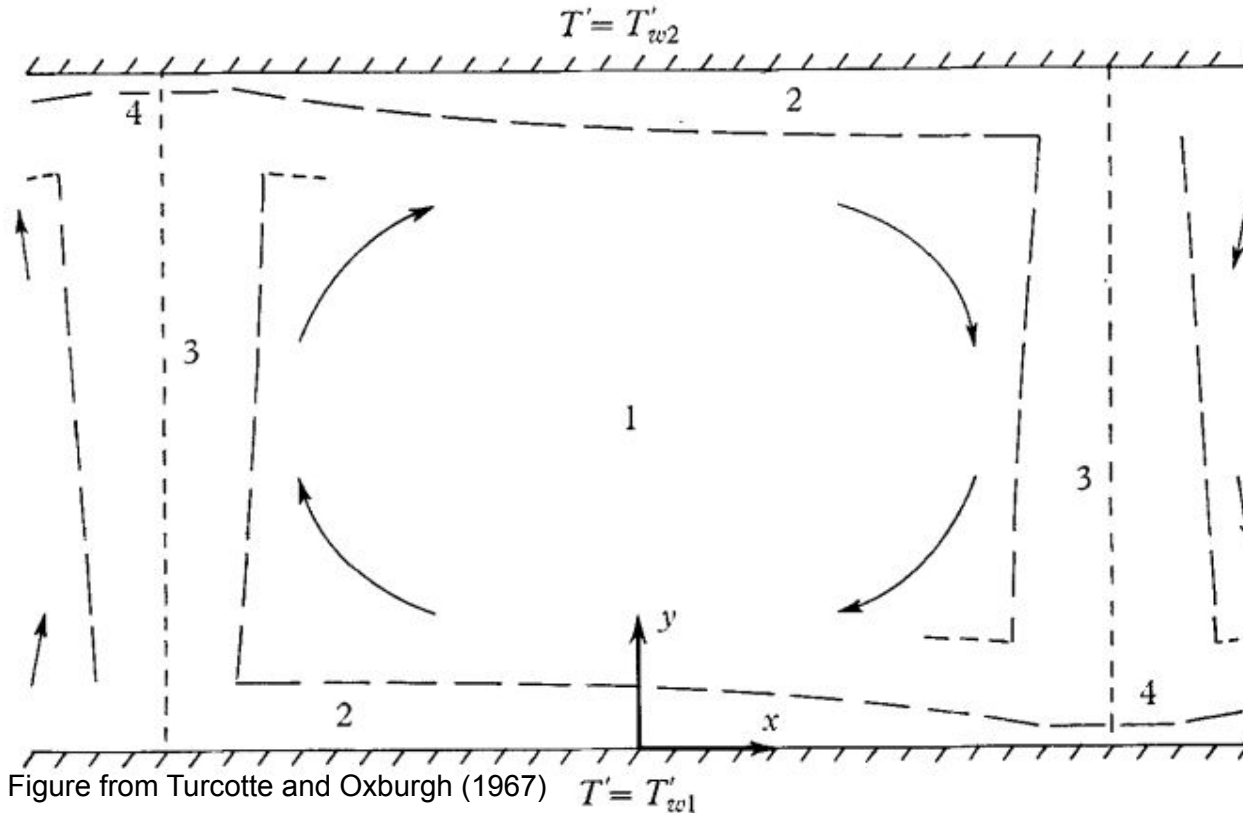


# Thermal modelling of the lithosphere and asthenosphere

Yijun Wang

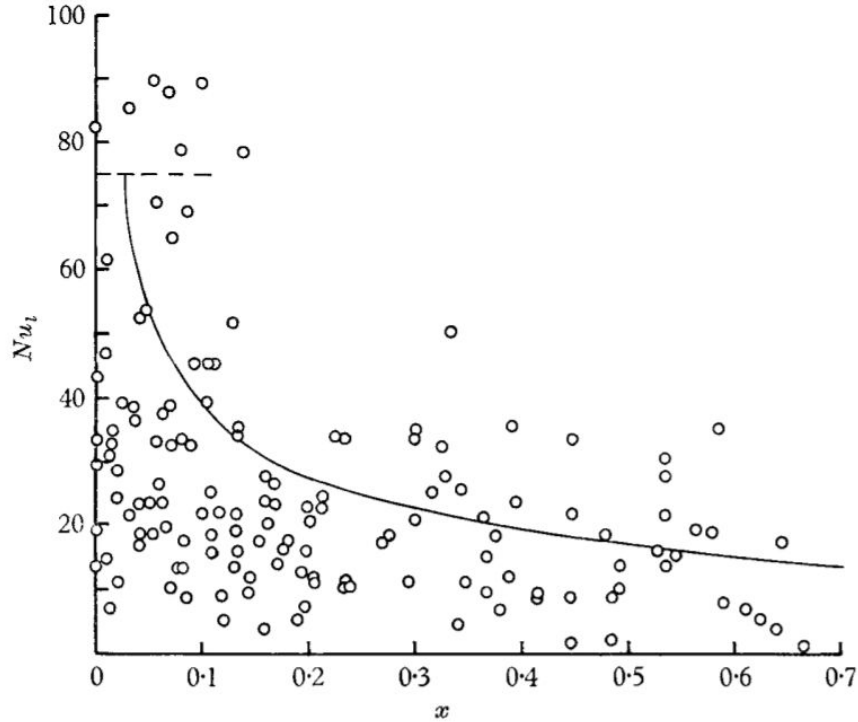
# Turcotte and Oxburgh, 1967



- Lithosphere as a conductive