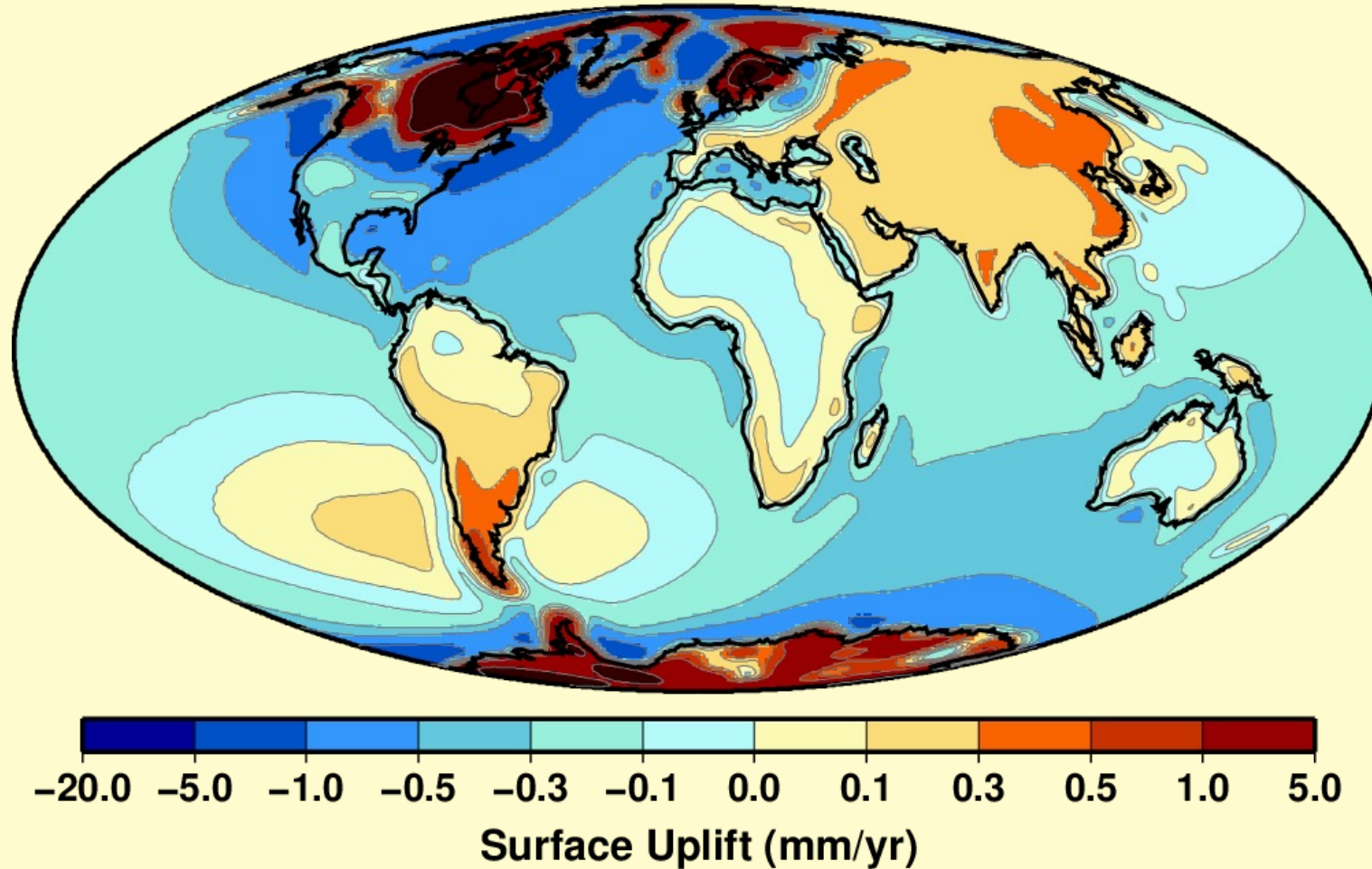
Viscous deformation

Elastic Response of the Earth to Surface Loads: Recent Ice Melt: Instantaneous Response



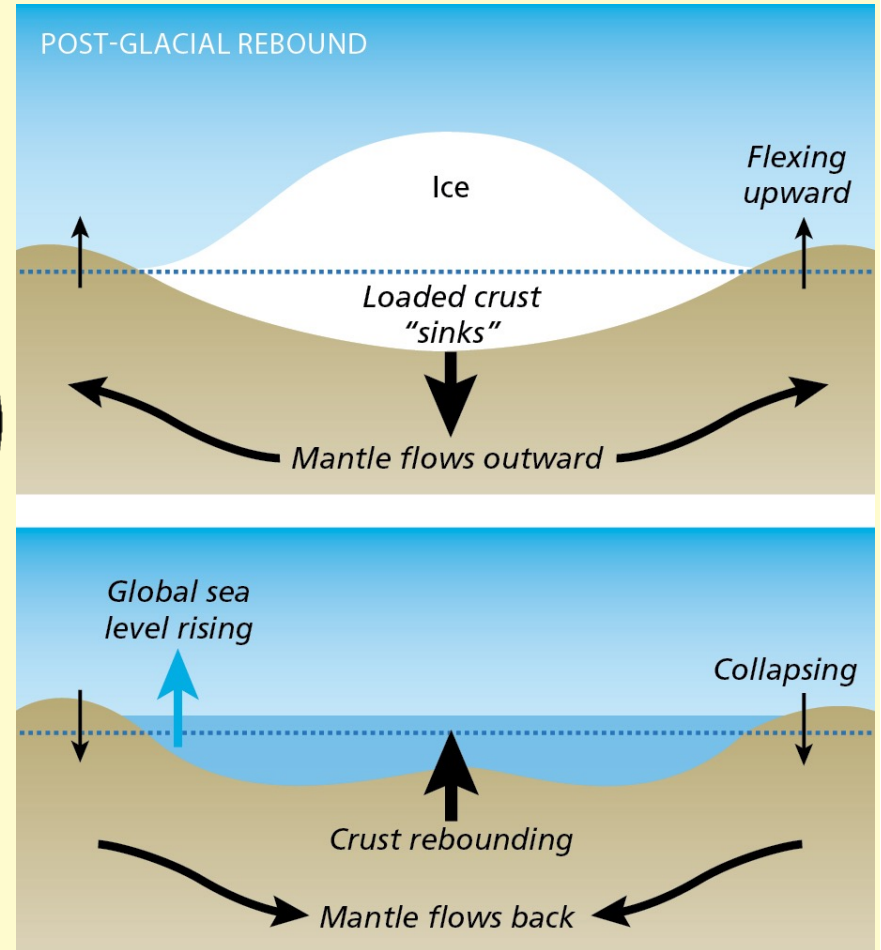
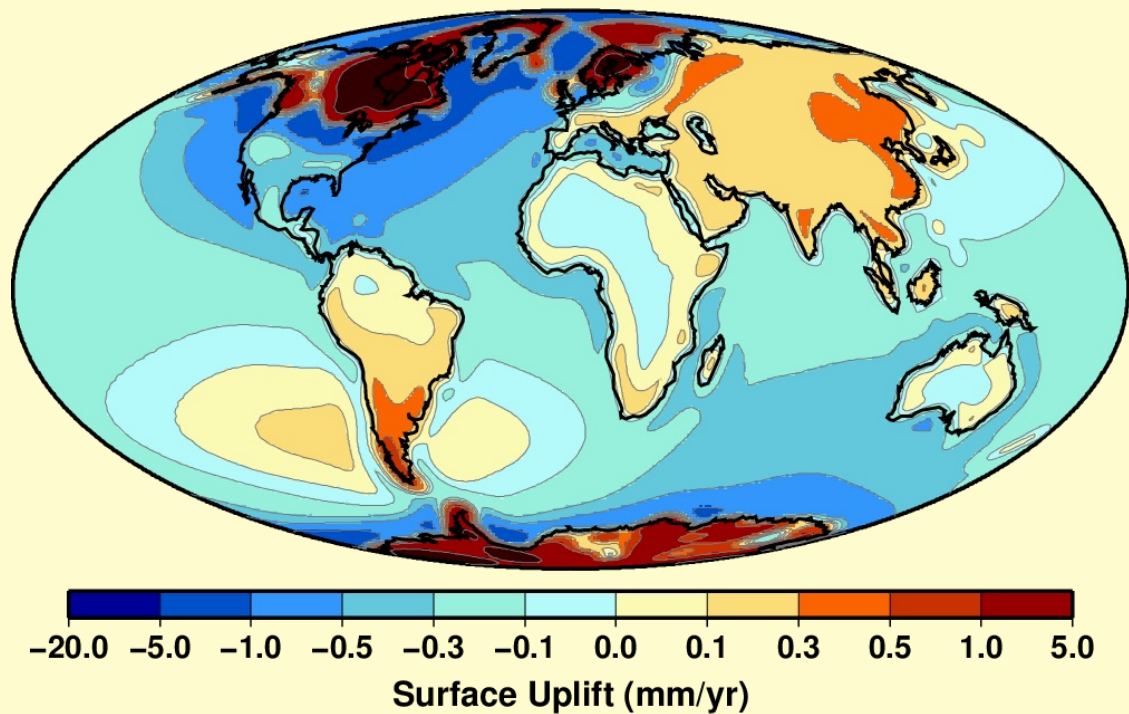
Conrad [2013]

Viscous Response of the Earth to Surface Loads: Postglacial Rebound after Last Ice Age ($\sim 10^4$ years ago)



Paulson et al. [2007]

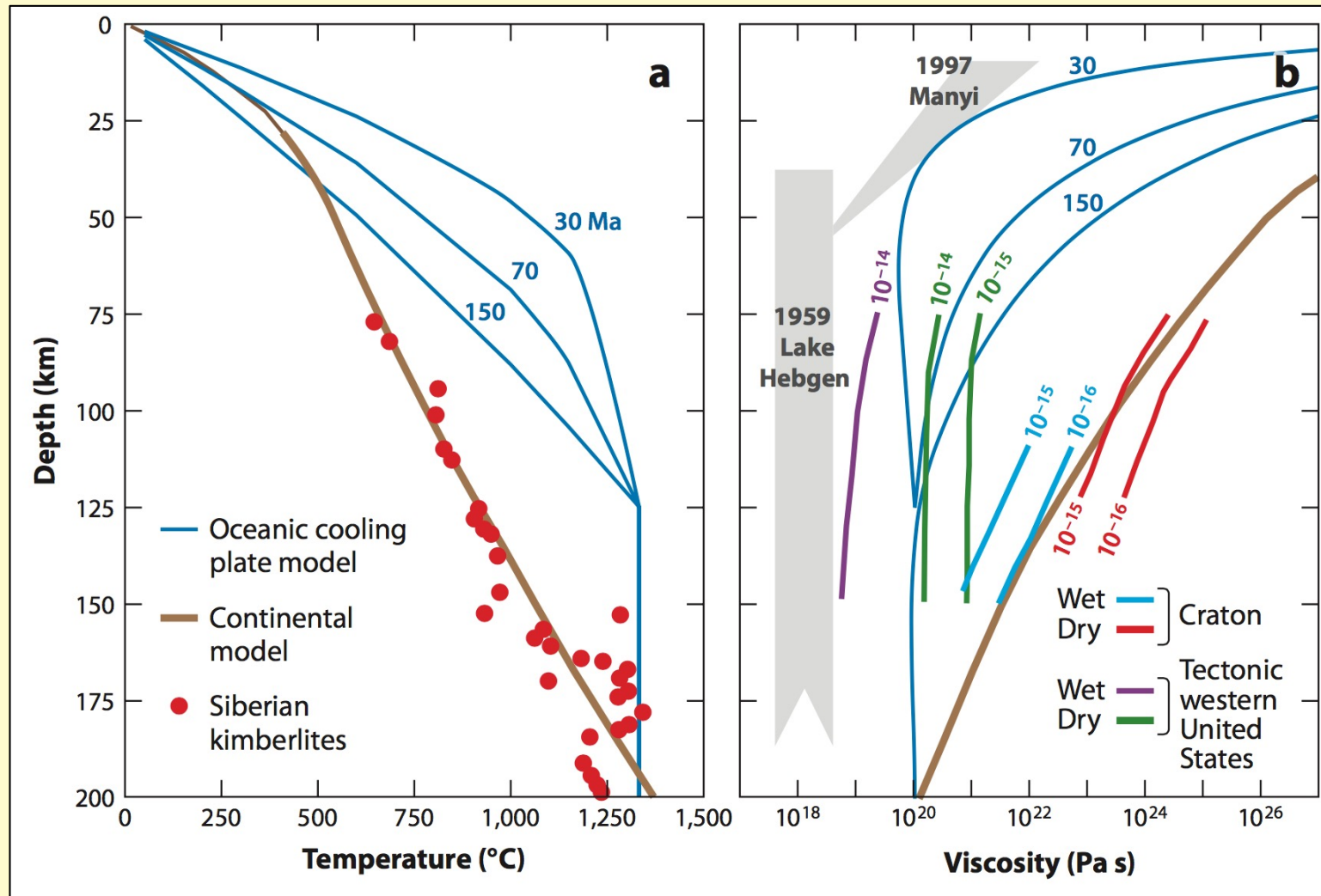
Viscous Response of the Earth to Surface Loads: Postglacial Rebound after Last Ice Age (~10⁴ years ago)



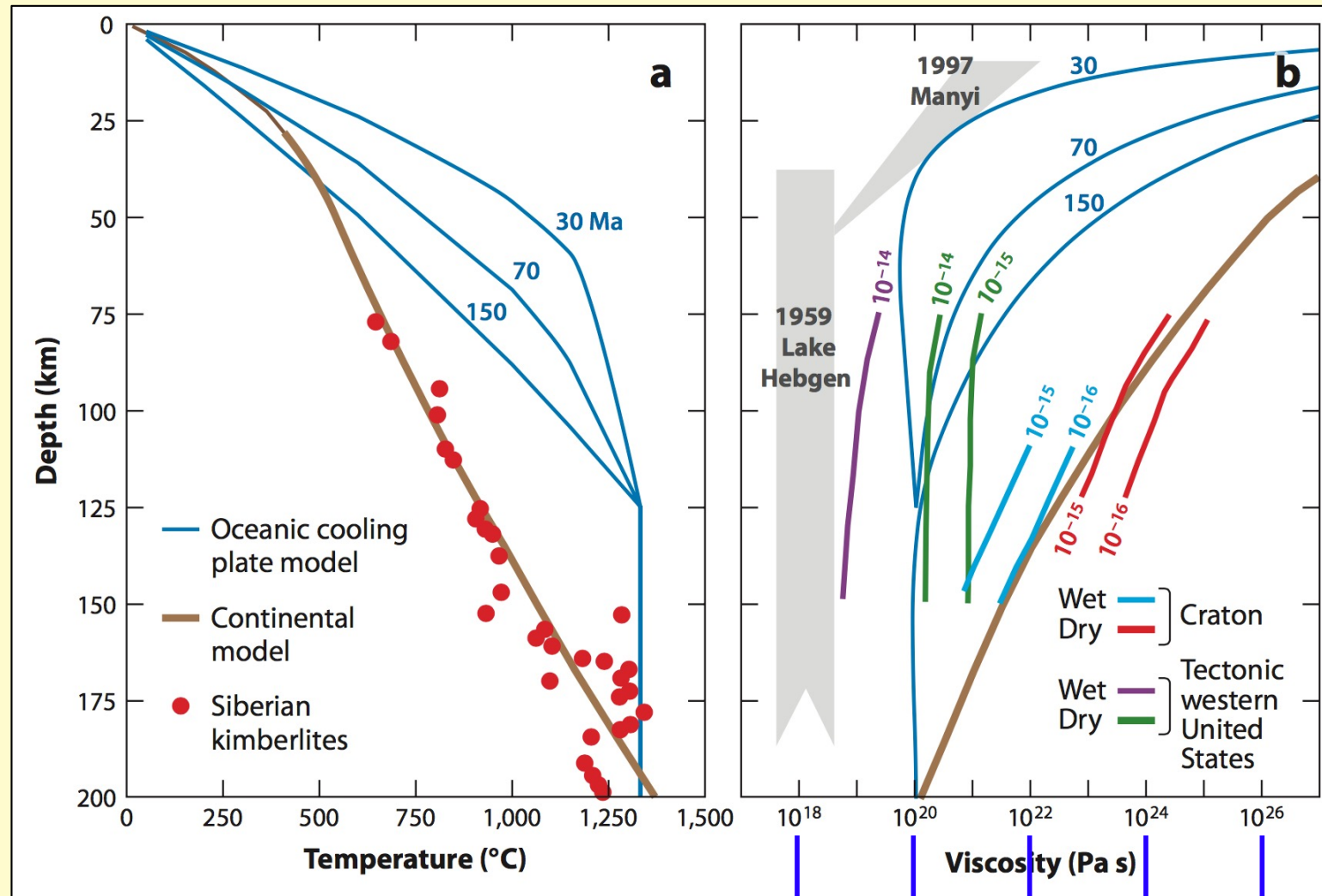
We can determine Earth's viscosity profile using postglacial rebound.

Paulson et al. [2007]

Viscosity Profile of the Lithosphere and Asthenosphere

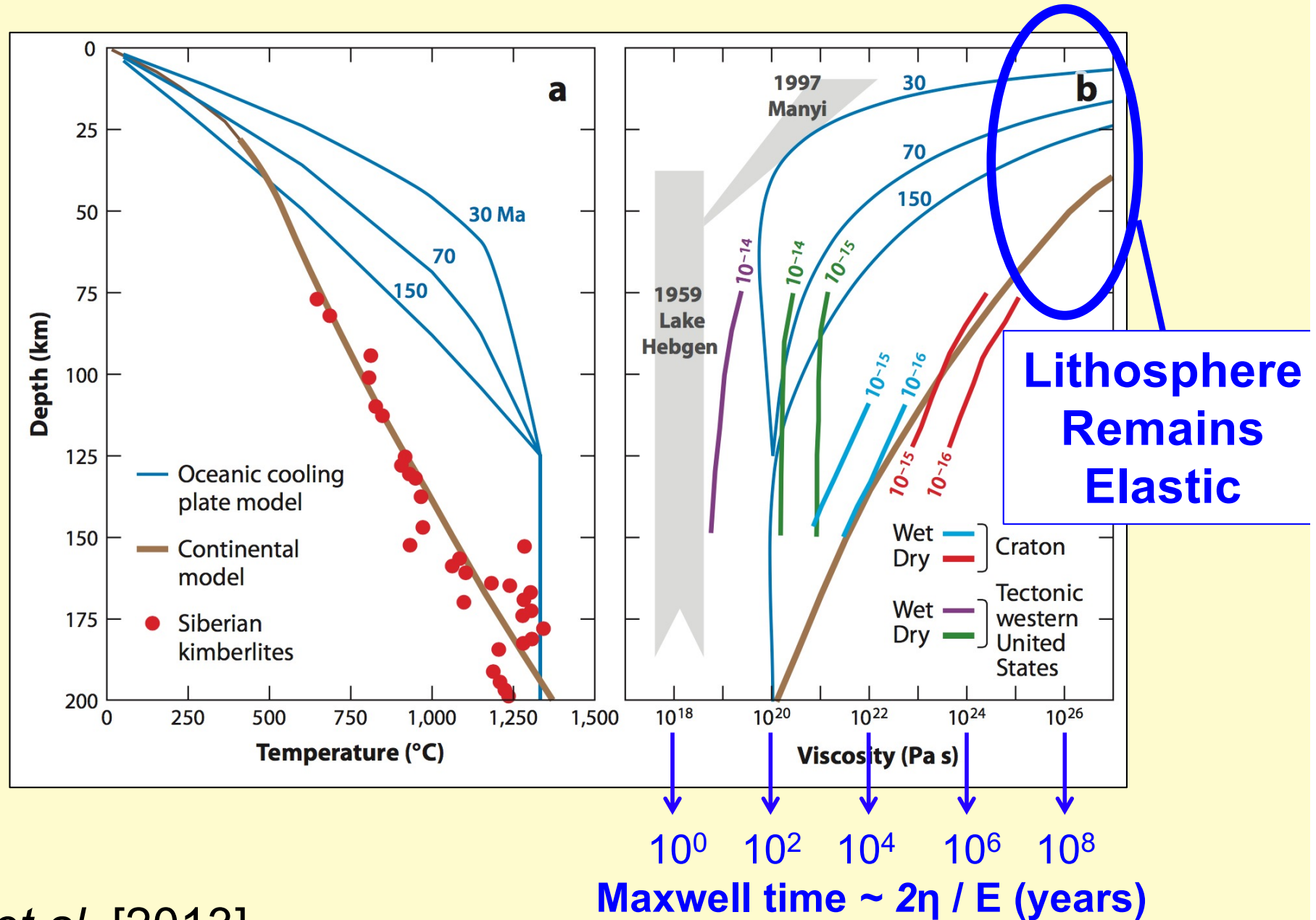


Viscosity Profile of the Lithosphere and Asthenosphere



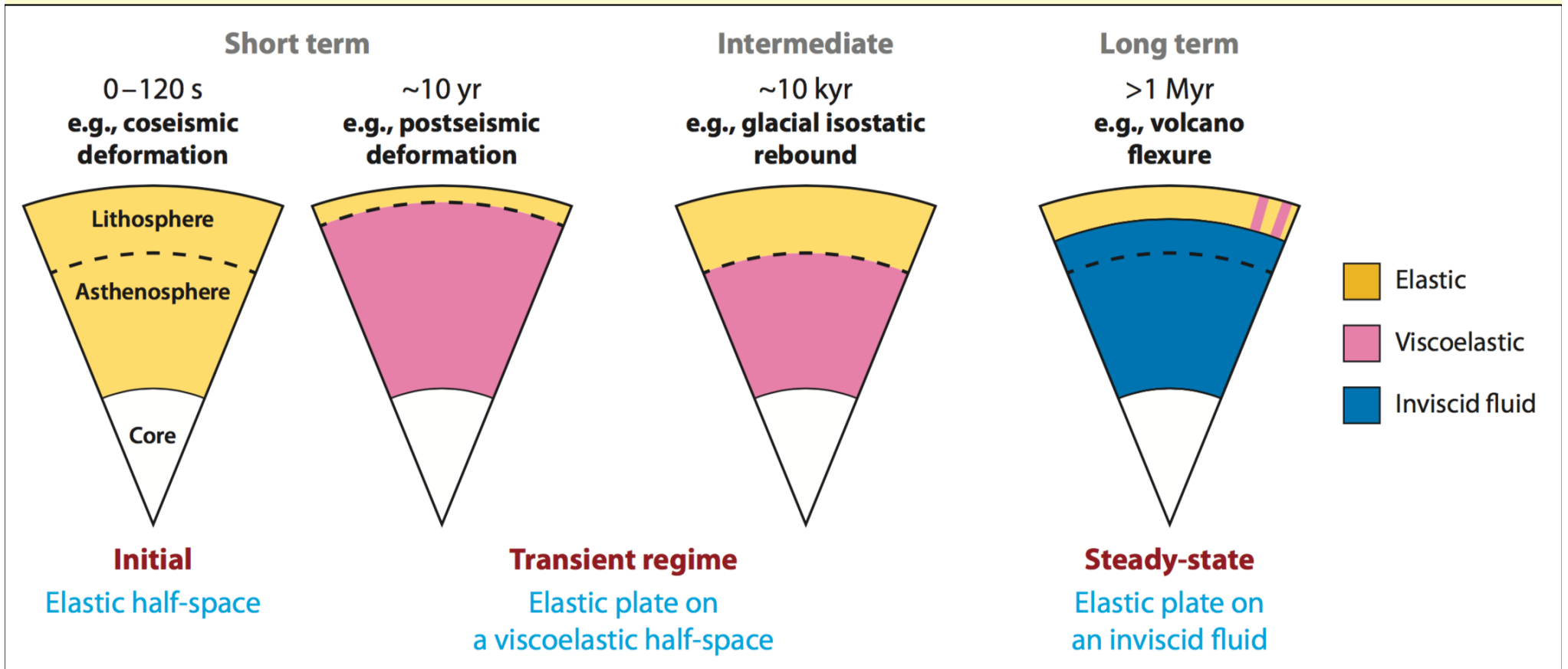
Watts et al. [2013]