thermal boundary layer above a convective mantle

Figure from Turcotte and Oxburgh (1967)  $T' = T'_{w1}$

# Turcotte and Oxburgh, 1967

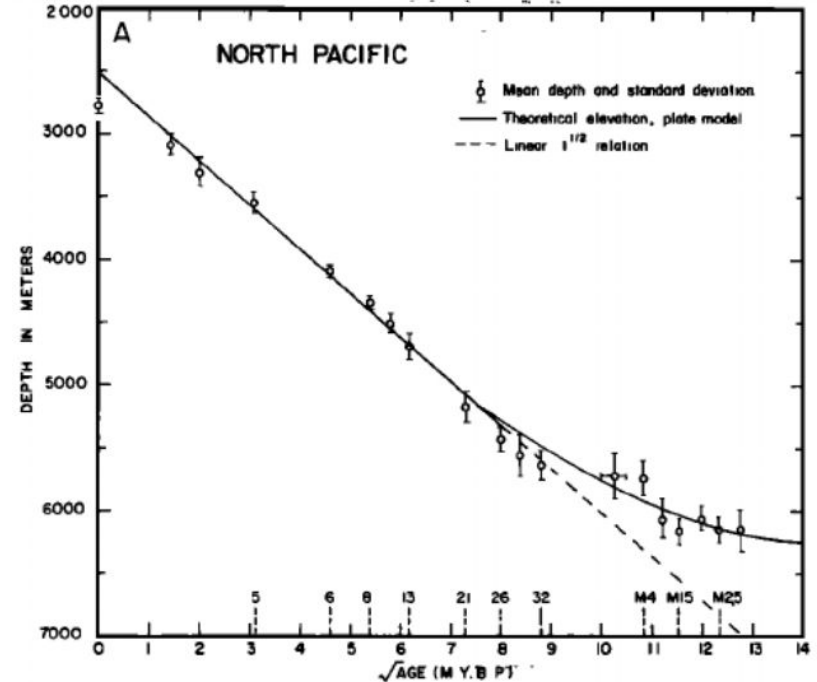
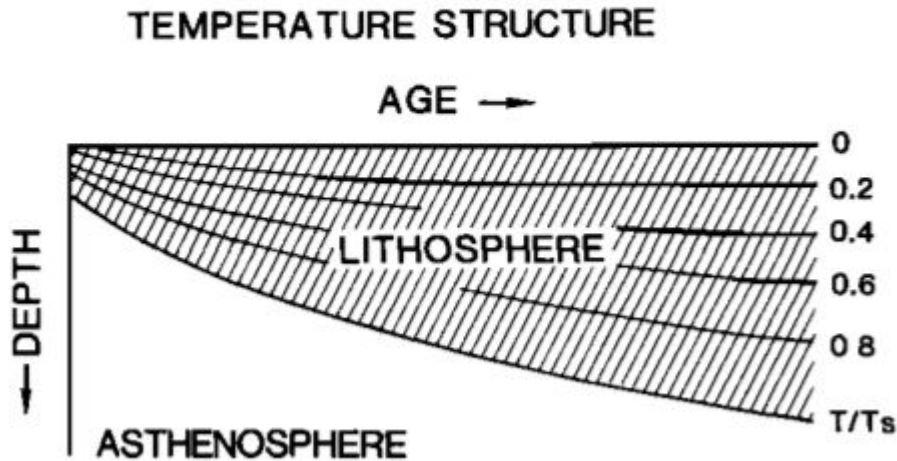


- Conservation equations
- Nusselt number
- East Pacific Rise from Lee and Uyeda (1965)
- Calculate velocity of continental drift

Figure from Turcotte and Oxburgh (1967)

# Simple conductive cooling model

## a. SIMPLE COOLING ' $\sqrt{t}$ '



Figures (left) from Renkin and Sclater (1988) and (right) from Parsons and Sclater (1977)

# McKenzie, 1967

- Accounts for flattening of oceanic seafloor
- Constant temperature at the bottom and constant thickness of the layer
- Knows thermal properties of the layer, calculate the temperature and depth-age relationship

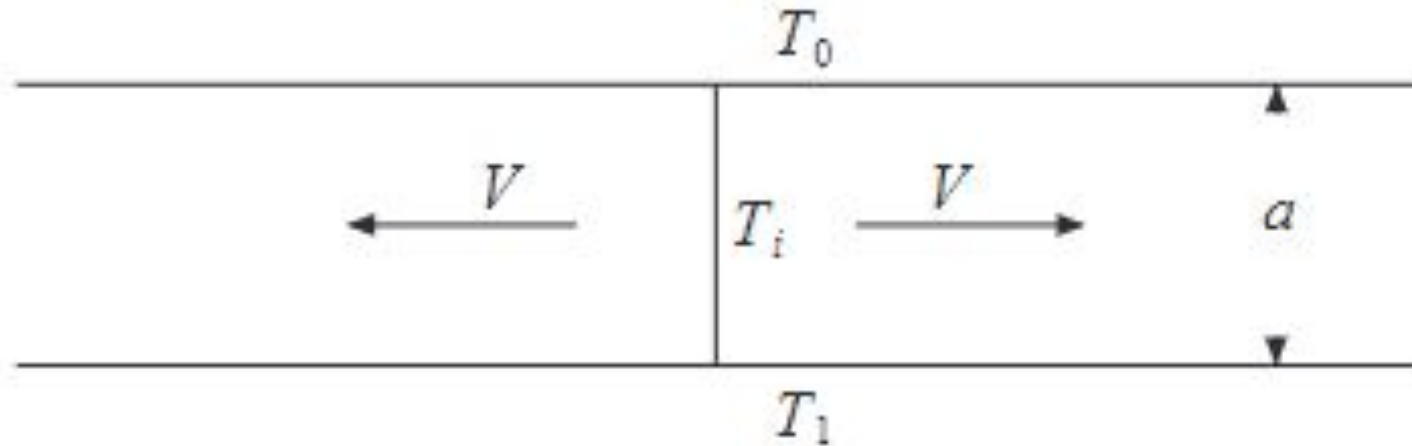


Figure from McKenzie et al. (2005)

## Governing equations

$$\rho C_P \left[ \frac{\partial T}{\partial t} + \mathbf{v} \cdot \nabla T \right] = \kappa \nabla^2 T + H$$

$$\nabla \cdot \mathbf{u} = 0$$

$$\frac{\mu}{\rho} \nabla^2 \mathbf{u} - g \mathbf{a}_z - \frac{1}{\rho} \nabla P = 0$$

# Plate model

- Parsons and McKenzie, 1978
- Convection that transport heat to the bottom of the mechanical layer