Viscosity Profile of the Lithosphere and Asthenosphere

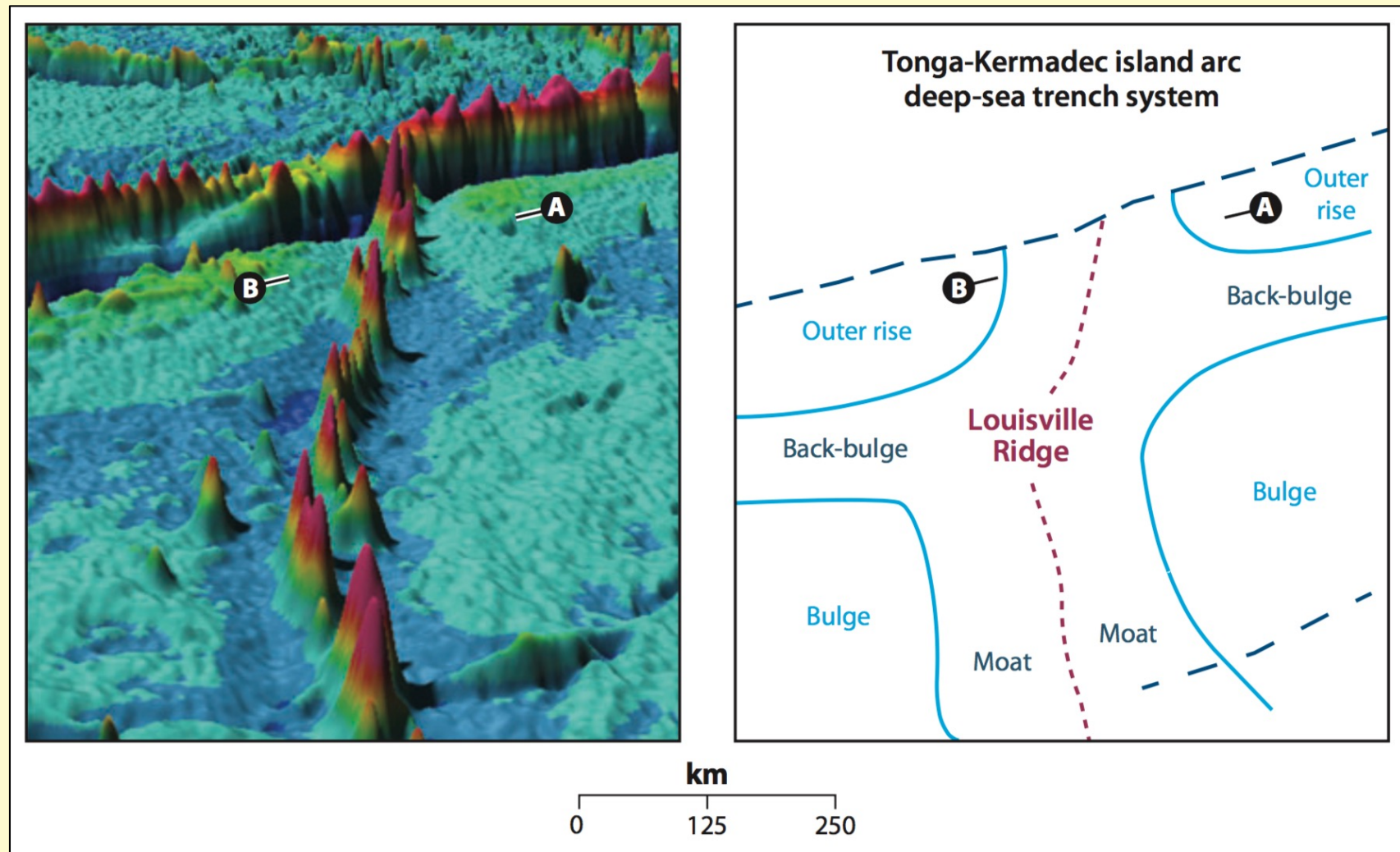


Watts et al. [2013]

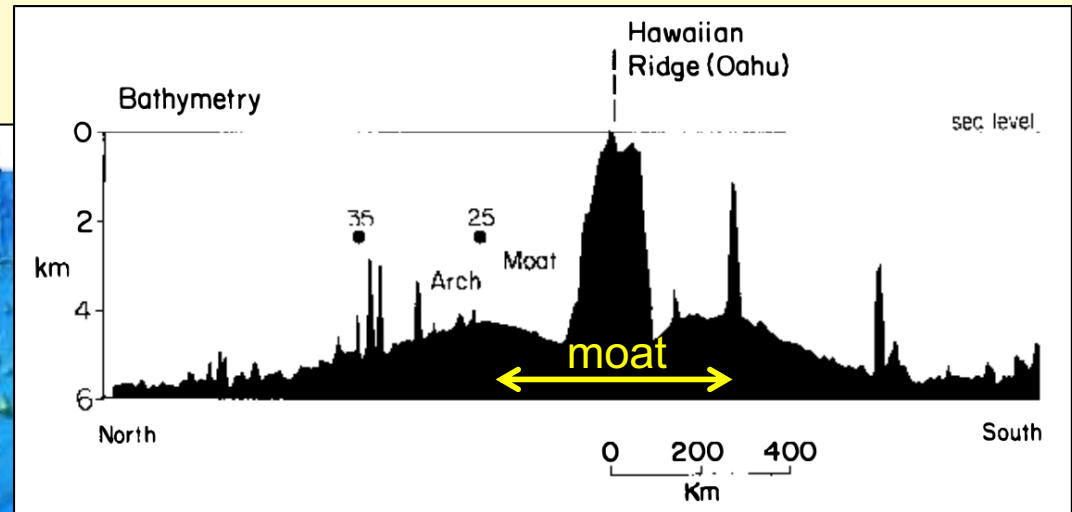
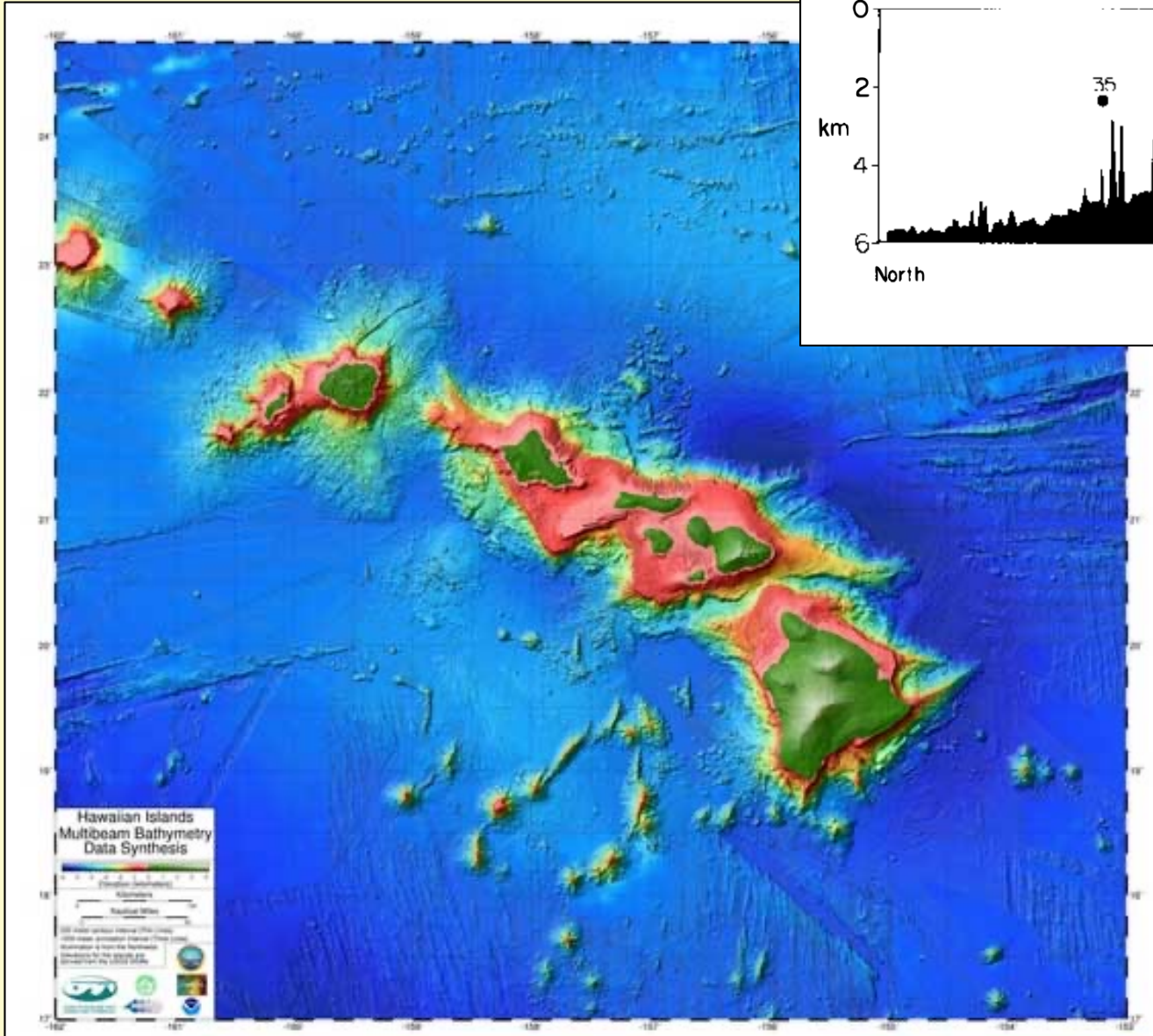
Timescale of loading determines the Earth's response: elastic vs. viscous



The lithosphere has a flexural response to different volcanic loads, and subduction bending



How does an elastic plate respond to an applied load?