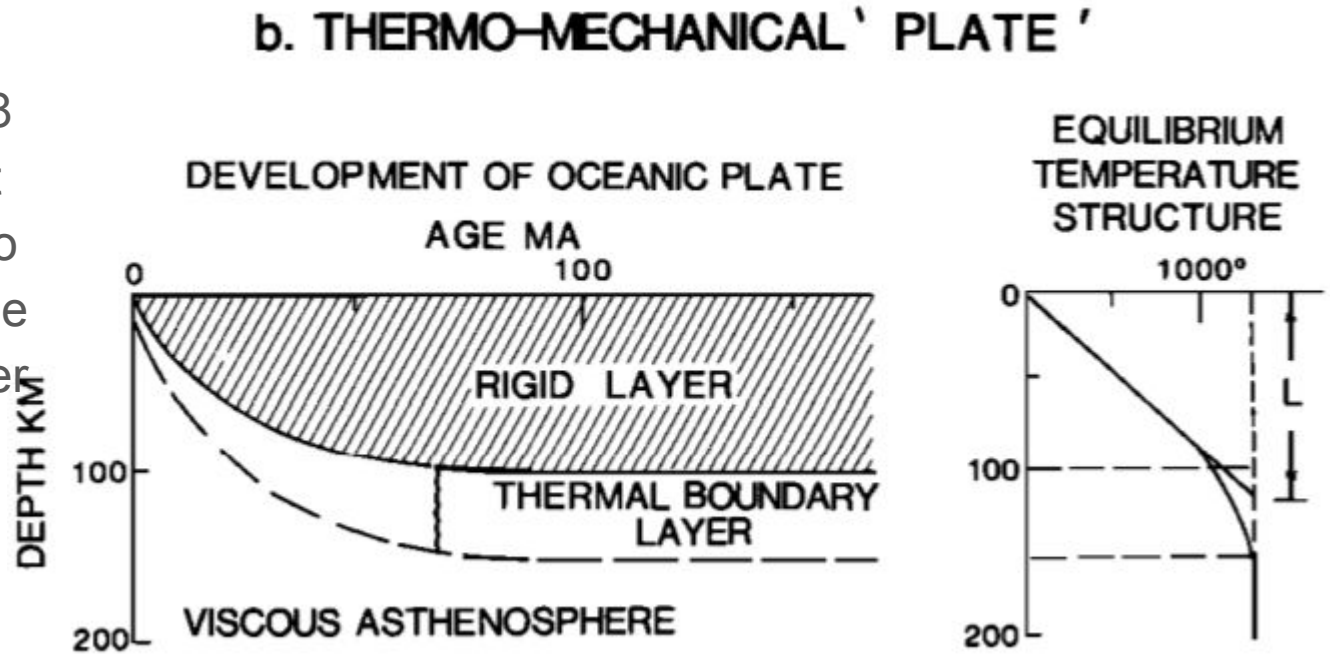
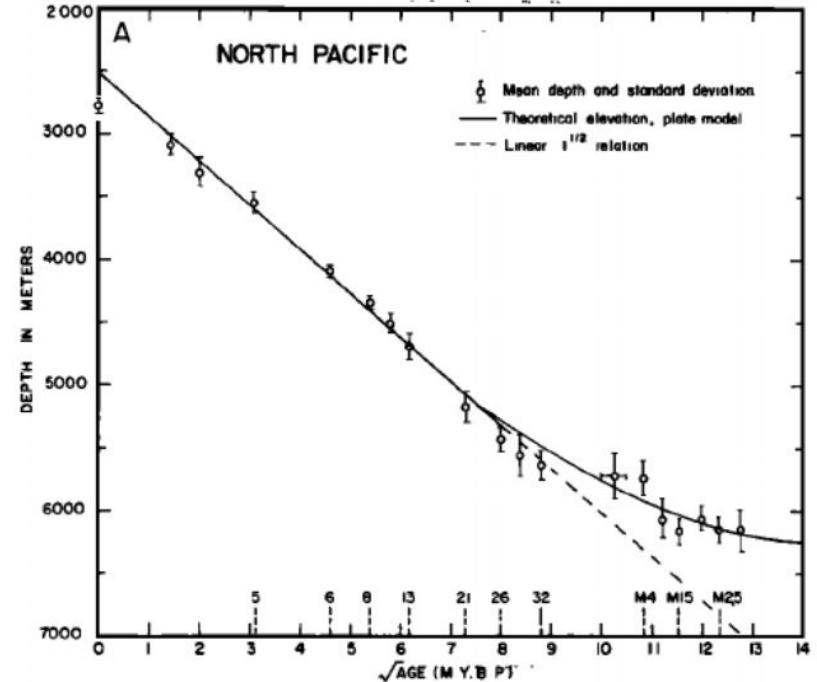
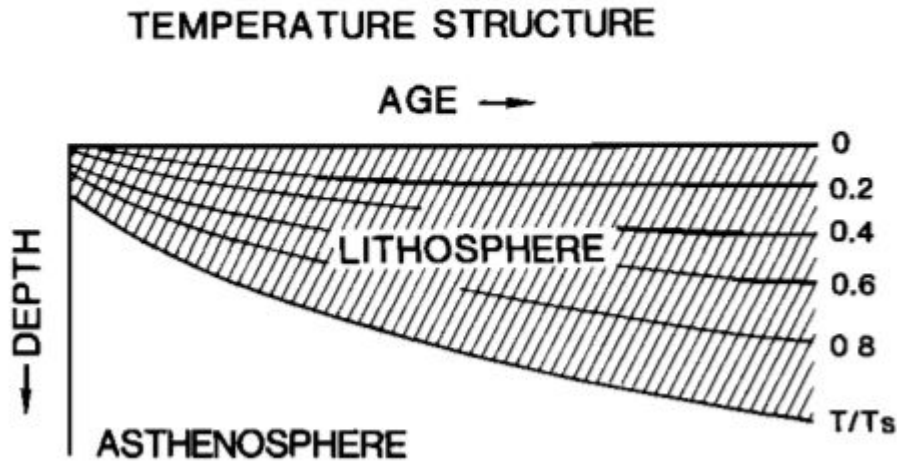


Figure from Renkin and Sclater (1988)

# Simple conductive cooling model

## a. SIMPLE COOLING ' $\sqrt{t}$ '



Figures (left) from Renkin and Sclater (1988) and (right) from Parsons and Sclater (1977)



# Hot spot as a mechanism that transport heat

- Studies with hot spot in different oceans
  - Heestand and Crough (1981)
  - Hayes (1988)
- Different filtering methods used to exclude seafloor influenced by volcanism obtain different depth-age relationship

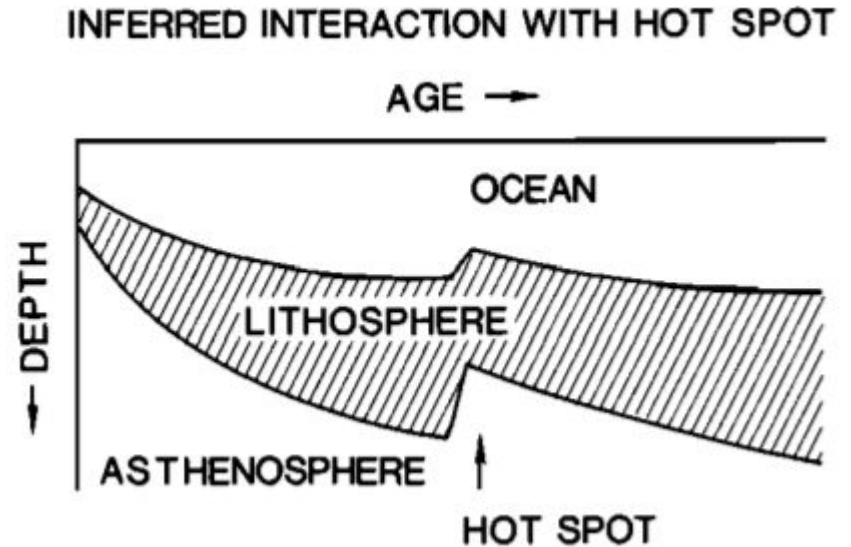


Figure from Renkin and Sclater (1988)

# Hot spots

- Hillier and Watts (2005)
  - By eye
- Crosby et al. (2006)
  - Absence of gravity anomaly