Watts & Daly [1981]

Width of the “moat”
scales with the
elastic thickness

$$w_m \sim h_e^{3/4}$$

For Hawaii:

Moat width: $w_m \sim 500$ km

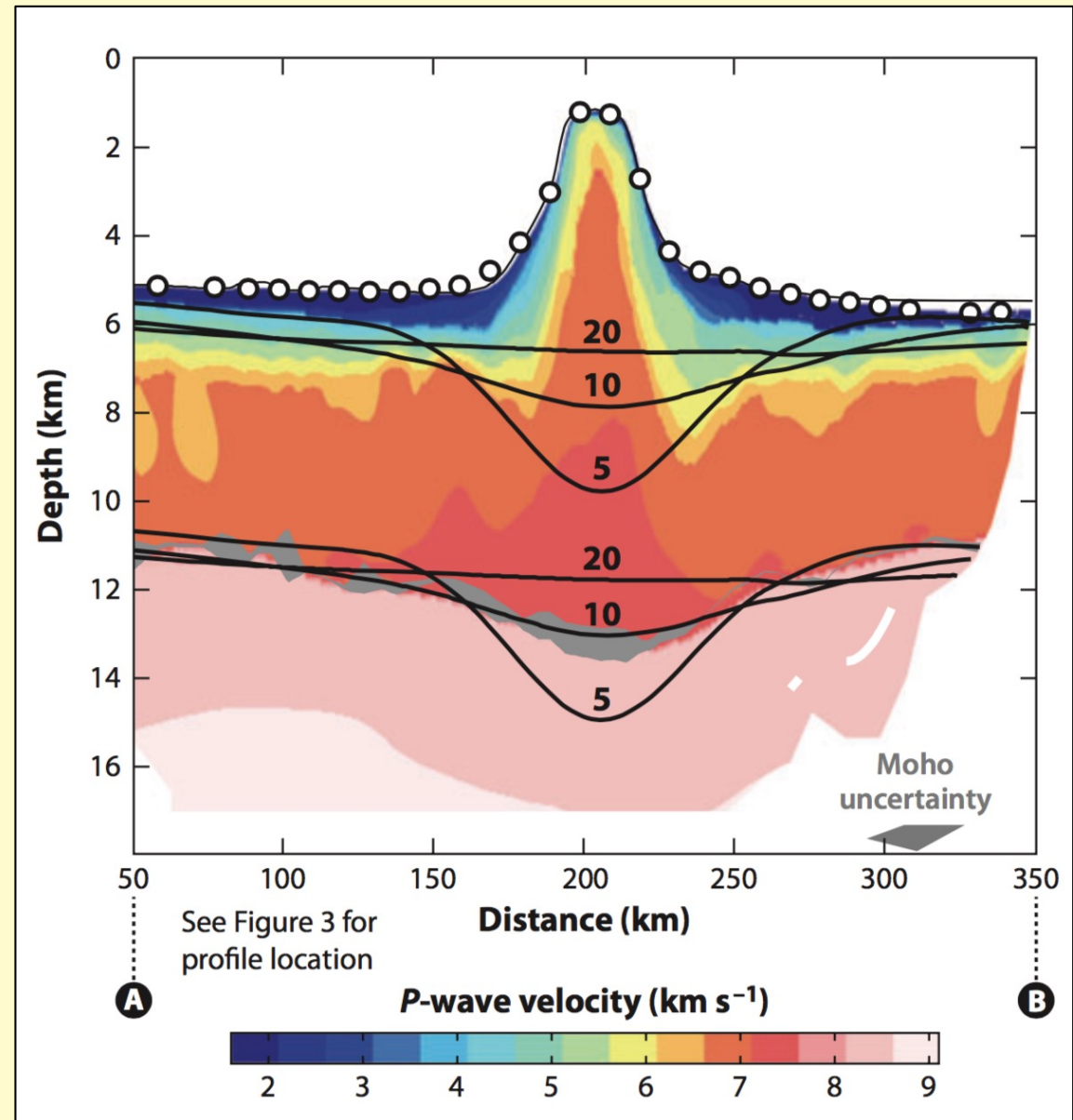
Elastic plate thickness:

$$h_e \sim 35$$
 km

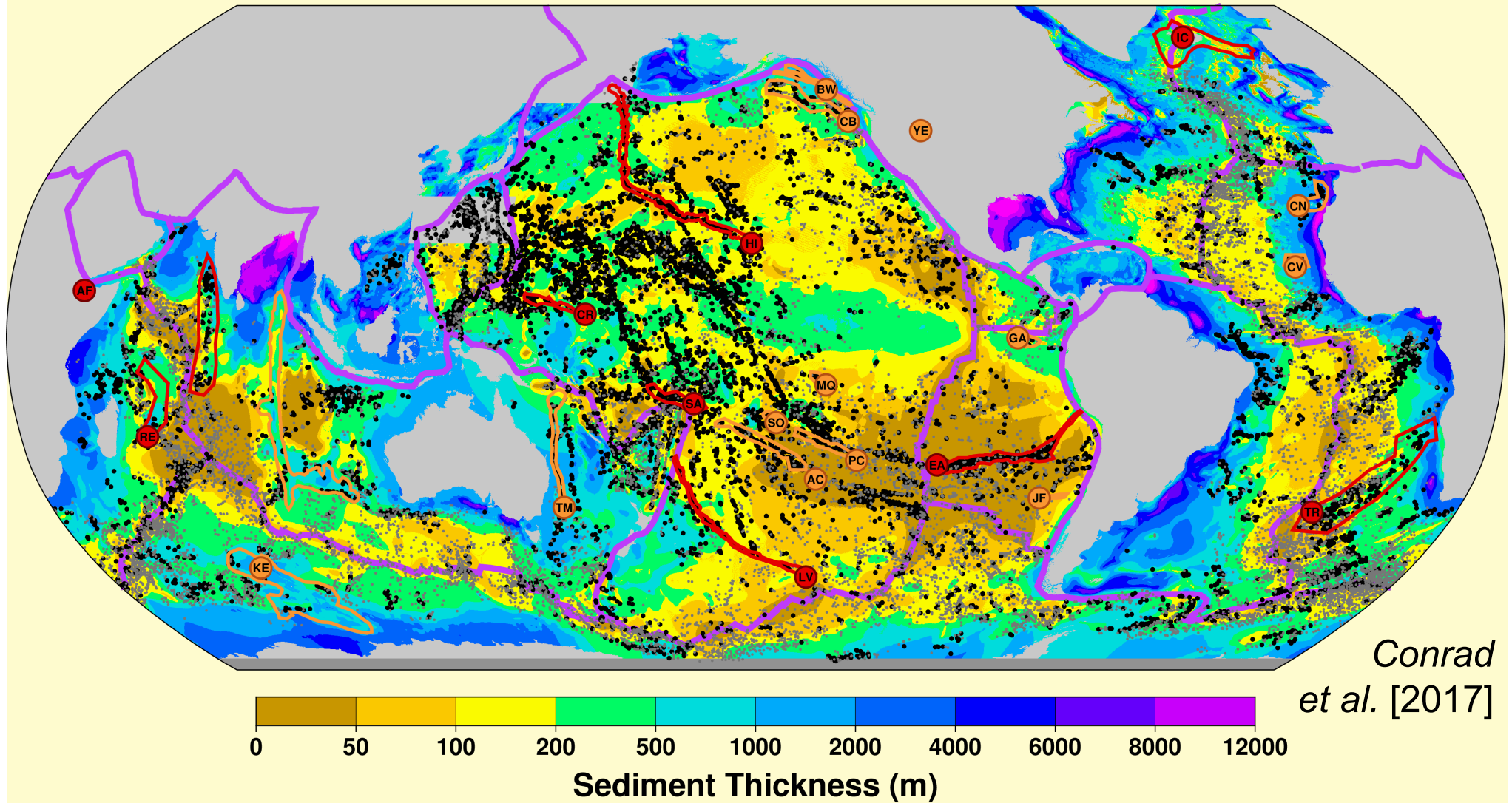
How does an elastic plate respond to an applied load?

Response of plate
with different
elastic thicknesses
(given in km)

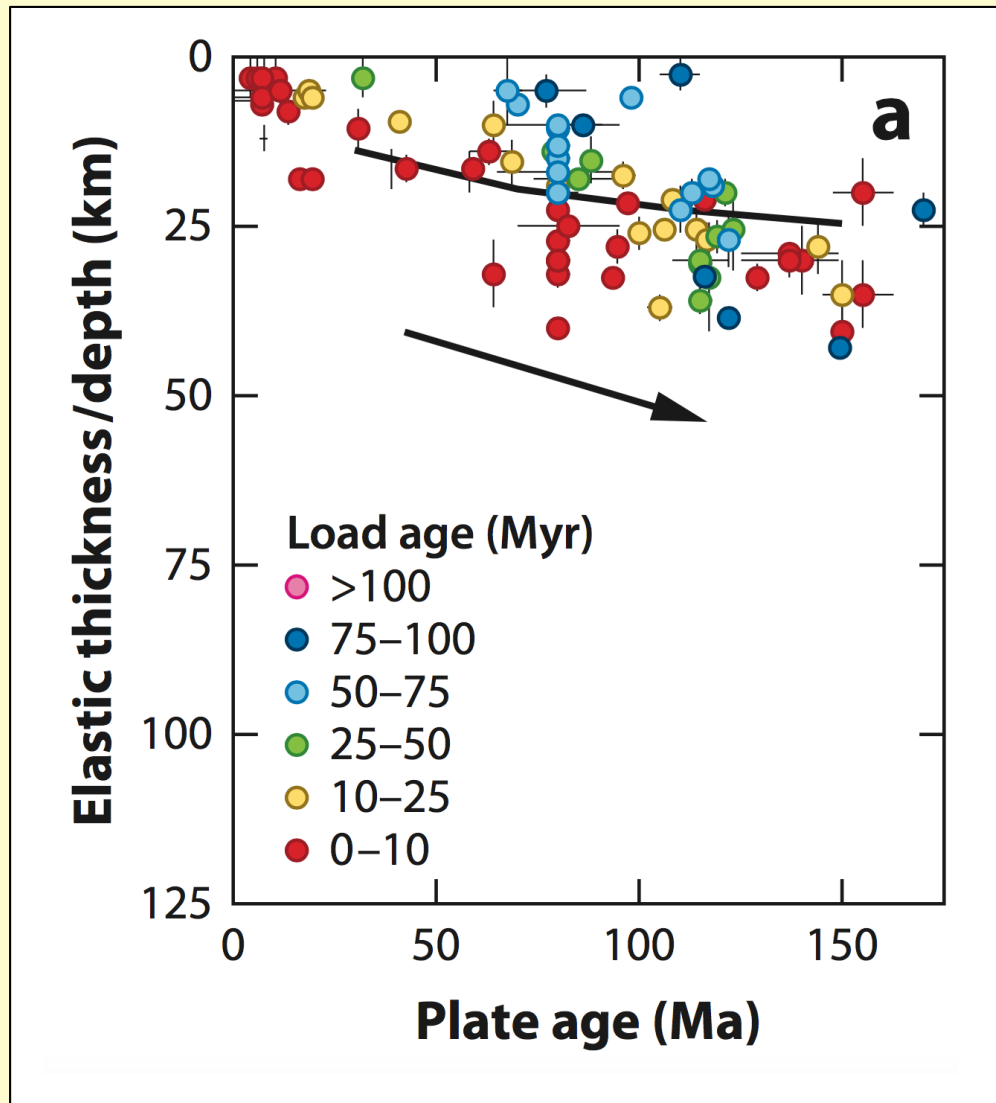
- Deflection scales
as $w_m \sim h_e^{3/4}$
- Wider deflection
for thicker elastic
plate