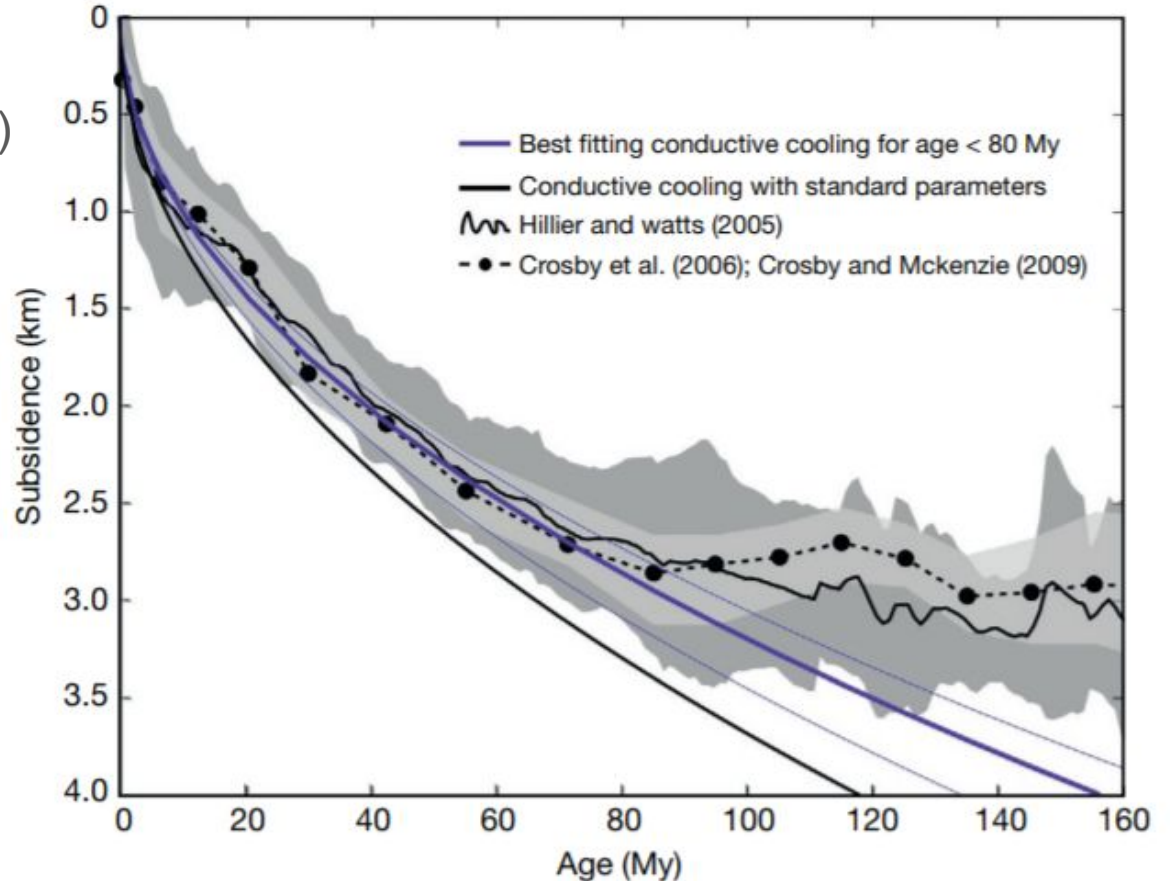


Figure from Parmentier et al. (2015)