We can use seamounts as point loads to measure the elastic plate thickness in many places



Elastic thickness increases with plate age



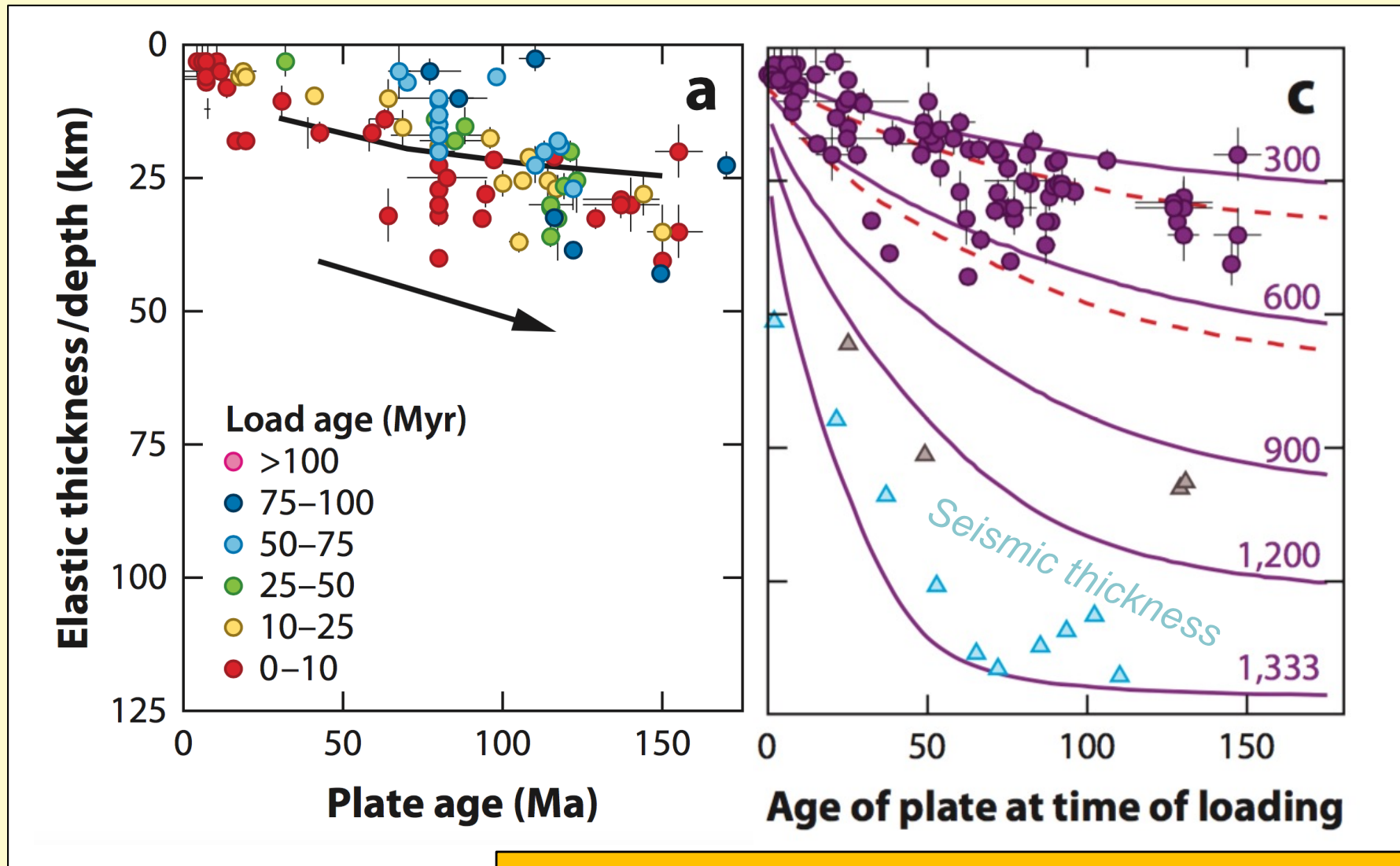
Much variation

Watts et al. [2013]

Estimate elastic thickness from the
seafloor deflection around seamounts

Elastic thickness increases with plate age

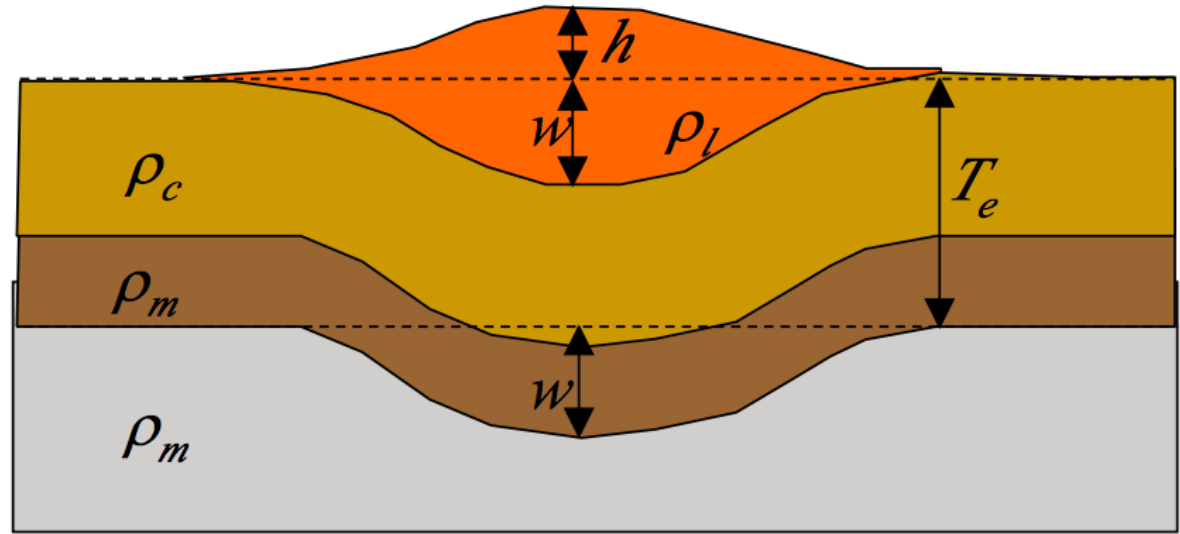
Elastic thickness follows an isotherm



Watts et al. [2013]

Seafloor deflection around seamounts suggests that temperature controls the elastic thickness:
The coldest 300-600 °C of the plate behaves elastically.

How does the lithospheric response depend on the width of the load?



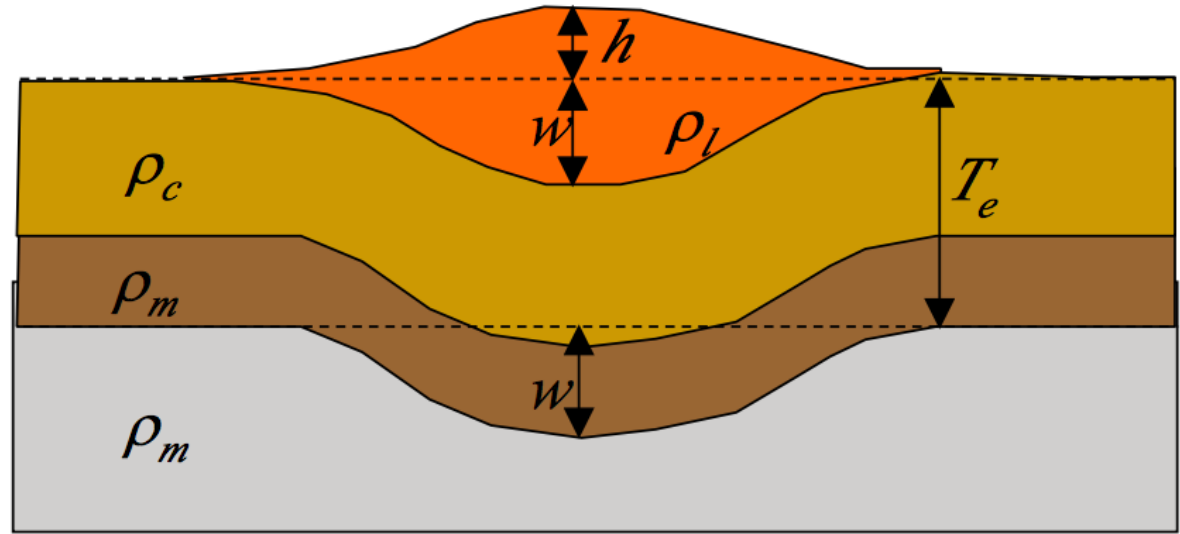
Force Balance Equation for a load on an elastic plate:

$$D \frac{d^4 w}{dx^4} + (\rho_m - \rho_l) g w = \rho_l g h$$

D is the (flexural) rigidity, T_e is the elastic thickness

$$D = \frac{ET_e^3}{12(1-\nu^2)}$$

How does the lithospheric response depend on the width of the load?



Force Balance Equation for a load on an elastic plate:

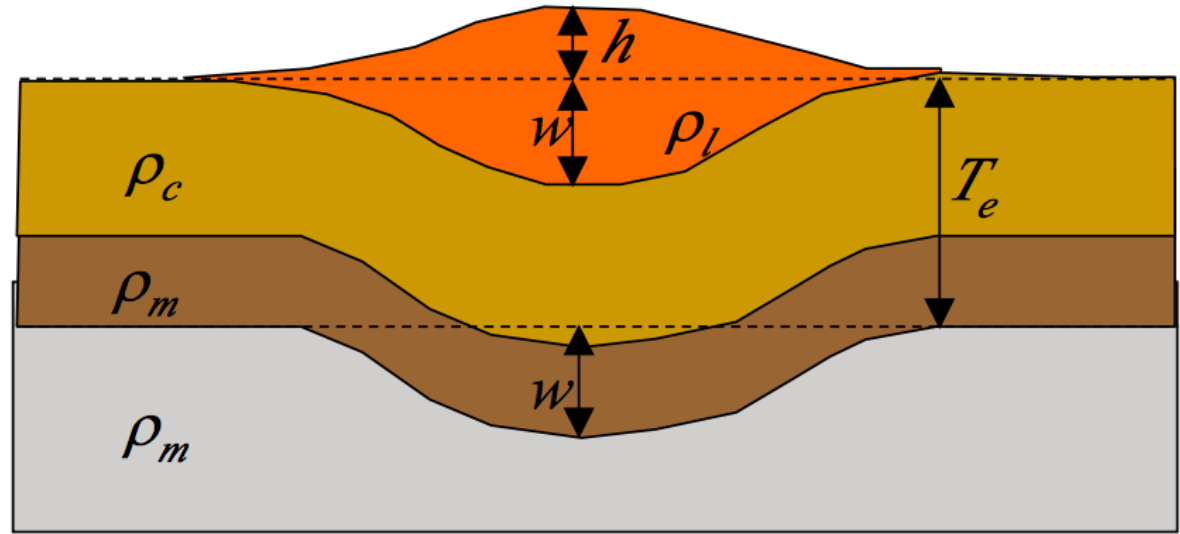
$$D \frac{d^4 w}{dx^4} + (\rho_m - \rho_l) g w = \rho_l g h$$

Assume periodic solution: $w = w_0 \sin(kx)$

$$w_0 = \frac{\rho_l}{\Delta\rho + \frac{Dk^4}{g}} h_0$$

Here $\Delta\rho = \rho_m - \rho_l$ and $k = 2\pi/\lambda$, where λ is the wavelength