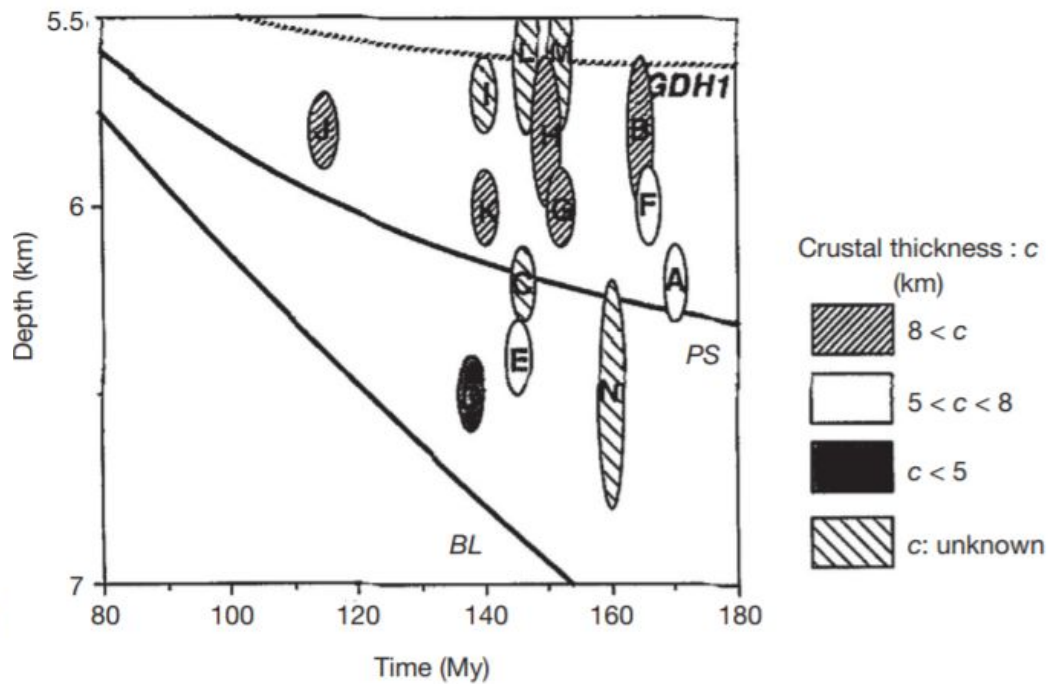
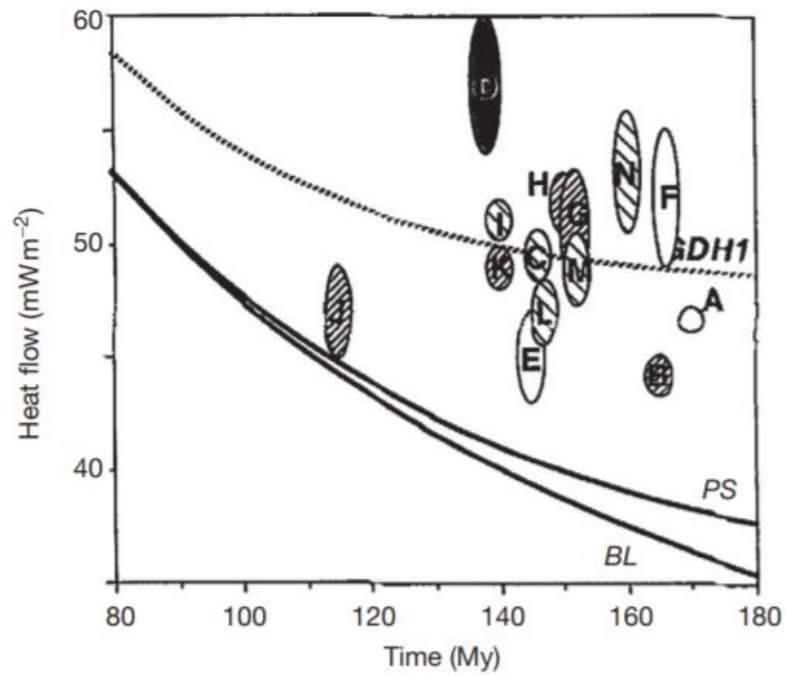
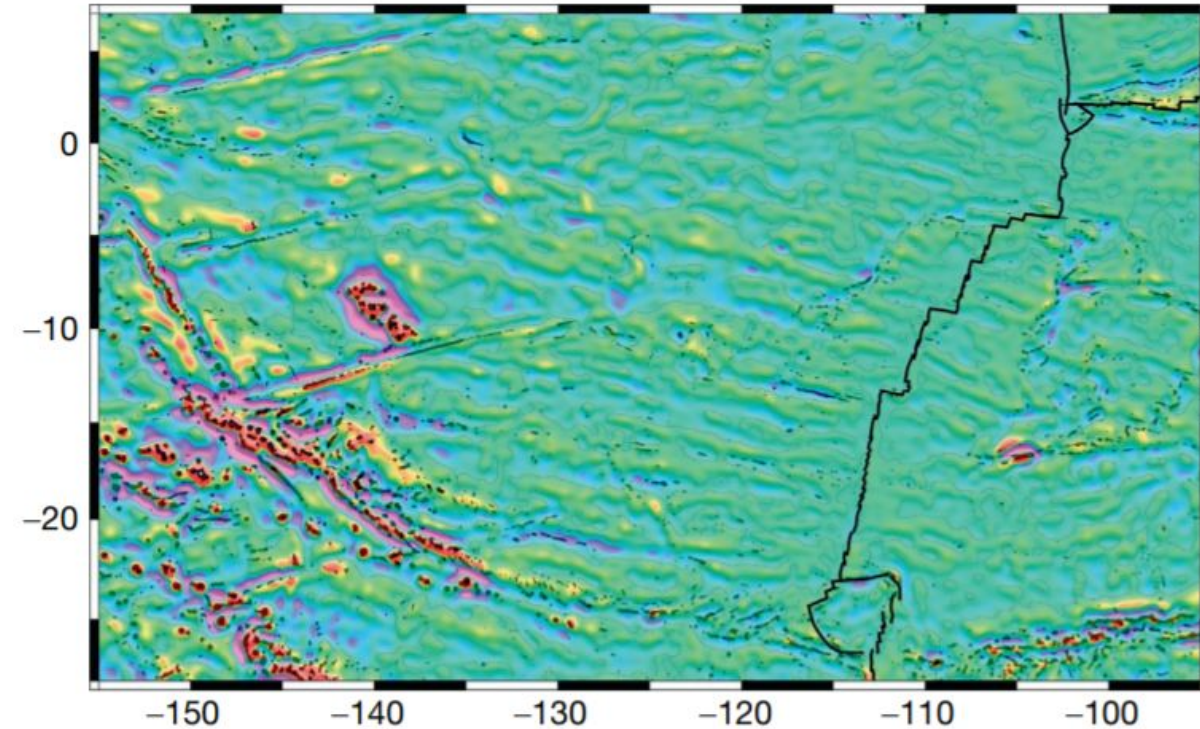