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Here $\Delta\rho = \rho_m - \rho_l$ and $k = 2\pi/\lambda$, where λ is the wavelength

Short wavelengths: Elastic flexure

$$\lambda \ll \lambda_e \rightarrow w_0 \text{ is small}$$

Elastic wavelength
 $\lambda_e \sim 400 \text{ km}$ if $T_e = 25 \text{ km}$

Long wavelengths: No elastic strength

$$\lambda \gg \lambda_e \rightarrow w_0 = h_0 \rho_l / \Delta\rho \text{ (isostatic compensation)}$$

Elastic Thickness of Continental Lithosphere:

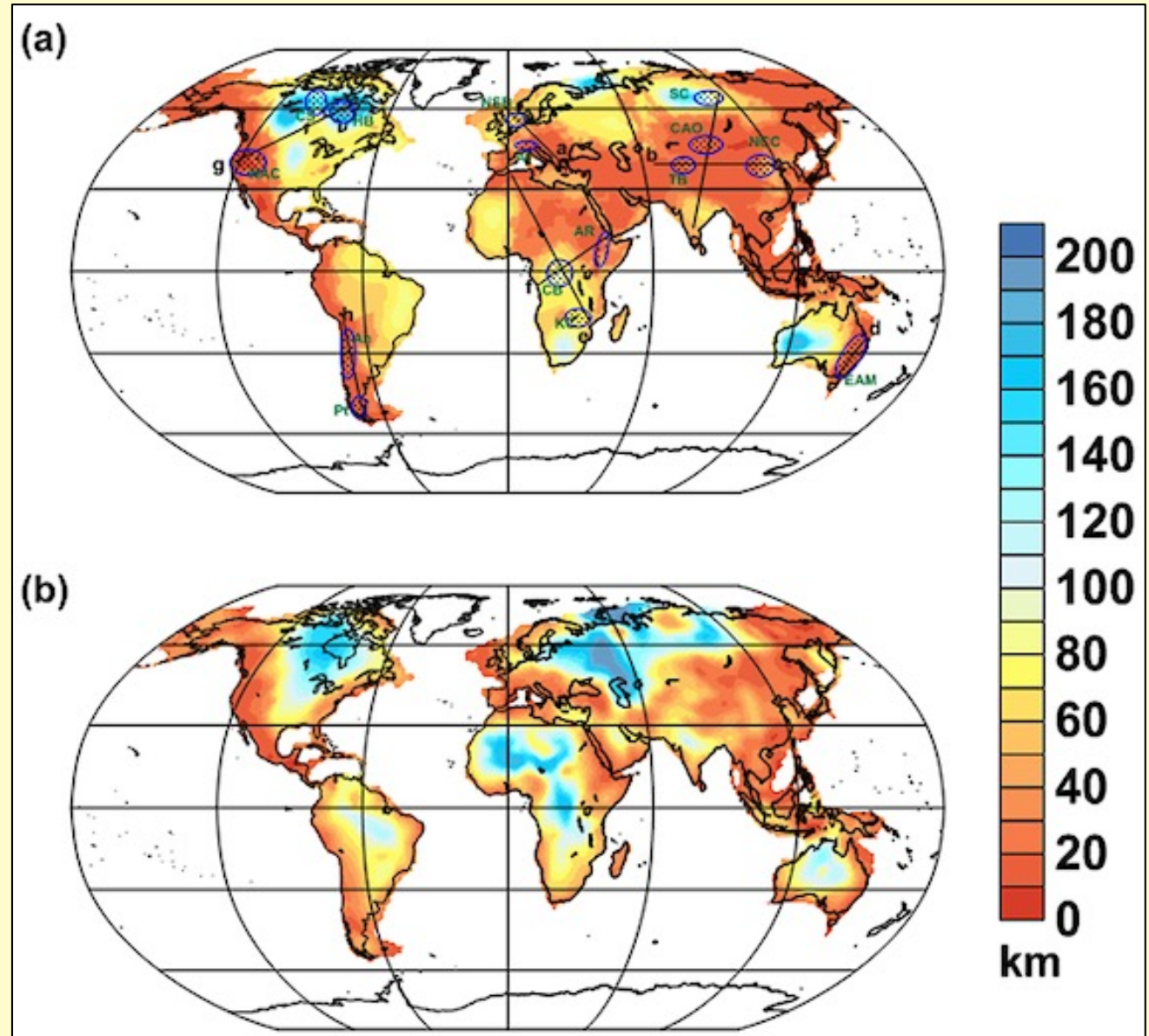
→ Strong Cratons

Elastic thickness
based on
rheology model

Elastic thickness
based on
topography to
gravity ratio

Isostatic topography
→ No gravity anomaly

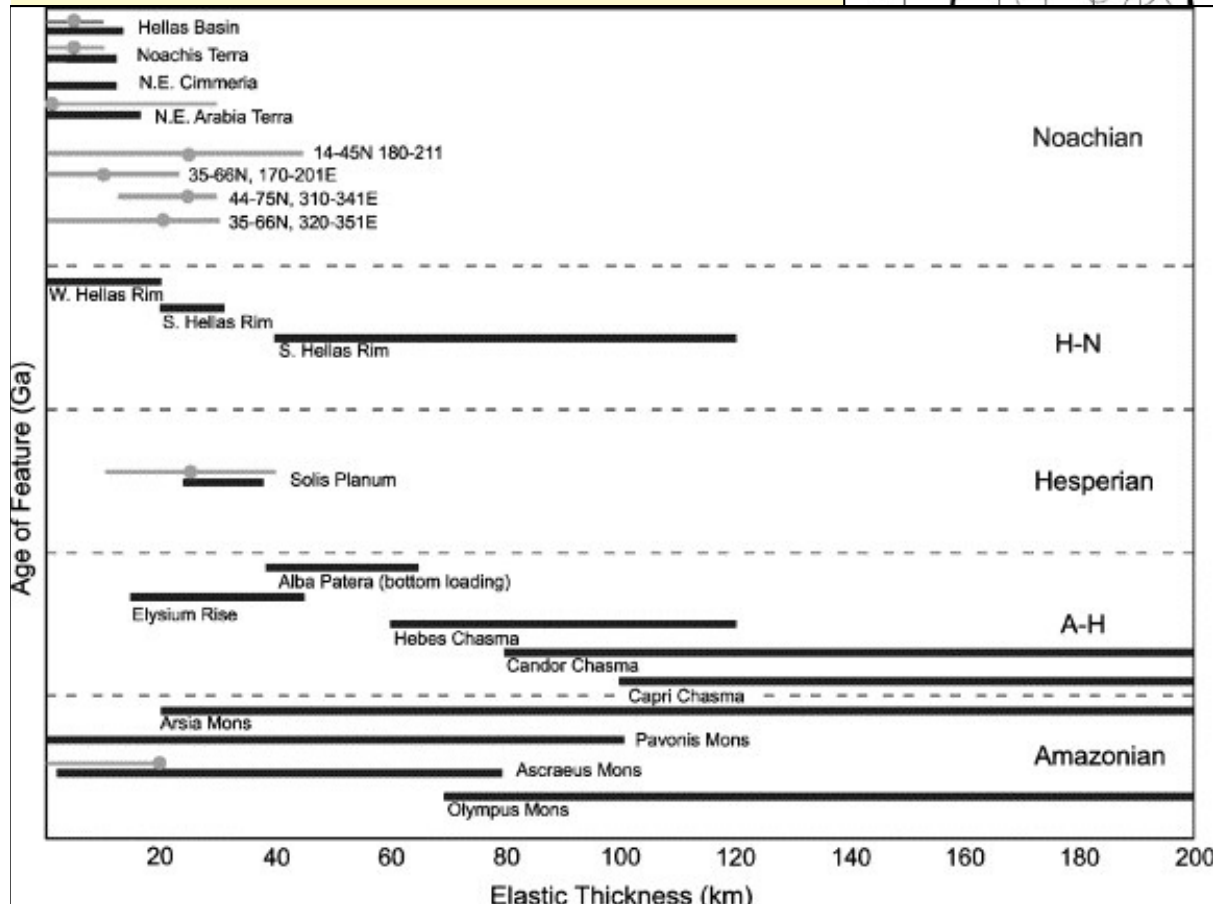
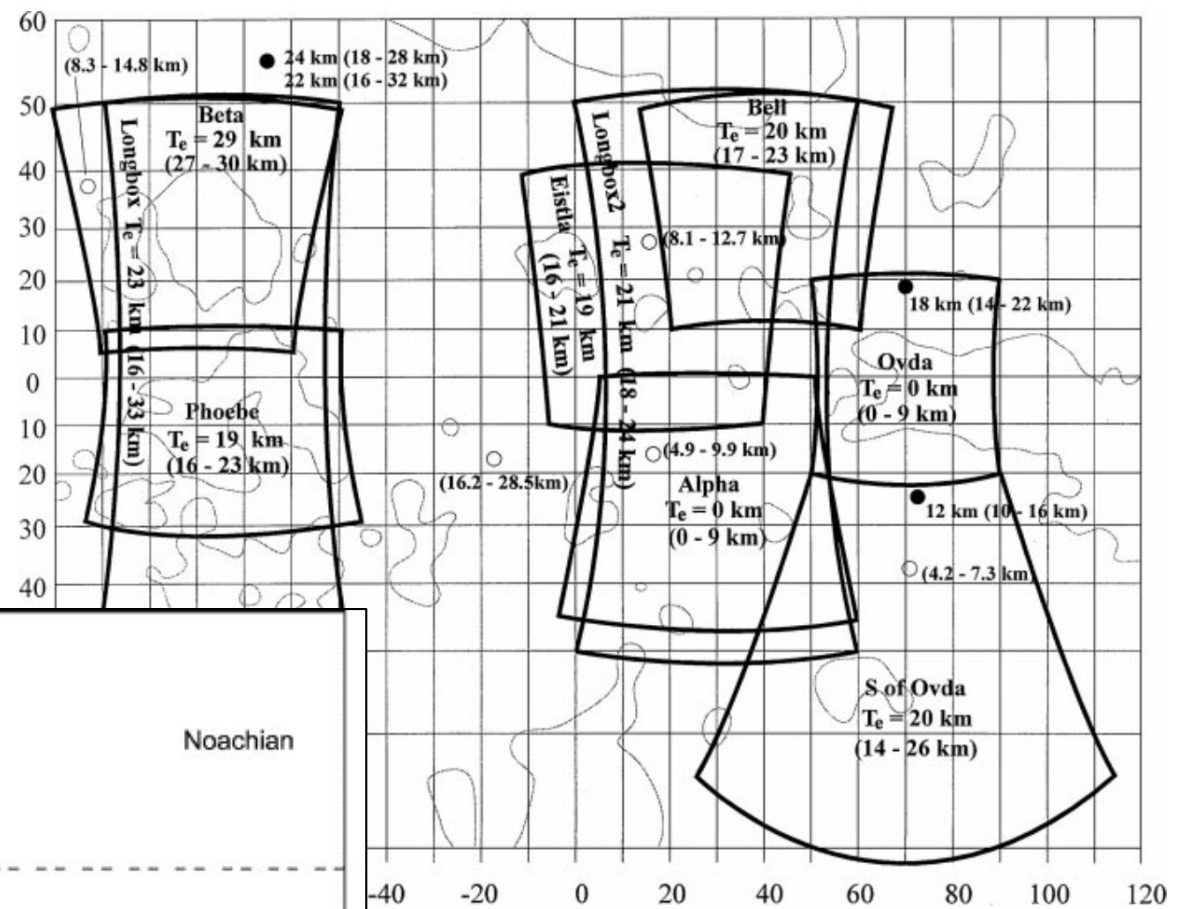
Elastic support
→ Gravity anomaly
correlates to topography



Tesauro et al. [2012]

Venus Elastic Thickness

Barnett et al. [2000]

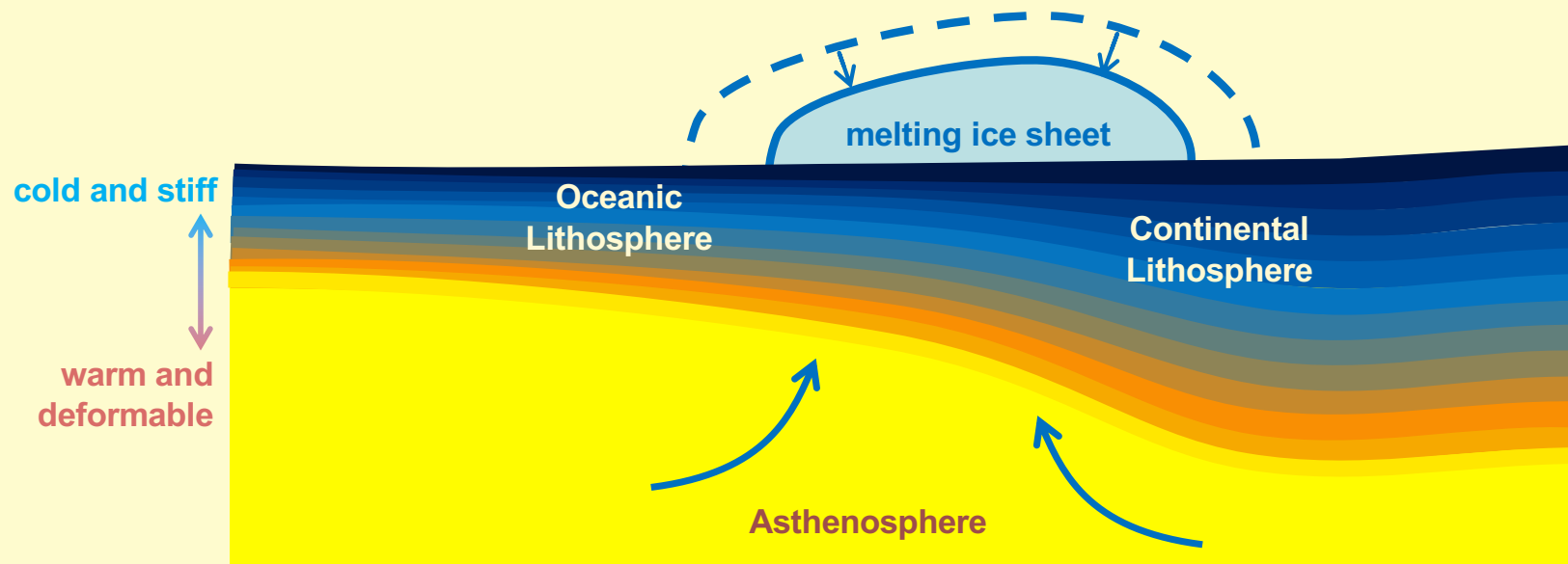


Mars Elastic Thickness

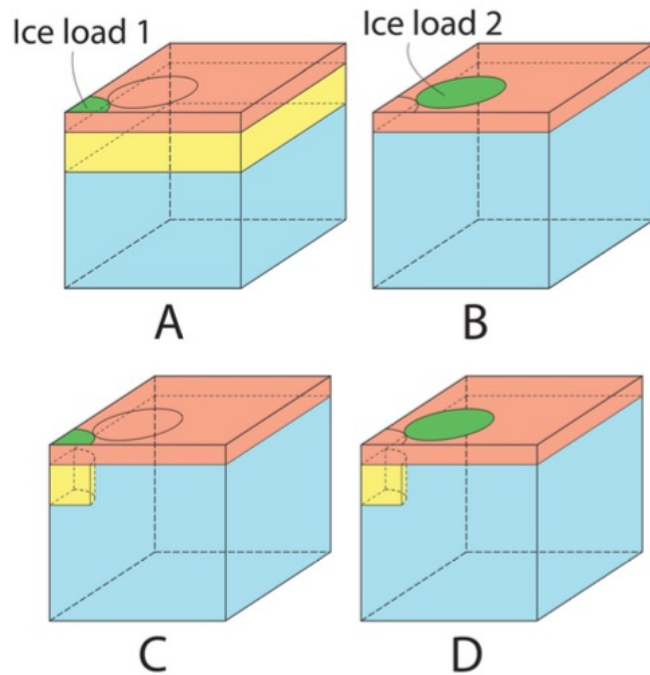
Hoogenboom & Smrekar [2006]

Glacial Isostatic Adjustment:

A time-dependent viscoelastic response to loads



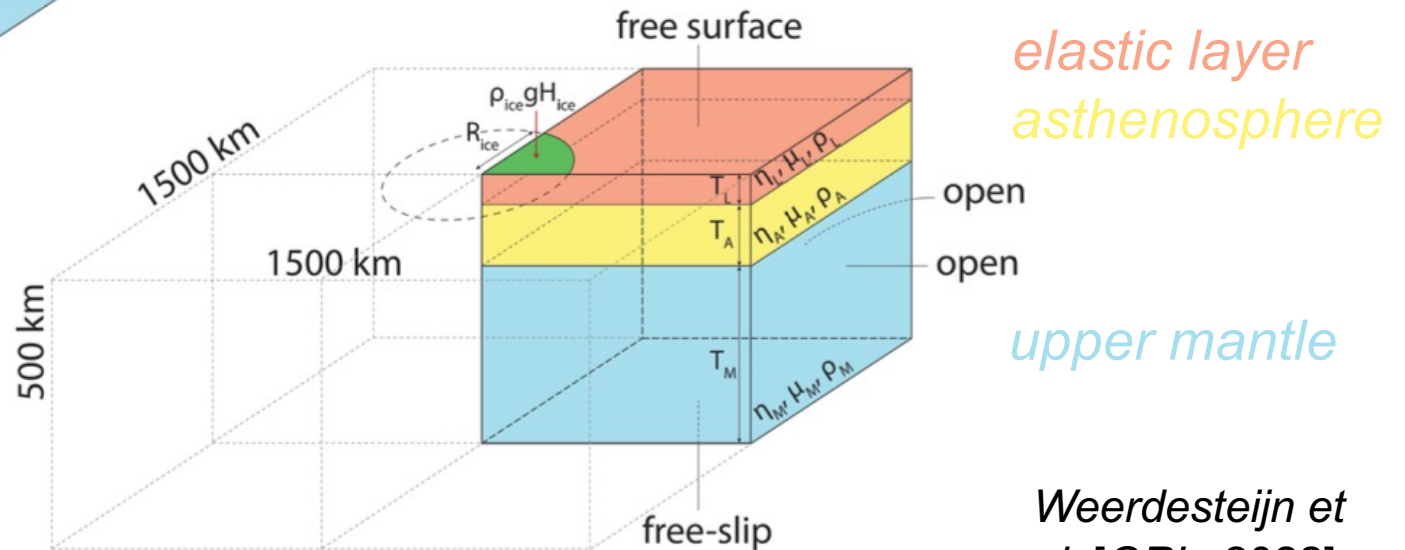
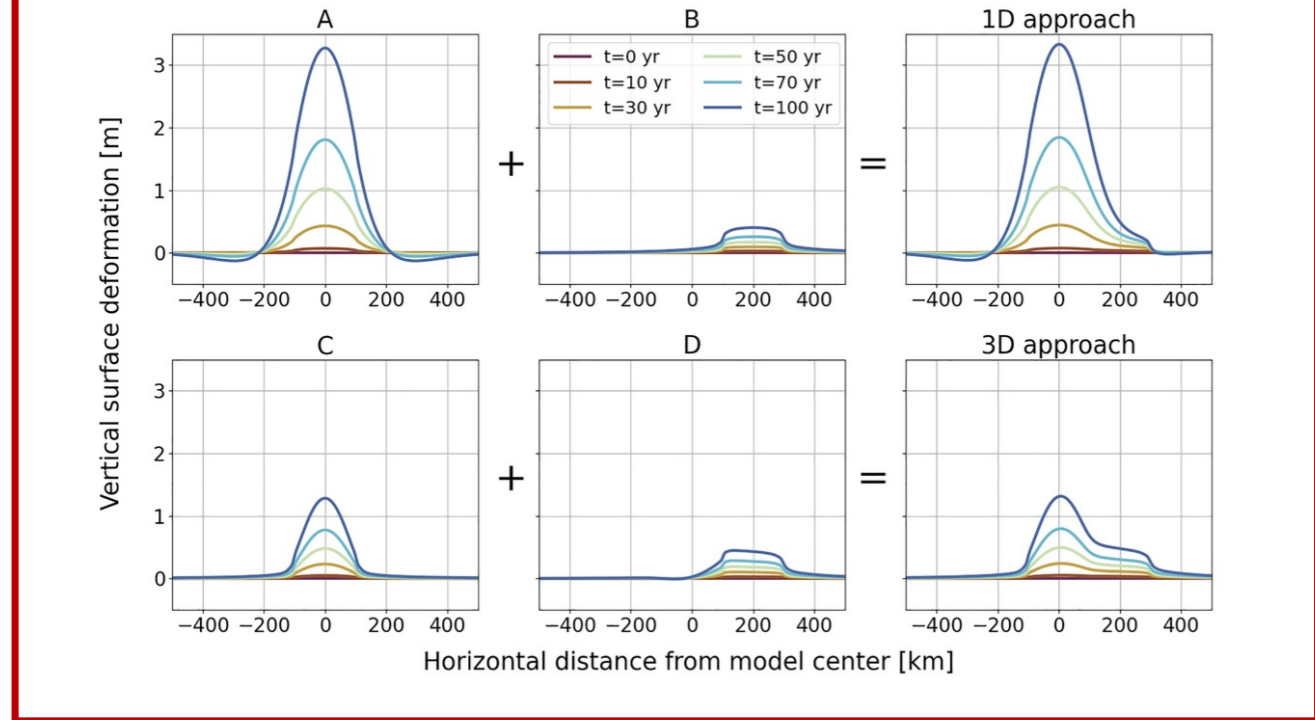
Ice load and Rheology Models