Figure from Nagihara et al. (1996)

# Small scale convective instability



- Haxby and Weissel, (1986), Pacific
- Onset time

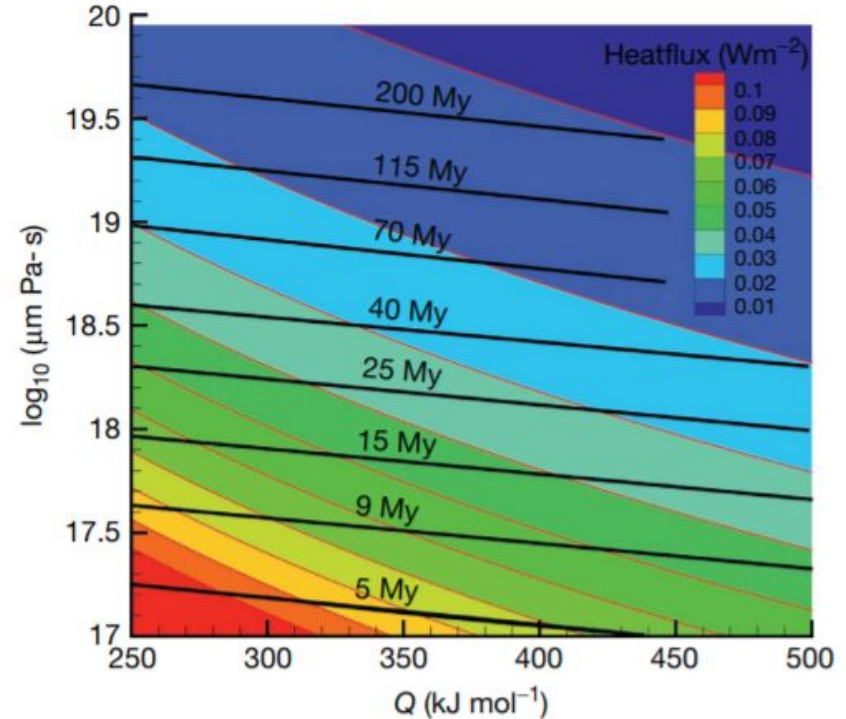
Figure from Sandwell and Fiaiko (2004)

## Doin and Fleitout, 1996

- a uniform heat flux supplied to the bottom of the thermal boundary layer explains depth–age relationship well

# Davaille and Jaupart, 1994

- scaling laws for convective onset time (age) and heat flow in a viscous fluid with strongly temperature-dependent viscosity cooled from above
- $Q = 250 \text{ kJ/mol}$ ,  $T = 1300 \text{ C}$ , predicts a onset time of 52-65 Ma.
- Numerical models predicts shorter onset times.



# Onset time of small-scale convective instability

- Spreading rate
- Sandwell and Schubert (1980) in Atlantic and SE Indian ocean
- High shear strain produces smaller grain size which is in favor of fast diffusion creep with lower viscosity or higher stress in favor of dislocation creep with large  $Q$
- Water content

## 2D vs. 3D models

- van Hunen et al. (2003) studied 3D convective instability beneath a moving plate



# Summary

- Parmentier, E. M. (2015), 7.08 - The Dynamics and Convective Evolution of the Upper Mantle A2 - Schubert, Gerald, in Treatise on Geophysics (Second Edition), edited, pp. 319-337, Elsevier, Oxford, doi:10.1016/B978-0-444-53802-4.00131-7.
- There isn't one mechanism of heat transfer that can account for all observations related to heating the lithosphere everywhere on Earth.
- The relative importance of hot spot and convective instability depends on spreading rate and mantle rheology.