Viscoelastic Model

Load is from 1 m/yr of ice melting

Deflection vs. time for the models

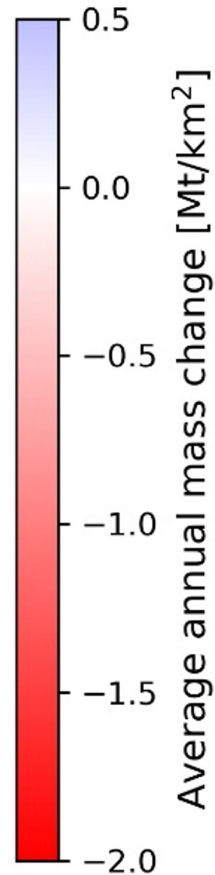
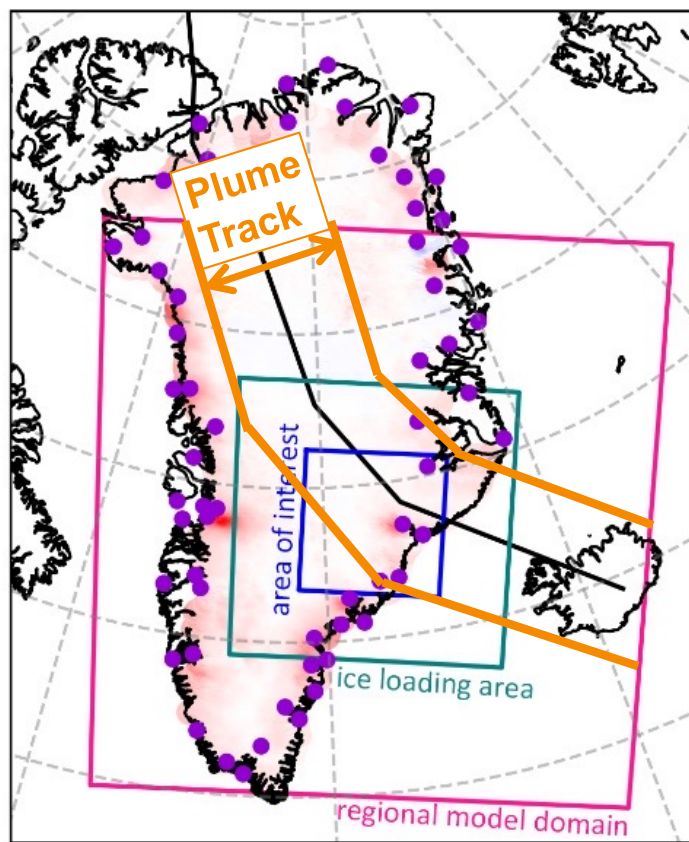


elastic layer
asthenosphere
upper mantle

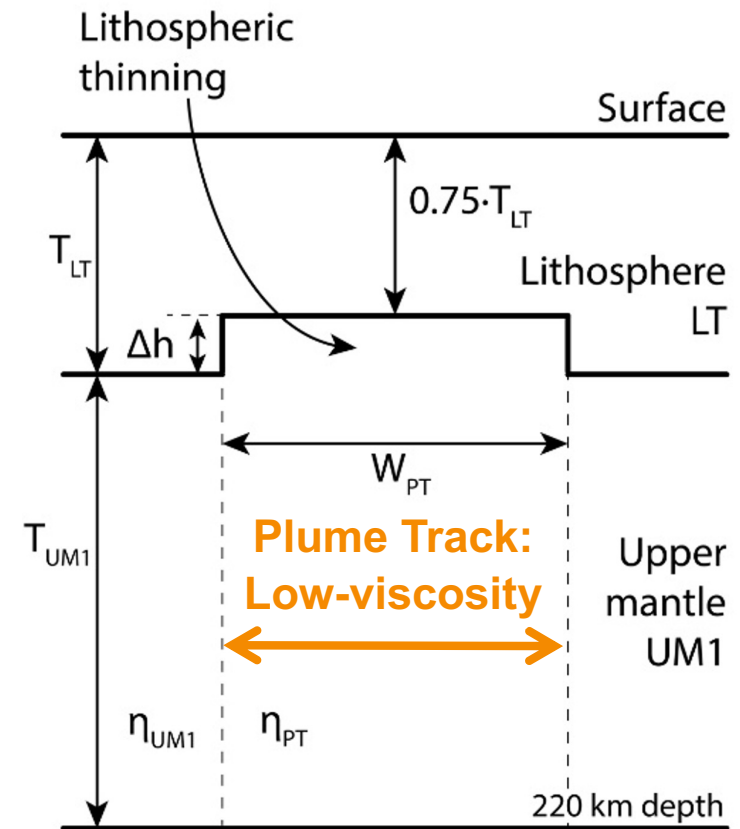
Weerdesteijn et al. [GRL, 2022]

Regional Models: Application to Greenland

Ice load inferred from satellite altimetry
[Simonsen et al., 2021]



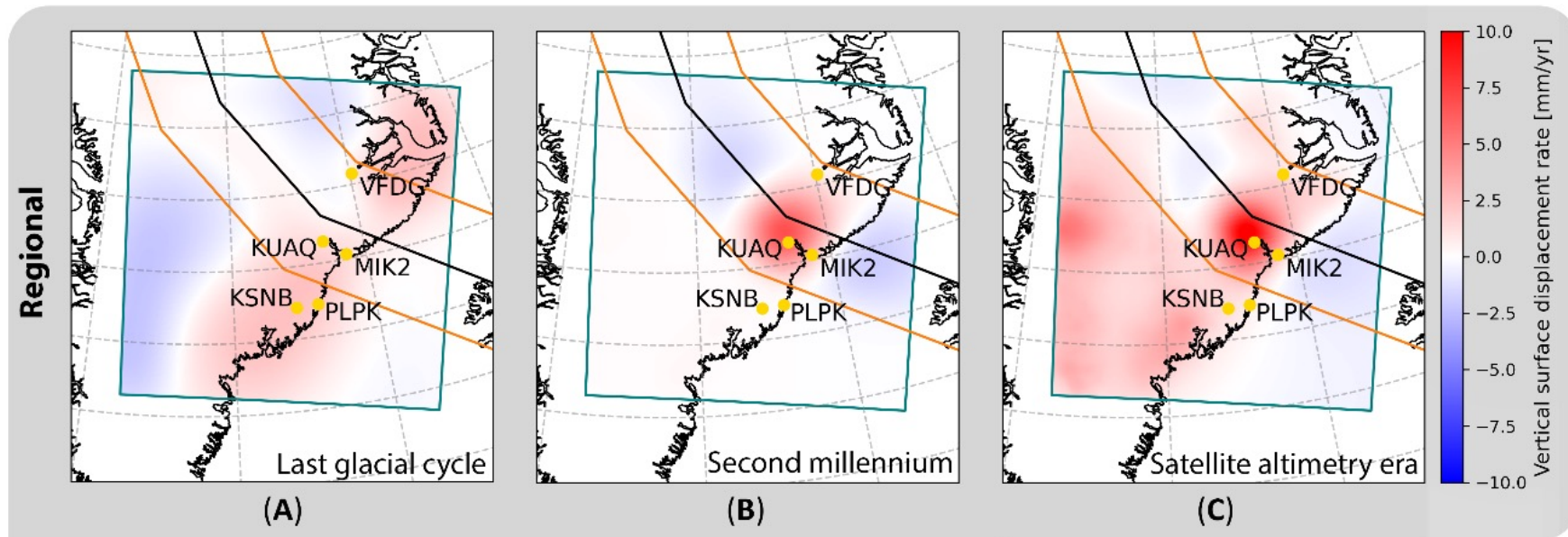
Model Profile



Weerdesteijn et al.
[submitted, 2024]

RESULTS

Regional models of GIA uplift



Last Glacial Cycle:
loads since 122 ka
[ICE-6G_C; Argus et al., 2014]

→ For past loads, the lithosphere relaxed long ago along the plume track

2nd Millennium
last 1000 years
[inferred based on Adhikari et al., 2021]

→ For recent loads, the lithosphere is actively adjusting

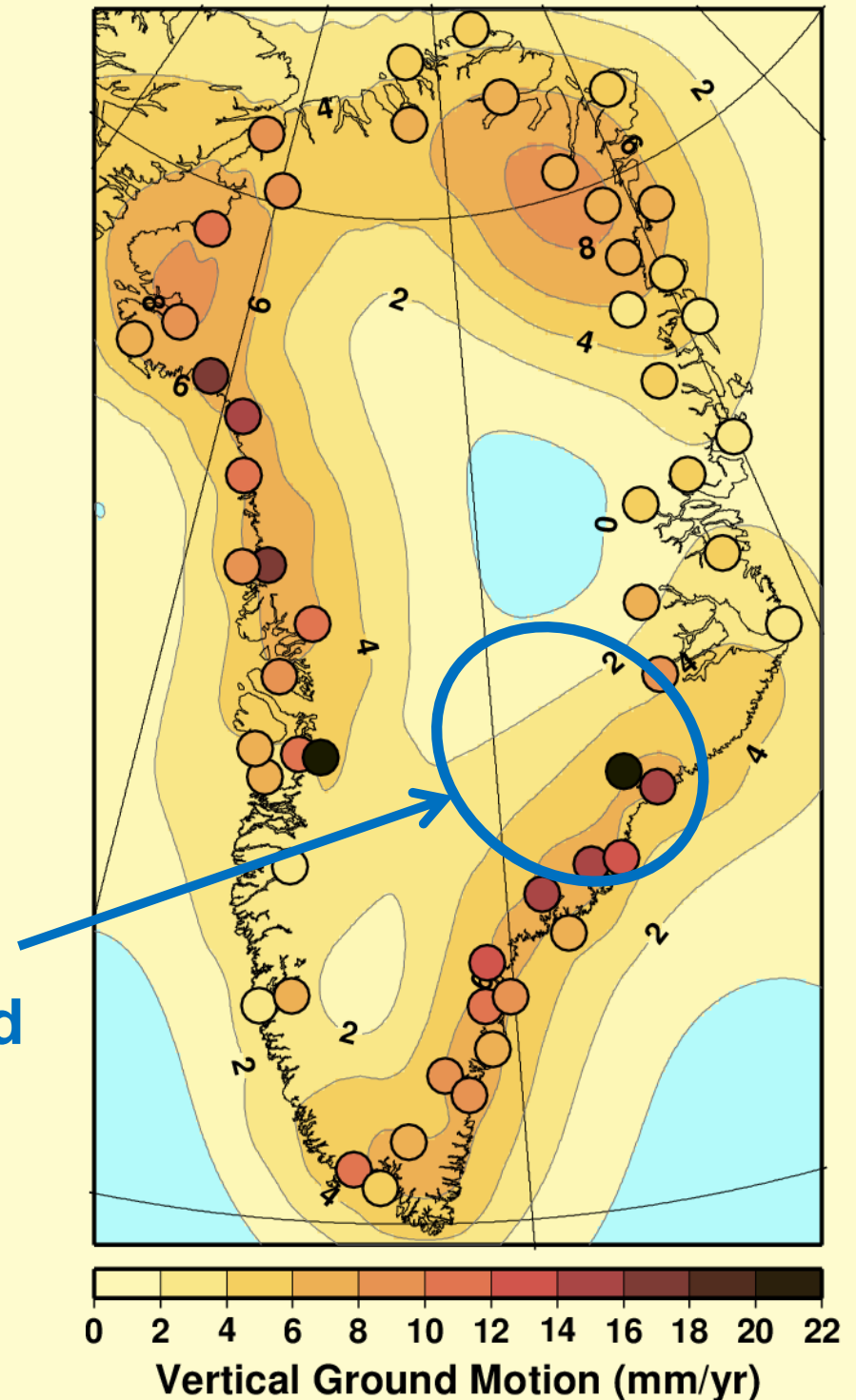
Satellite Altimetry Era:
loads 1990-2020
[Simonsen et al., 2014]

Rapid uplift along coastal Greenland

- Measured by GPS [e.g. Khan et al., 2016]
- Some stations are rising faster than expected

Is Southeast Greenland rising faster because of heat left behind by the Iceland Plume?

- Thinner elastic lithosphere
- Low viscosity asthenosphere



Conclusions: The Elastic Lithosphere

- The top (cold) part of the lithosphere behaves elastically
- Elastic stresses can support loads up to ~400 km wide
- Lithosphere flexure depends on elastic thickness.
- Elastic thickness depends on temperature and history.
- The viscoelastic response to a load is faster for hotter and thinner lithosphere.

